

Scaling of seismic and aseismic moments of natural and induced earthquakes

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Paul Selvadurai², Stefan Wiemer²

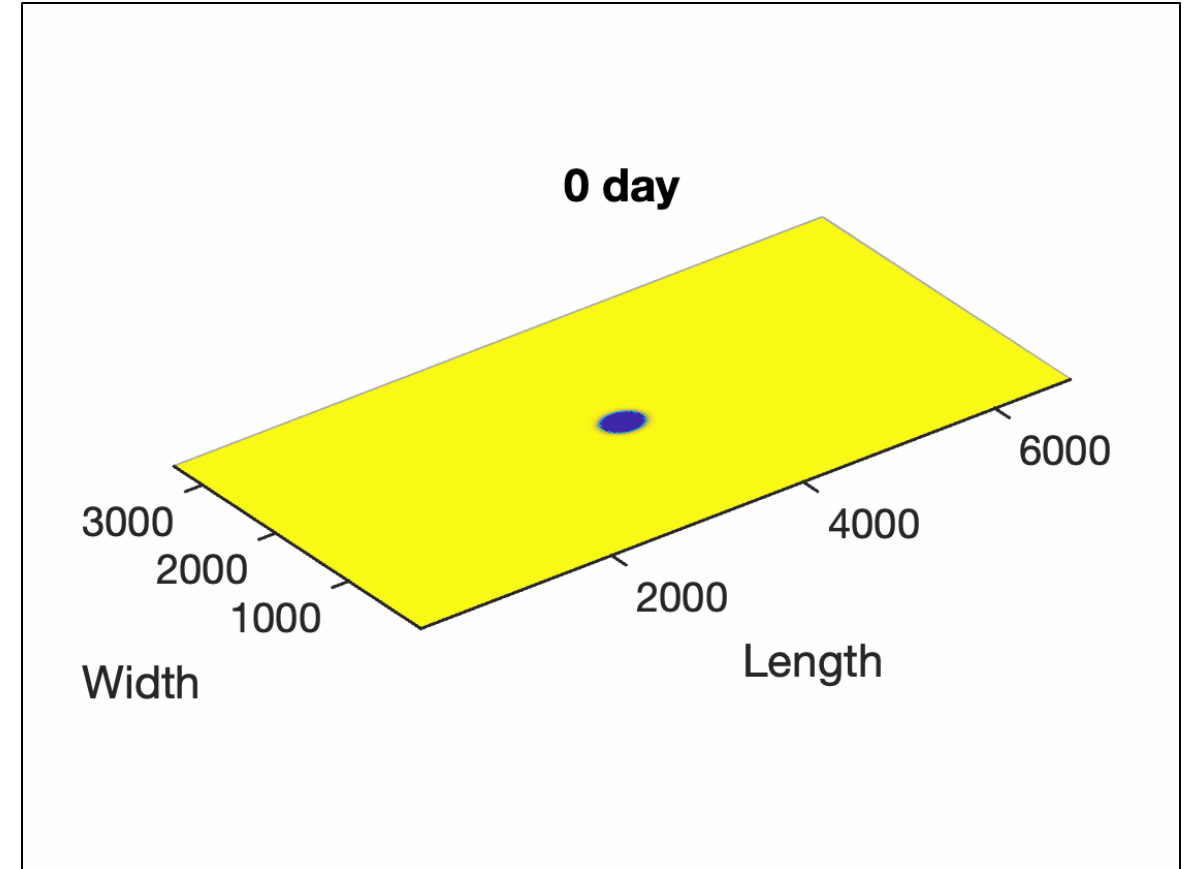
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- Database of slow slip transients + triggered seismicity:
 - Geodetic moments
 - Seismic moments of triggered seismicity
 - Duration
 - Source dimensions
- Source parameters of triggered seismicity **only during slow slip transient**

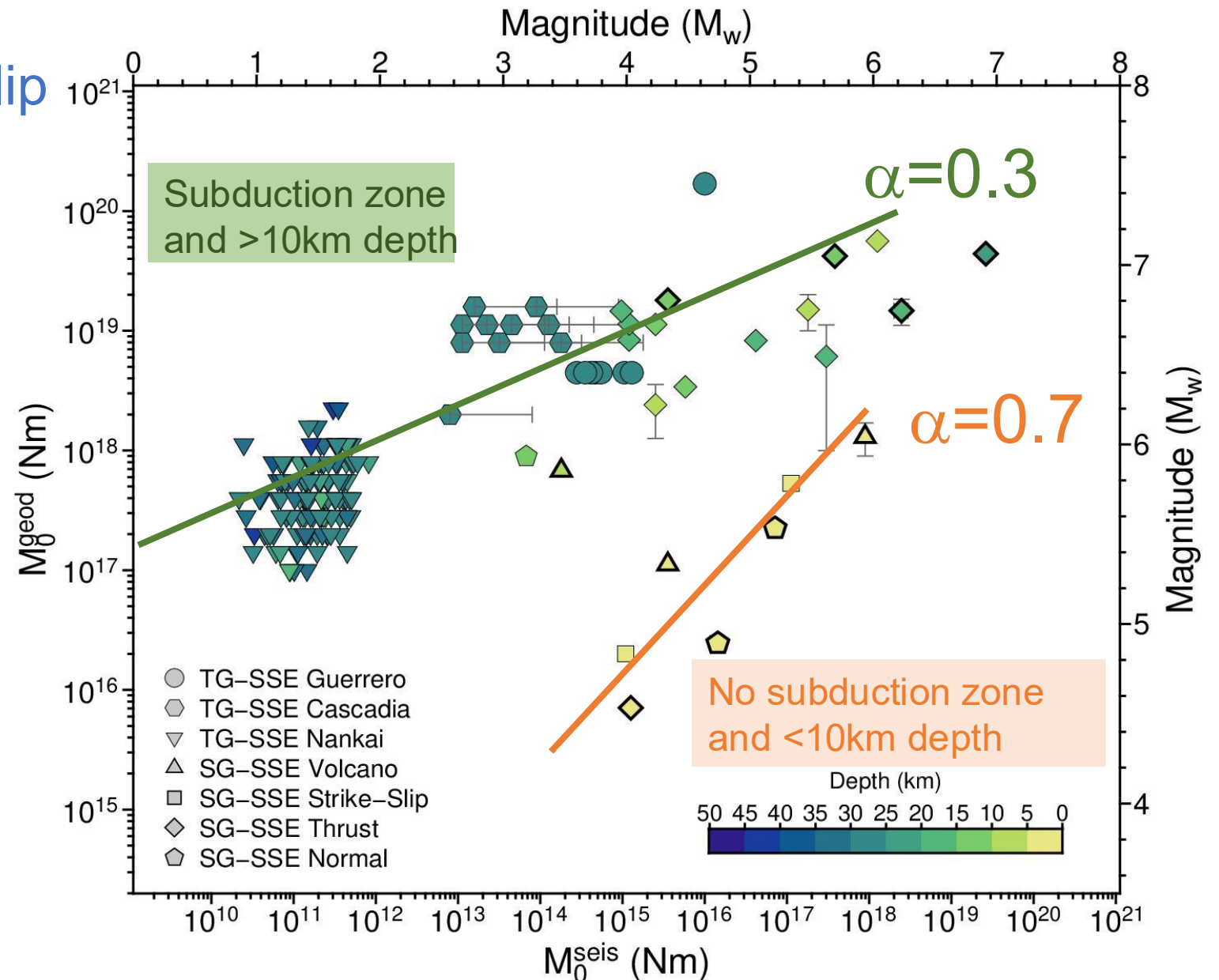


Aseismic and seismic moment release **do scale** during slow slip transients

$$M_0^{geod} = \beta (M_0^{seis})^\alpha$$

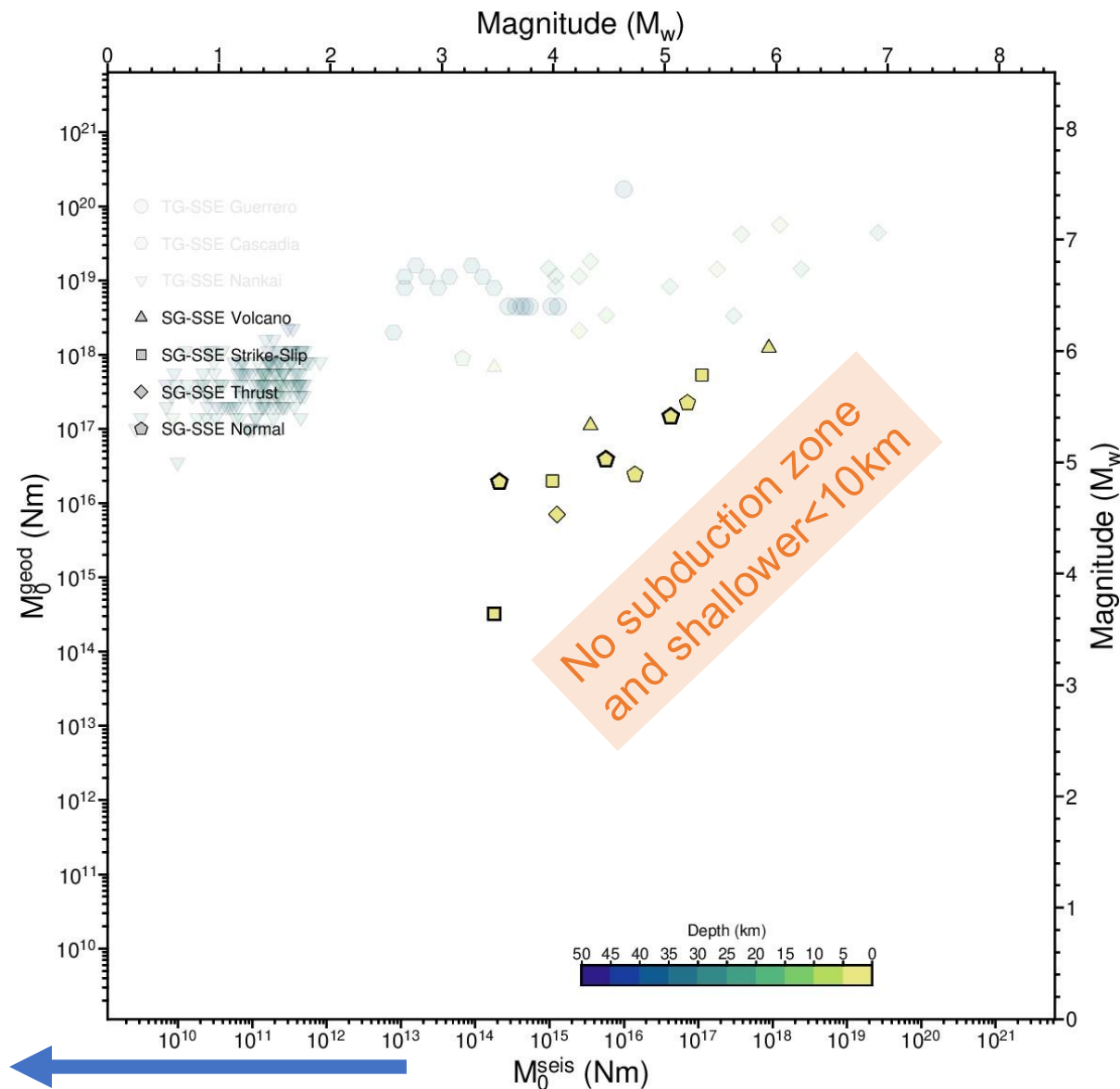
$$M_0^{geod} = M_0^{seis} + M_0^{aseis}$$

- TG-SSE Guerrero
- TG-SSE Cascadia
- ▼ TG-SSE Nankai
- ▲ SG-SSE Volcano
- SG-SSE Strike-Slip
- ◆ SG-SSE Thrust
- ◆ SG-SSE Normal



Aseismic and seismic scaling

In-situ and Lab
Exp



$$M_0^{geod} = \beta (M_0^{seis})^\alpha$$

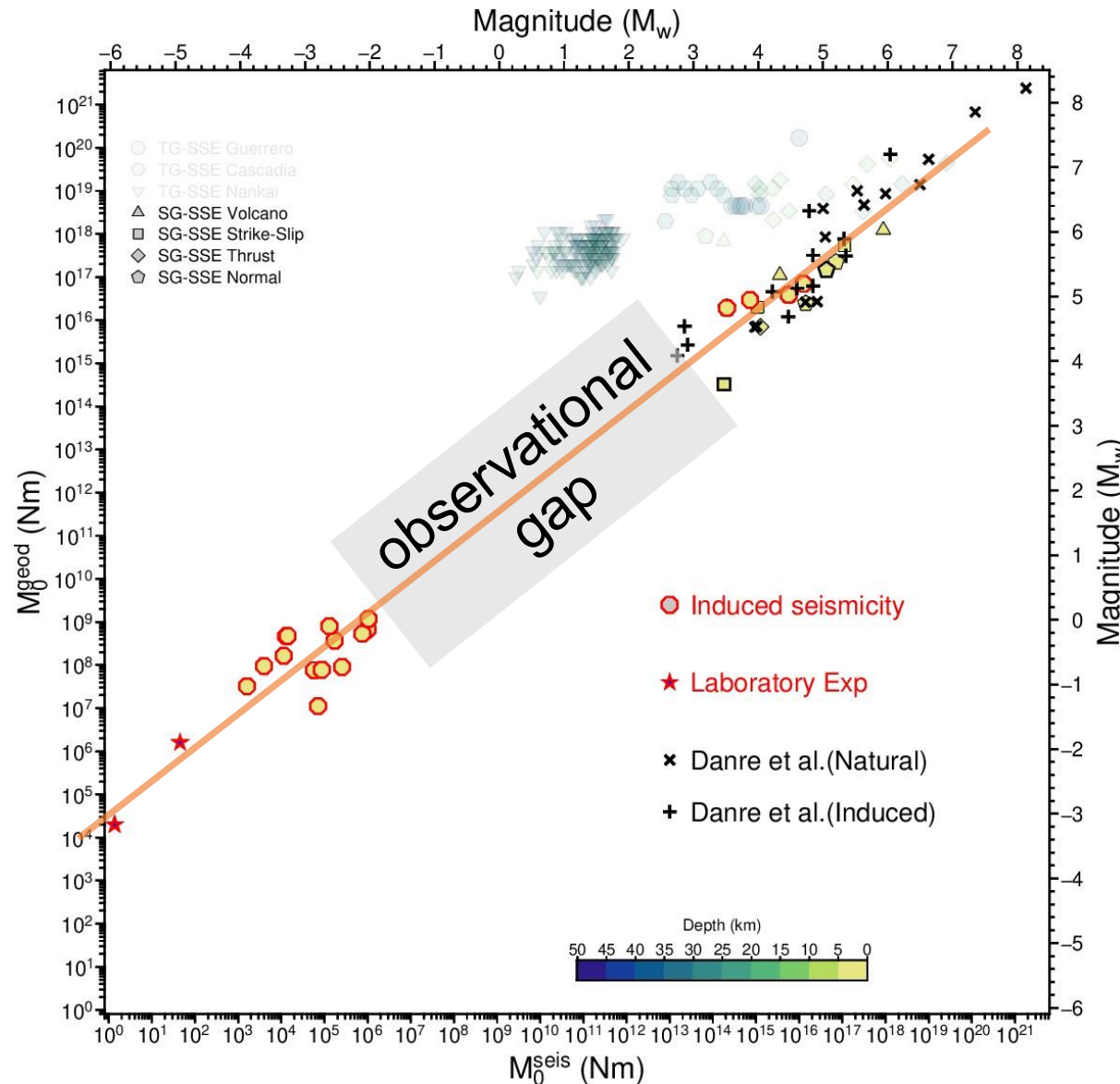
Added new data:

- Natural SSEs
- In-situ field scale injection experiments
- Lab experiments
- Performed simulation

Aseismic and Seismic slip: From nature to lab

$$M_0^{geod} = \beta (M_0^{seis})^\alpha$$

$$\alpha = 0.7$$



x Natural (Danré et al., 2025)

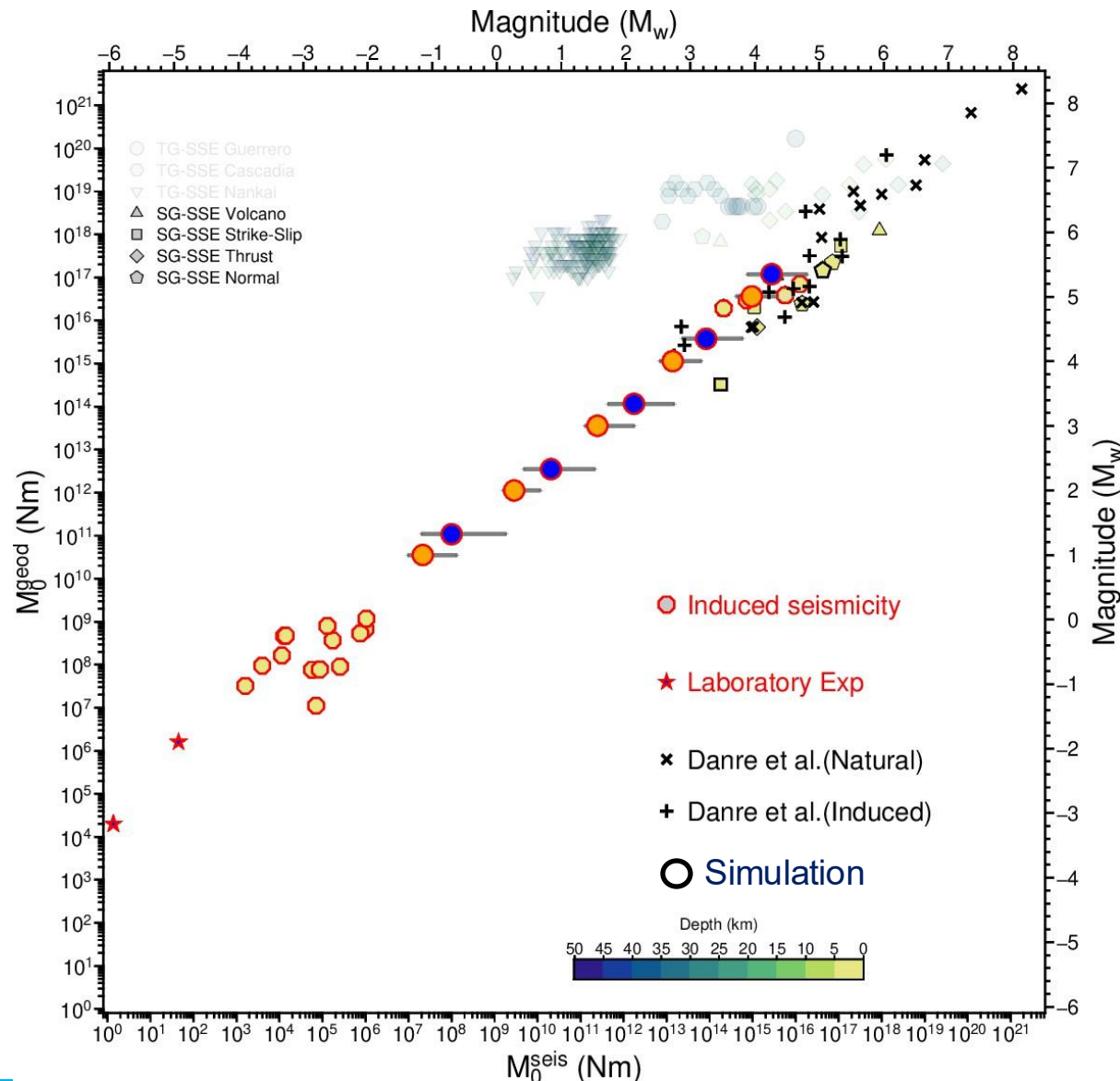
+ Induced (Danré et al., 2025)

★ Laboratory Exp

Aseismic and Seismic scaling

Geodetic vs Seismic Moment scaling:

$$M_0^{geod} = \beta (M_0^{seis})^\alpha$$



- Scaling -> The larger M_0^{geod} the larger the proportion of M_0^{seis}
- Scaling arises by a combination of:
- Confining stress increase with depth - > larger seismic moment release for deeper events
- Geometrical finiteness of the fault activated

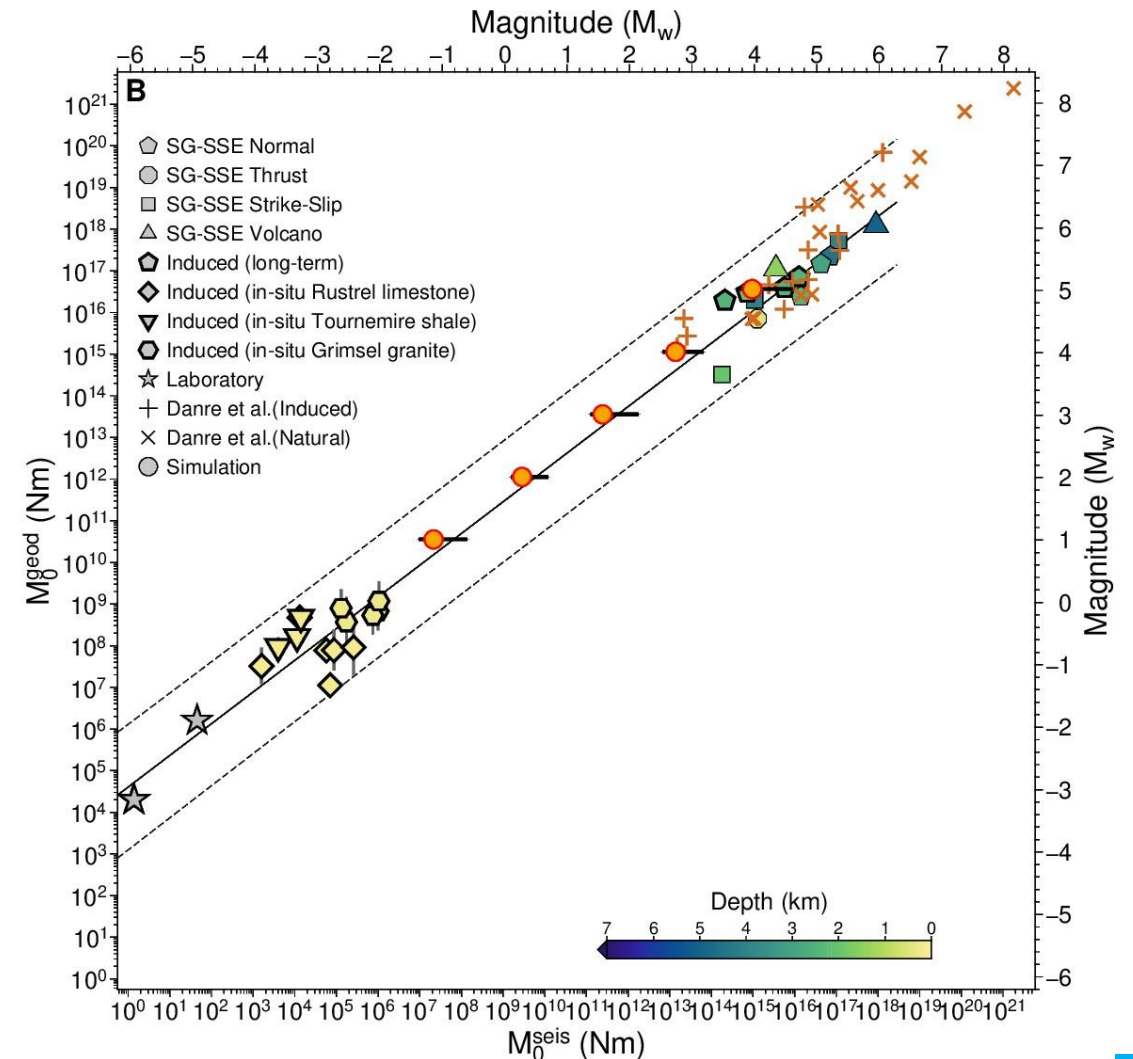
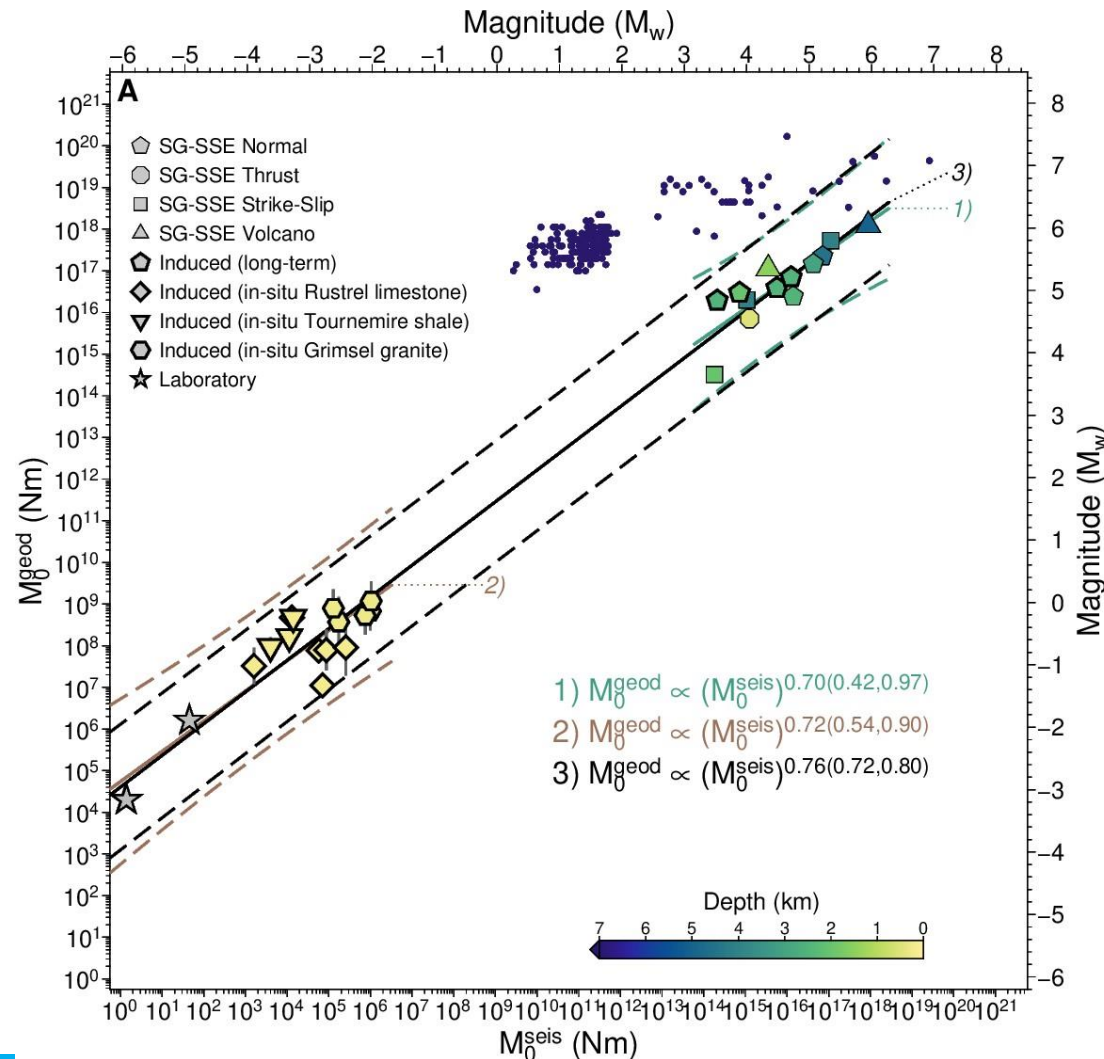
Results

Aseismic and Seismic scaling

Geodetic vs Seismic Moment scaling

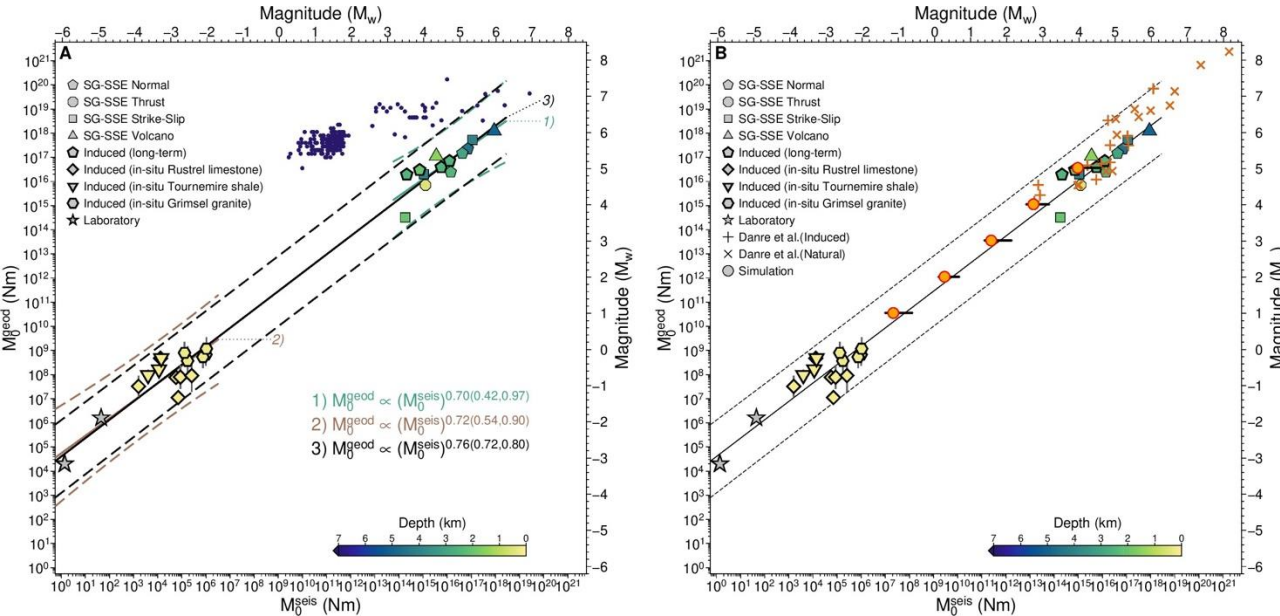
$$M_0^{geod} = 10^{4.6} (M_0^{seis})^{0.76}$$

$$M_0^{geod} = \beta (M_0^{seis})^\alpha \equiv \text{if and only if } M_0^{seis} \exists$$



Aseismic and Seismic scaling

Geodetic vs Seismic Moment scaling



$$M_0^{geod} = \beta (M_0^{seis})^\alpha$$

$\equiv$ if and only if $M_0^{seis} \exists$

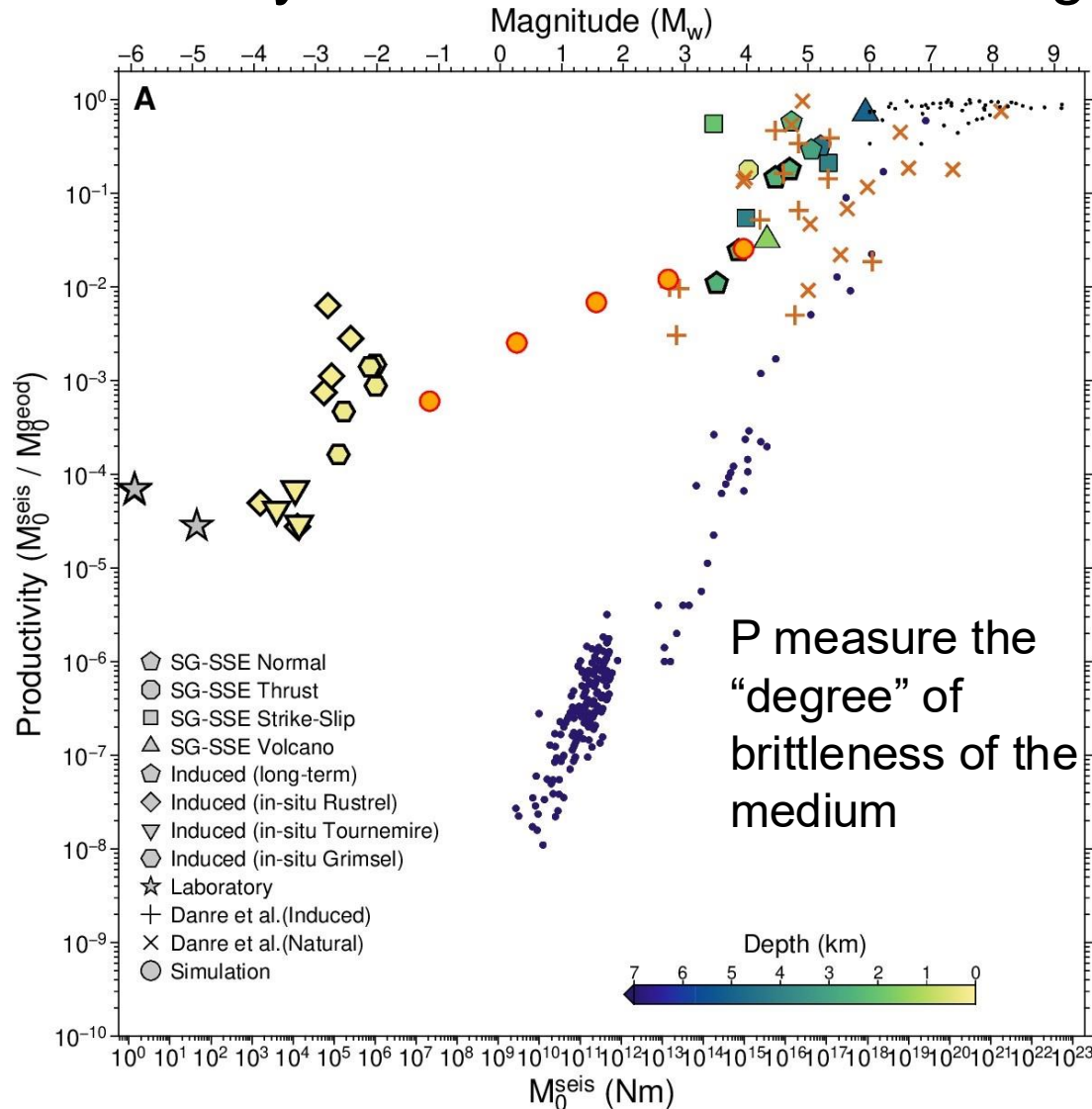
Eq. holds only for seismogenic slow slip transients

$$M_0^{geod} = M_0^{seis} + M_0^{aseis}$$

$$M_0^{geod} = \beta M_0^{seis} \text{ if } \alpha=1$$

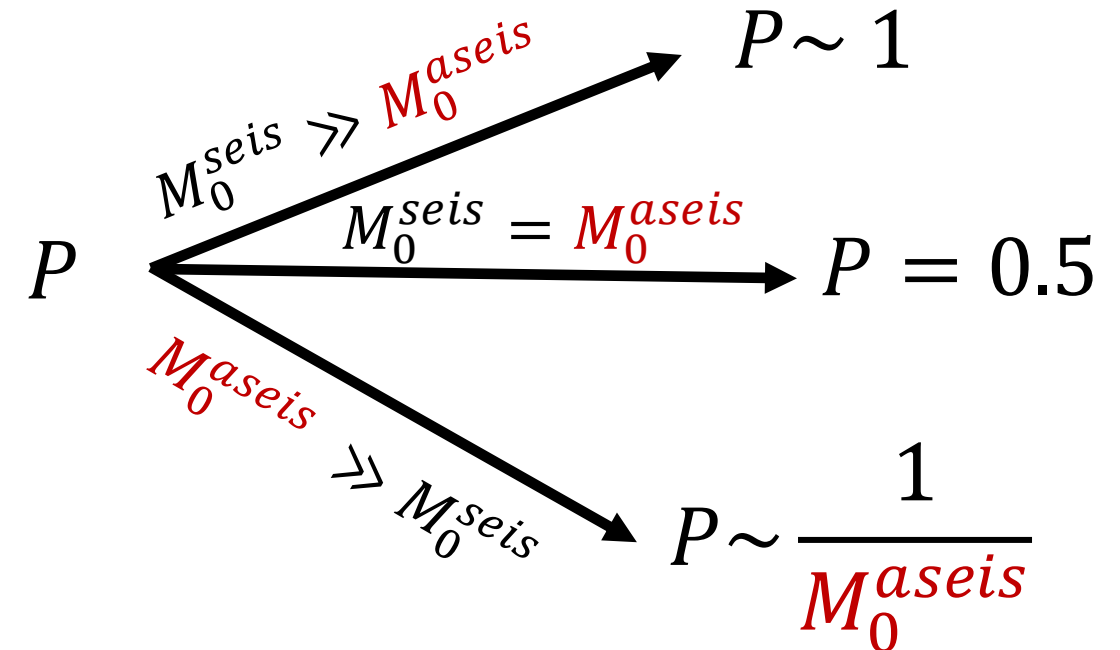
Aseismic and Seismic scaling

Productivity vs Seismic Moment scaling:



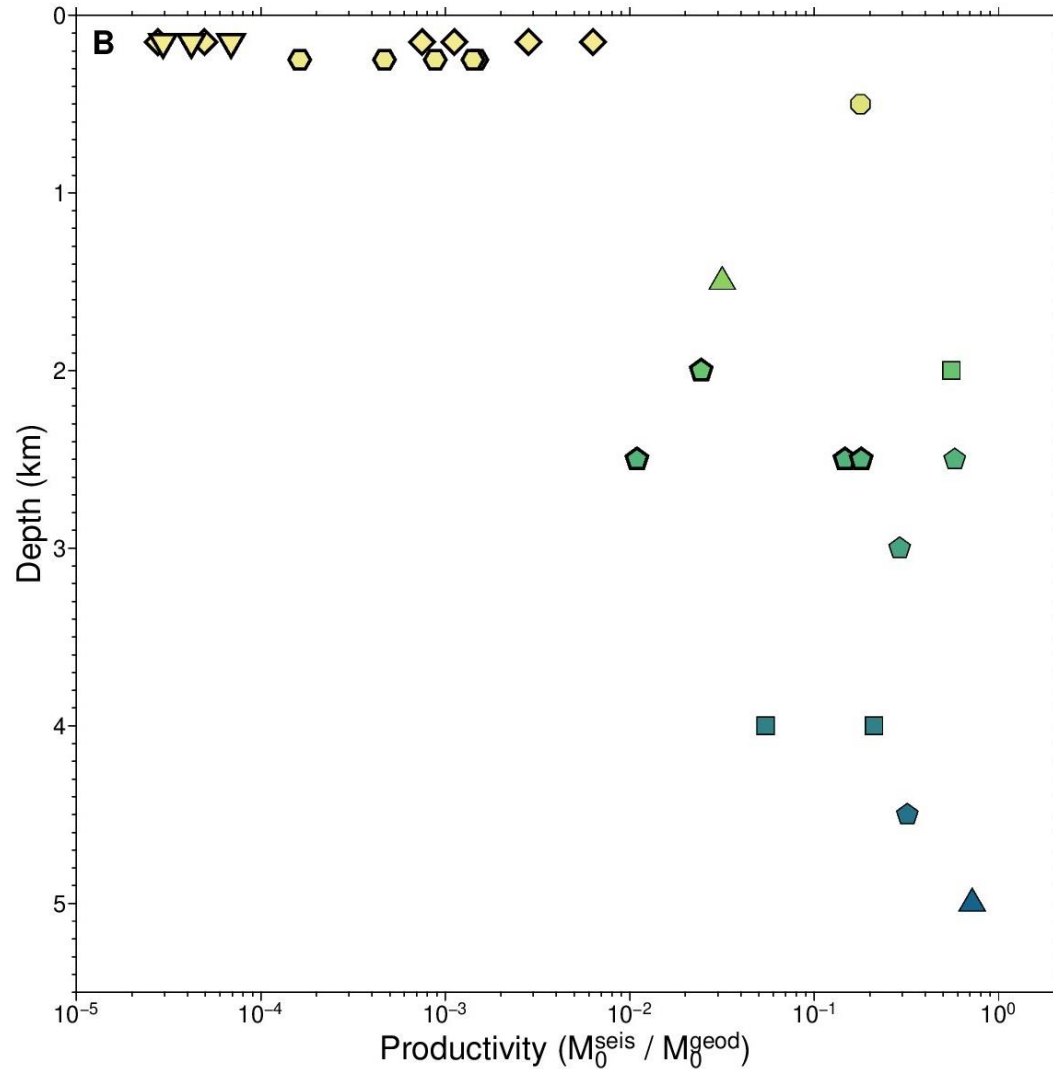
$$M_0^{\text{geod}} = \beta (M_0^{\text{seis}})^\alpha$$

$$P = \frac{M_0^{\text{seis}}}{M_0^{\text{geod}}} = \frac{M_0^{\text{seis}}}{M_0^{\text{seis}} + M_0^{\text{aseis}}} = \frac{(M_0^{\text{seis}})^{1-\alpha}}{\beta}$$



Aseismic and Seismic scaling

Productivity vs Seismic Moment scaling:



$$Depth \sim P$$

The deeper SSE the more seismic productive they are

$$M_0^{seis} \lesssim M_0^{aseis}$$

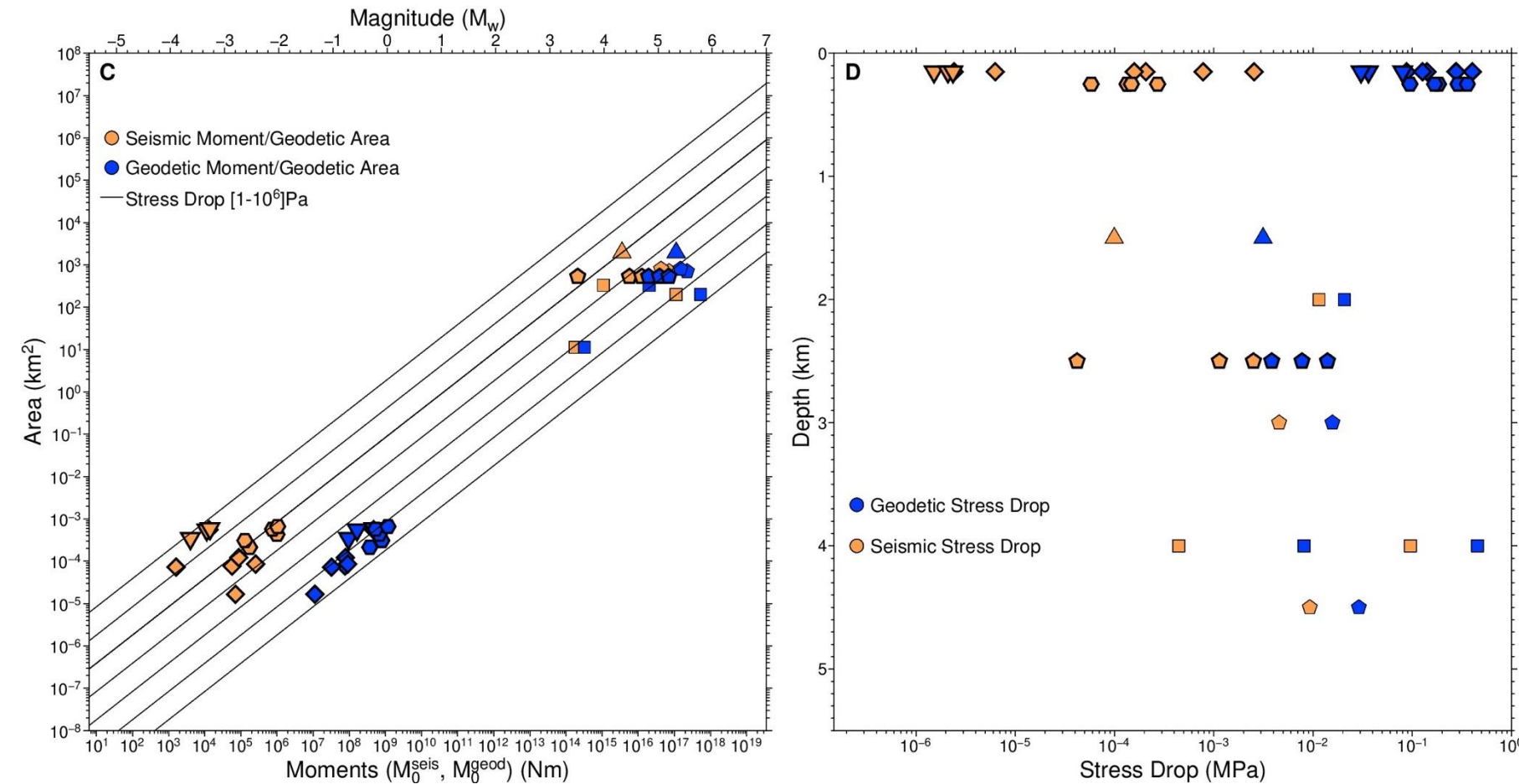
Spearman Correlation ~ 0.85

Aseismic and Seismic slip: From nature to lab

Stress drop

$$\tau_{s,g} = \frac{7}{16} \frac{M_0^{s,g}}{L_{geod}^3}$$

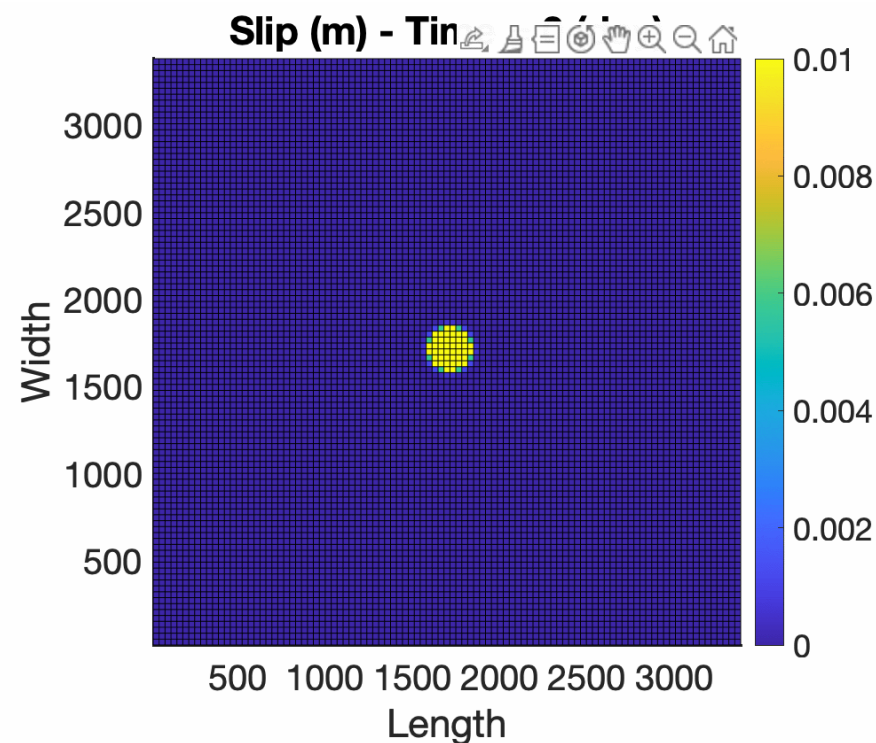
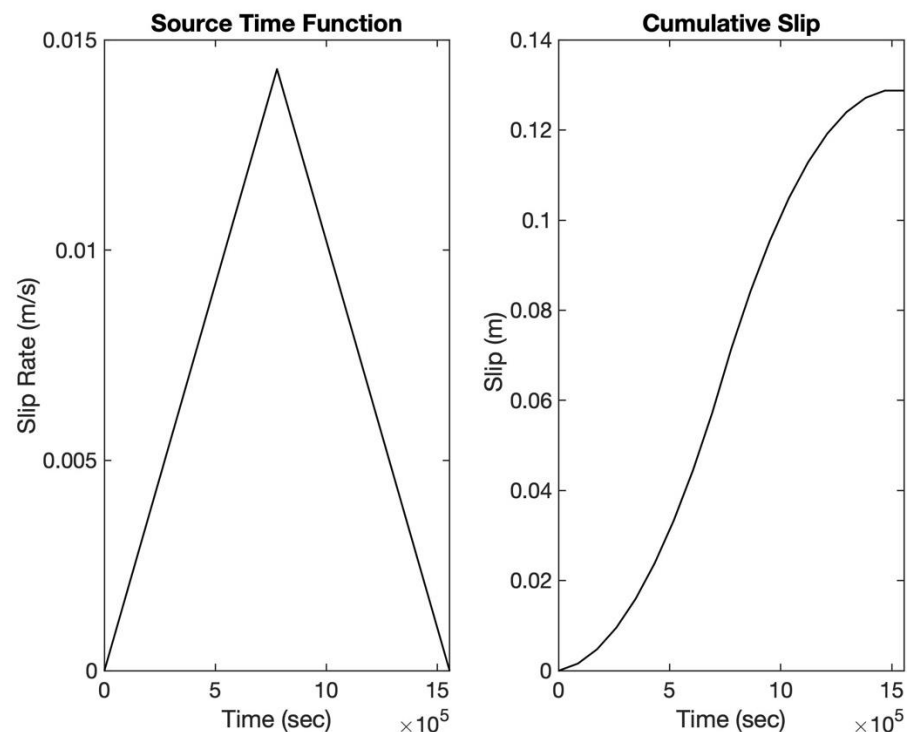
$$\tau \sim P$$



Simulation: Aseismic and Seismic slip

Poroelastic model (Wang and Kümpel, 2003)

Aseismic Source (SSE) w/ triangular Source Time Function and Gaussian slip...



...and calculate poro-elastic stresses in a homogeneous half-space

Simulation: Aseismic and Seismic slip

Physics-based stochastic earthquake simulator (SEED)

(Rinaldi and Nespoli, 2017)

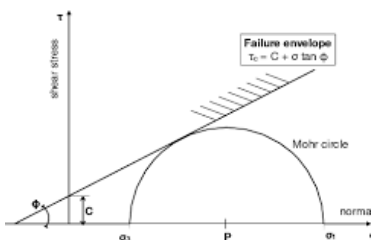
SEED = point source eqk

Initial
random position
(x,y,z)

Random
Focal
Mechanism

Initial random
stress state
 $\Delta\sigma$

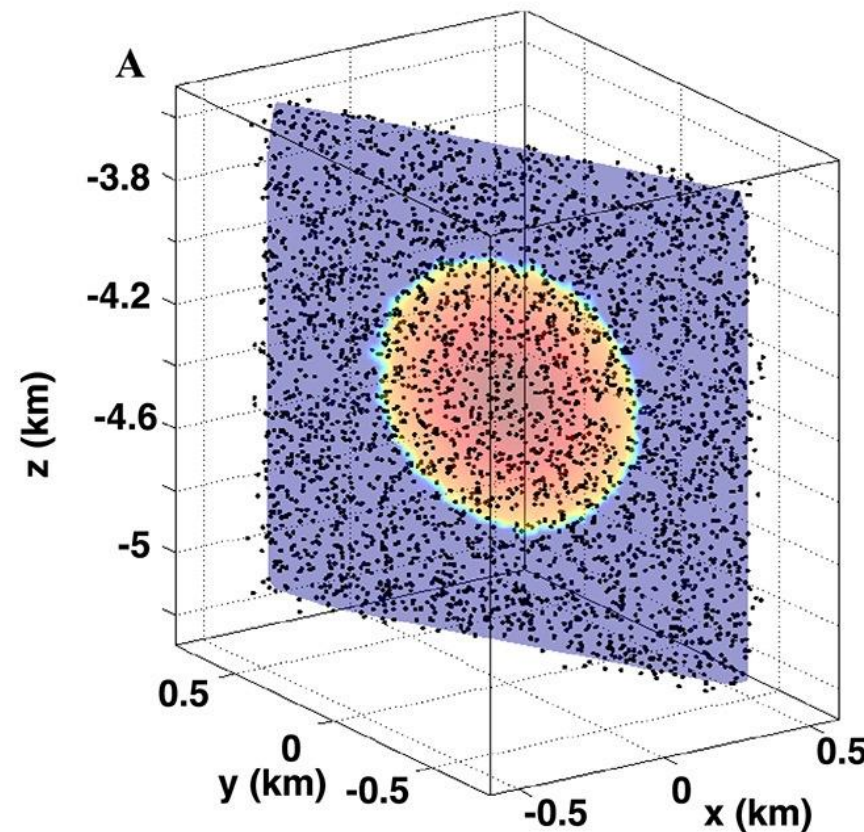
SEED fails according Coulomb failure criterion



b-value

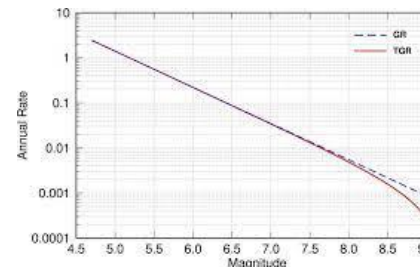
b-value $\sim \Delta\sigma$

b-value = 1



Magnitude
Truncated GR

Catalli et al., 2016



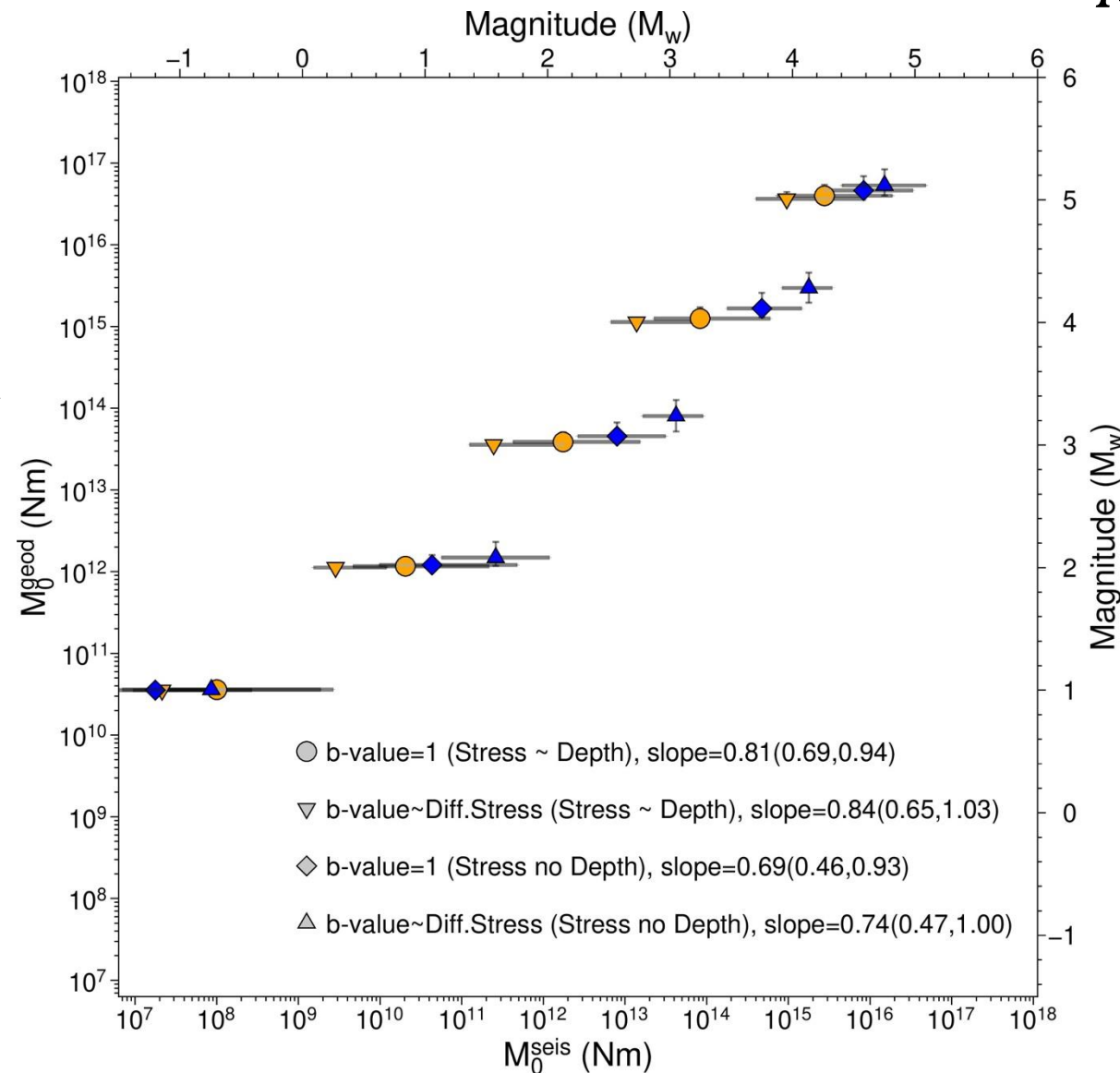
Aseismic and Seismic slip: From nature to lab + Simulations

Geodetic vs Seismic Moment scaling:

SSE magnitude
(source):

$M_w = 1, 2, 3, 4, 5$

- Computational domain scales w/ M_w
- SEED density constant
- $M_{\max}(\text{eqks}) = M_w(\text{SSE})$
- 200 catalogs for each magnitude class



$$M_0^{\text{geod}} = \beta (M_0^{\text{seis}})^\alpha$$

Simulation:

$$b \sim \Delta\sigma \text{ and } b = 1$$

w/ Depth-dependent stress or not

Finiteness of the fault matters, i.e. truncated GR

Aseismic and seismic scaling

Physical explanation

Physical mechanisms: fault pressurization (unclamping)

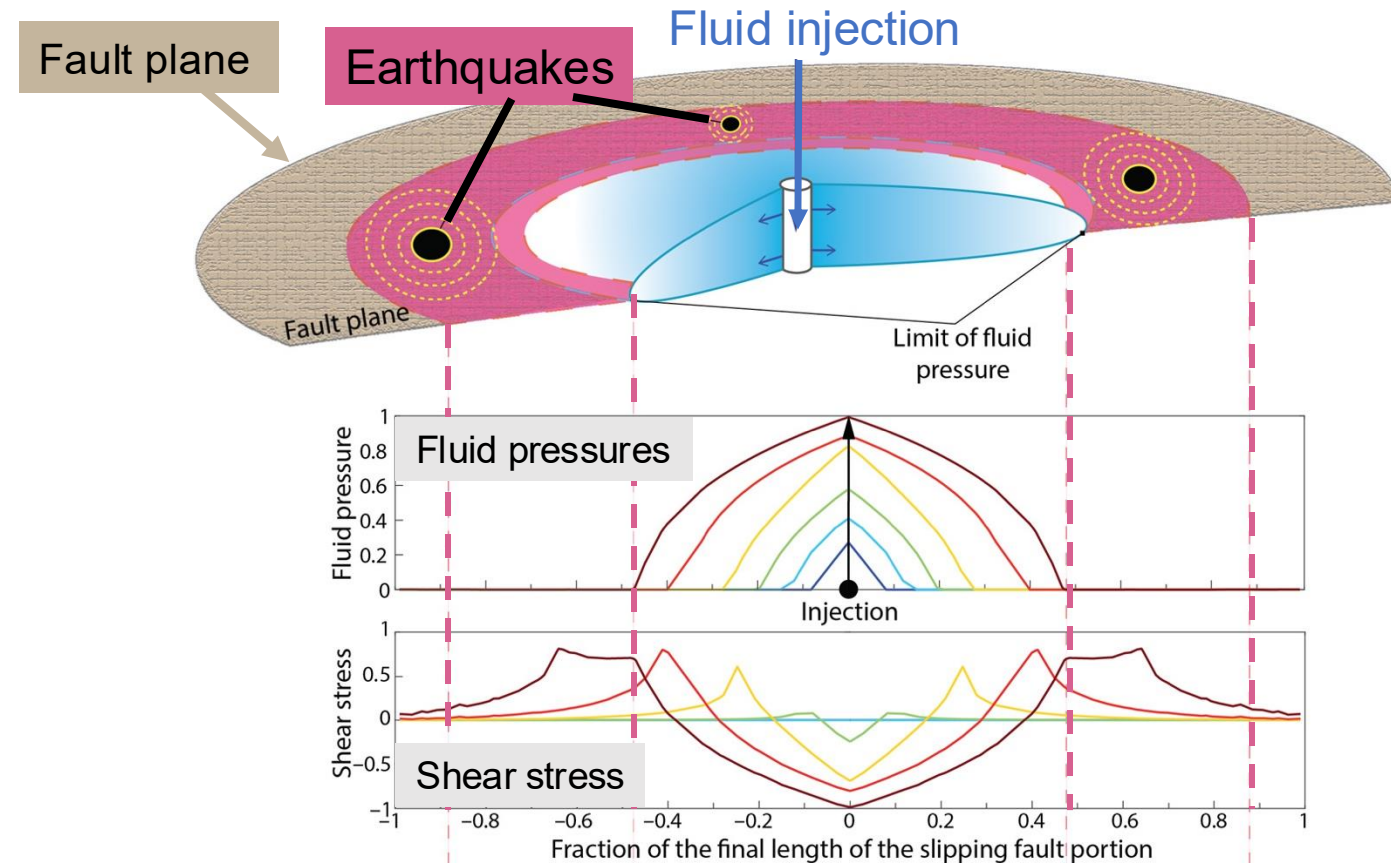
Pressure unclamps the fault and produces aseismic slip

High-pressure induces aseismic slip to expand (around the injection point)

Propagation of the aseismic rupture triggers seismicity due to shear stress increase at the rupture edges

$$\text{Shallower } M_0^{geod} = \beta (M_0^{seis})^{0.7}$$

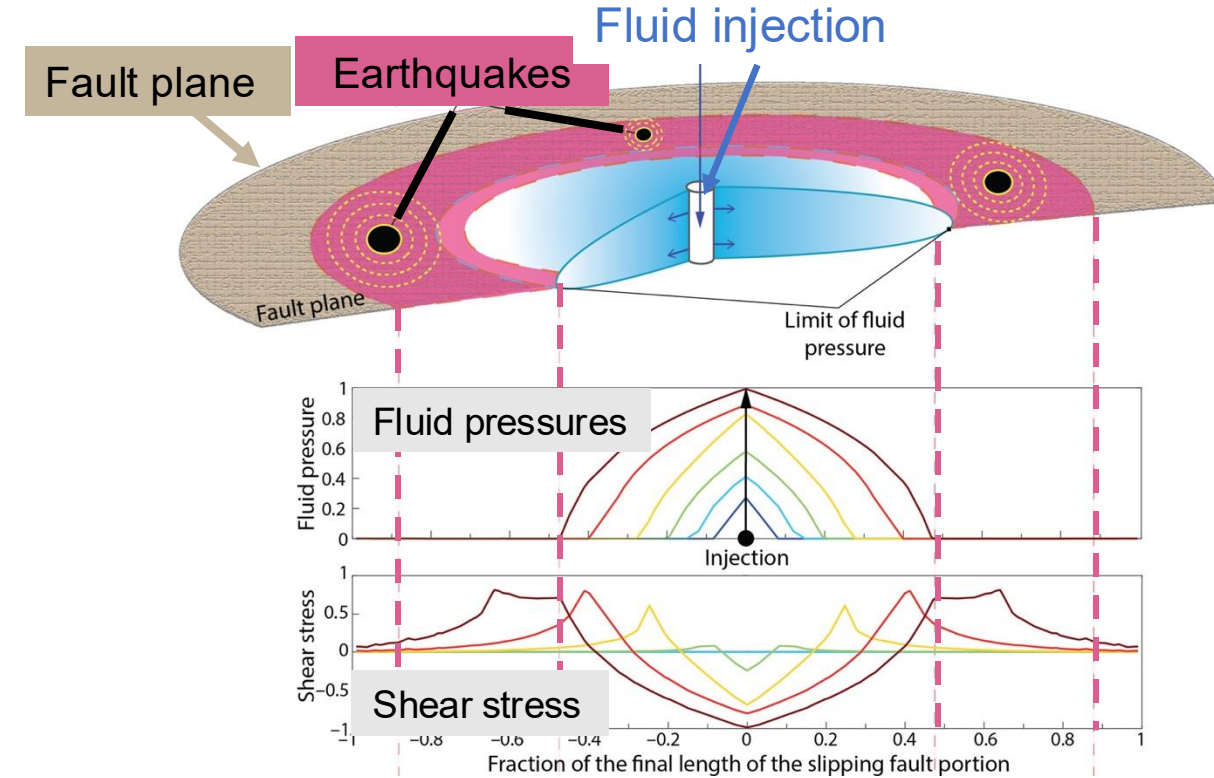
(outside subduction zones)



Cappa et al., 2019

Conclusions

- Seismic to aseismic moment scaling over 14-18 orders of magnitude (but observational gap??)
- If $M_a \gg M_s$ scaling holds over large range of orders of magnitude
- Geometrical finiteness of the fault constrains aseismic and seismic moment release
- Fluid-induced pressurization of fault is a viable mechanism to trigger slow and fast slip

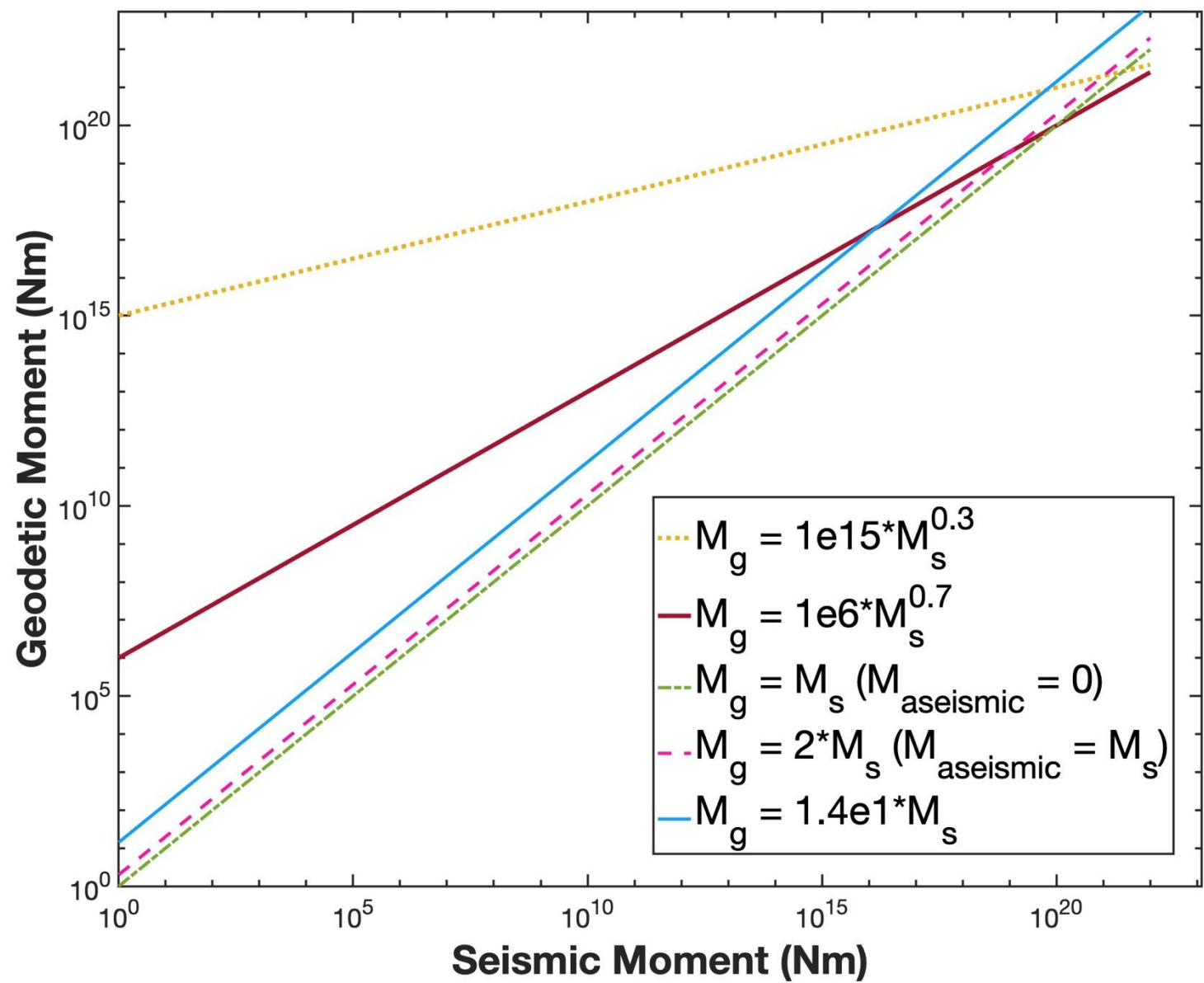


Cappa et al., 2019

Thank you

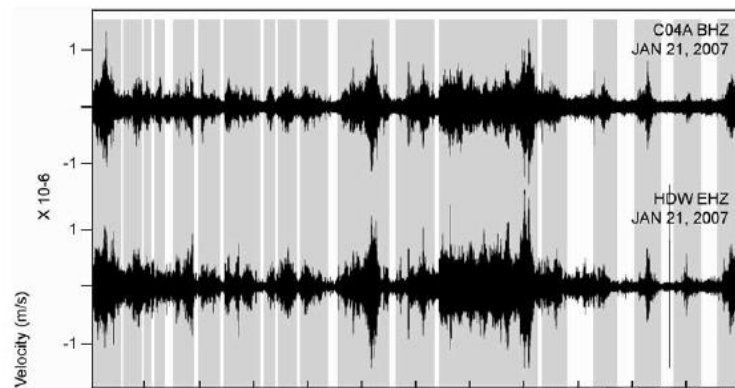
START OF BACK UP SLIDES:

please contact me at luigi.passarelli@ingv.it for further explanation since
slides are very minimalistic



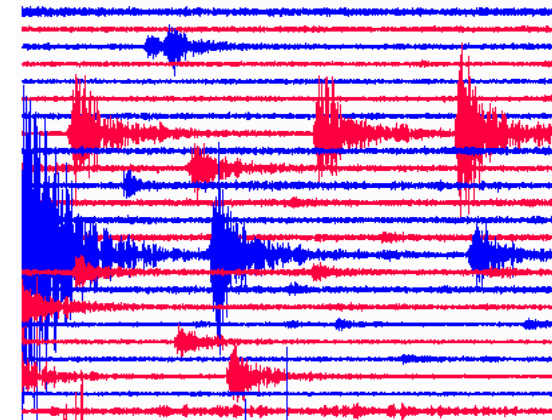
Slow slip transients are rupture that do not excite seismic wave and are often “associated” w/ “*seismic*” events:

Non-volcanic Tremor
(many tiny earthquakes)



Aguilar et al., 2009

Swarms of “ordinary”
(larger) earthquakes

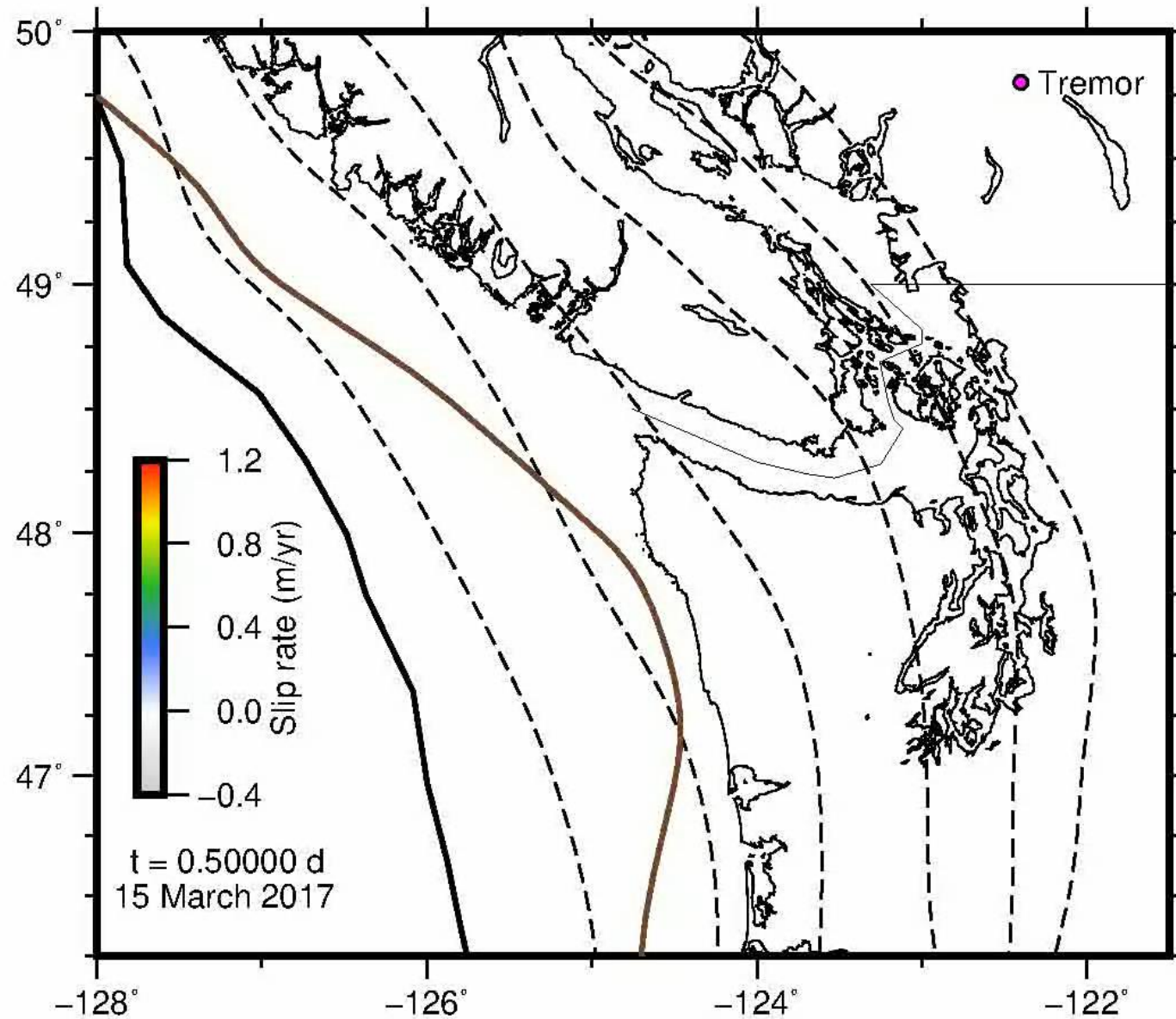


GEOFON TNTI stazione

and/or

Tremor-Genic Slow Slip Events
TG-SSEs

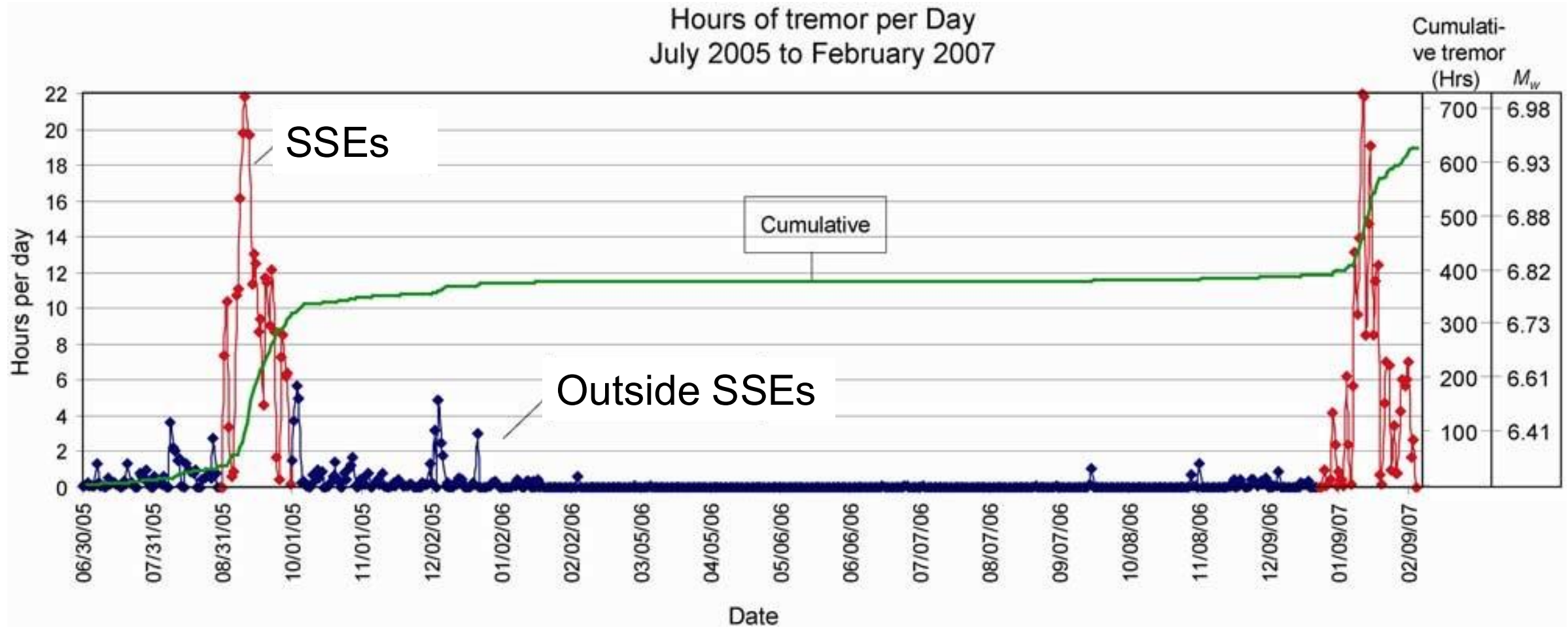
Swarm-Genic Slow Slip Events
SG-SSEs



Itoh et al., 2022

Slow Slip Events (SSEs)

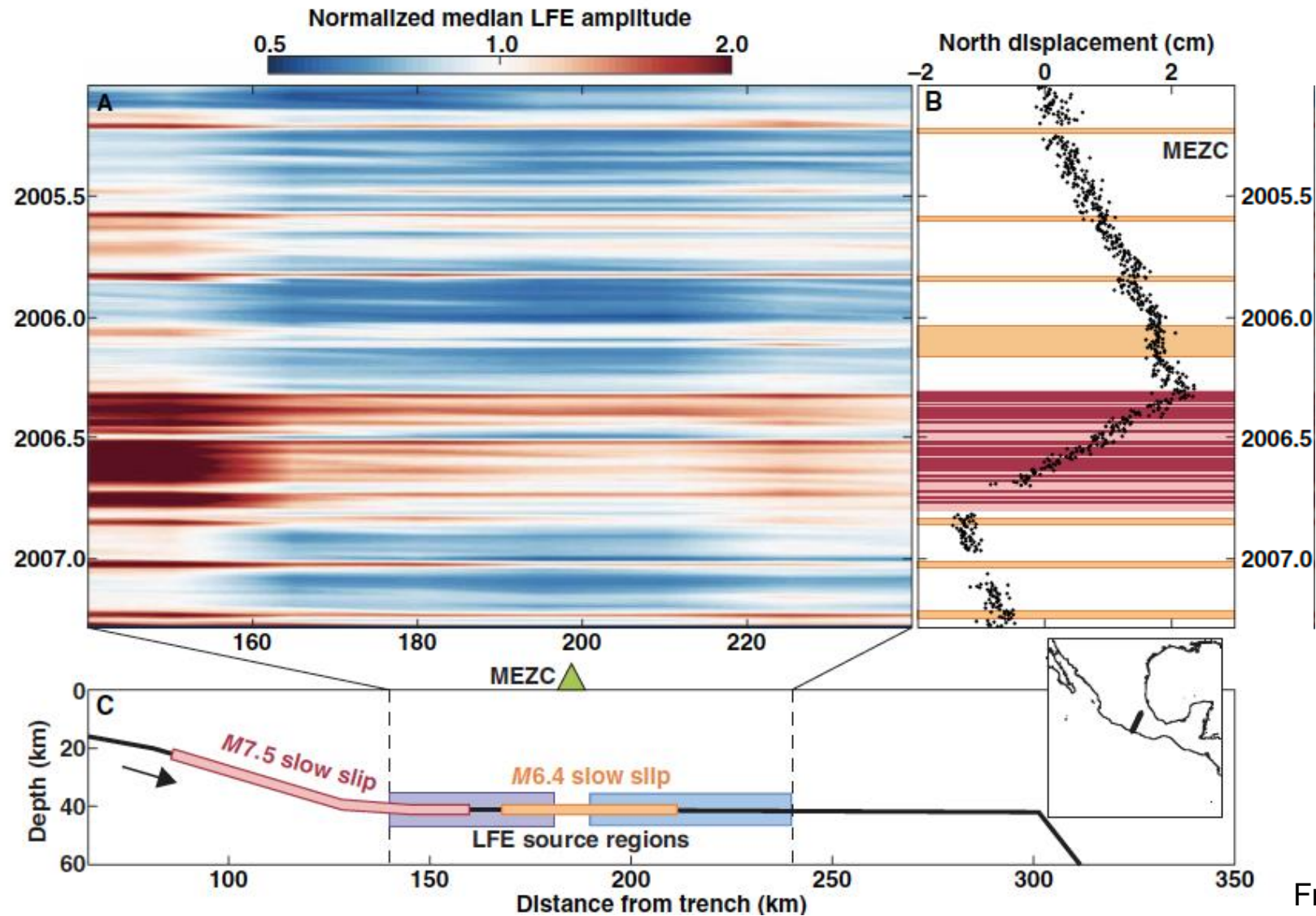
Tremor or LFEs counts increases during SSEs



Aguiar et al., 2009

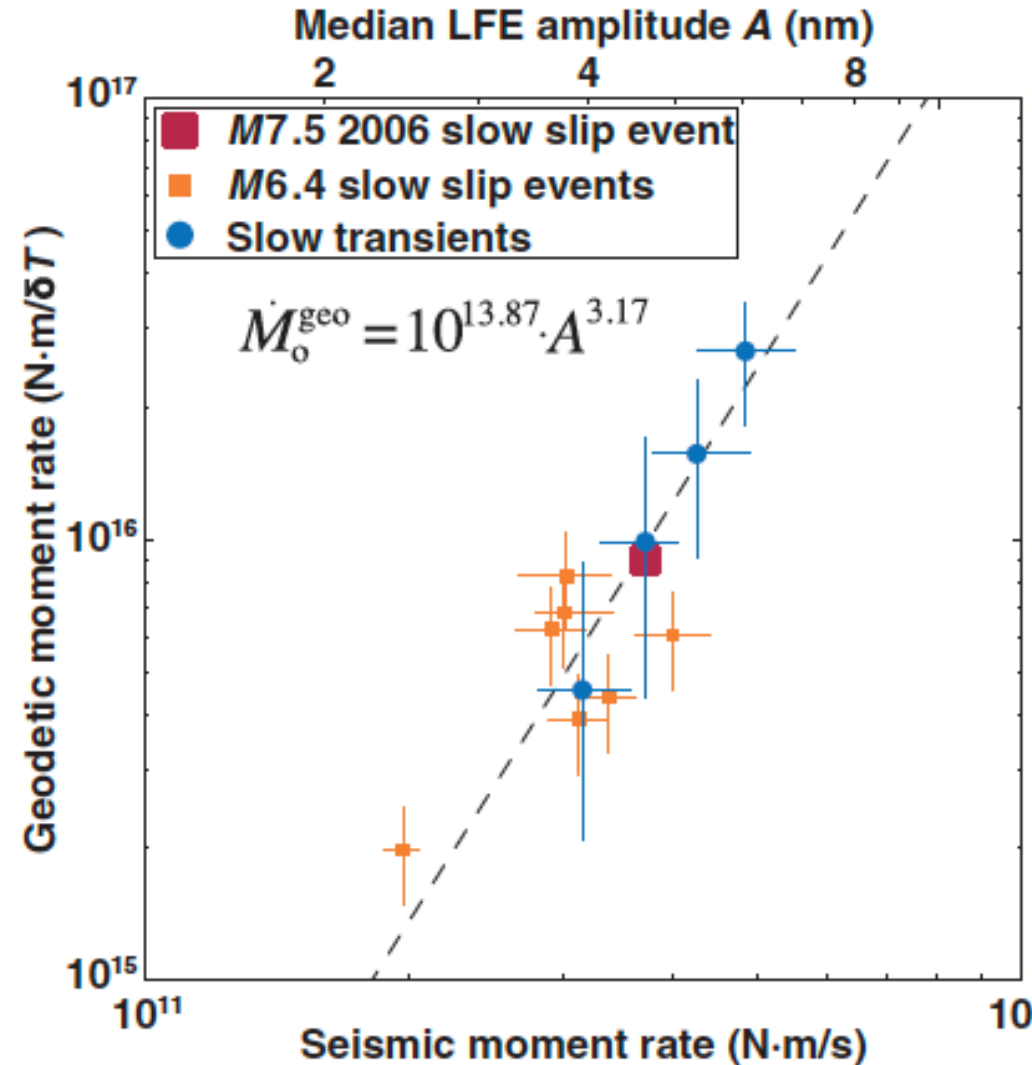
Slow Slip Events (SSEs)

Tremor or LFEs amplitudes (energy) increase during SSEs



Frank and Brodsky 2019

Tremor or LFEs moment (slip) rate correlates with moment (slip) rate of SSEs



Frank and Brodsky 2019

⇒ Fault structures where *slow and aseismic* and *fast and seismic* slip can be accommodated at the same time

⇒ Conceptual model:

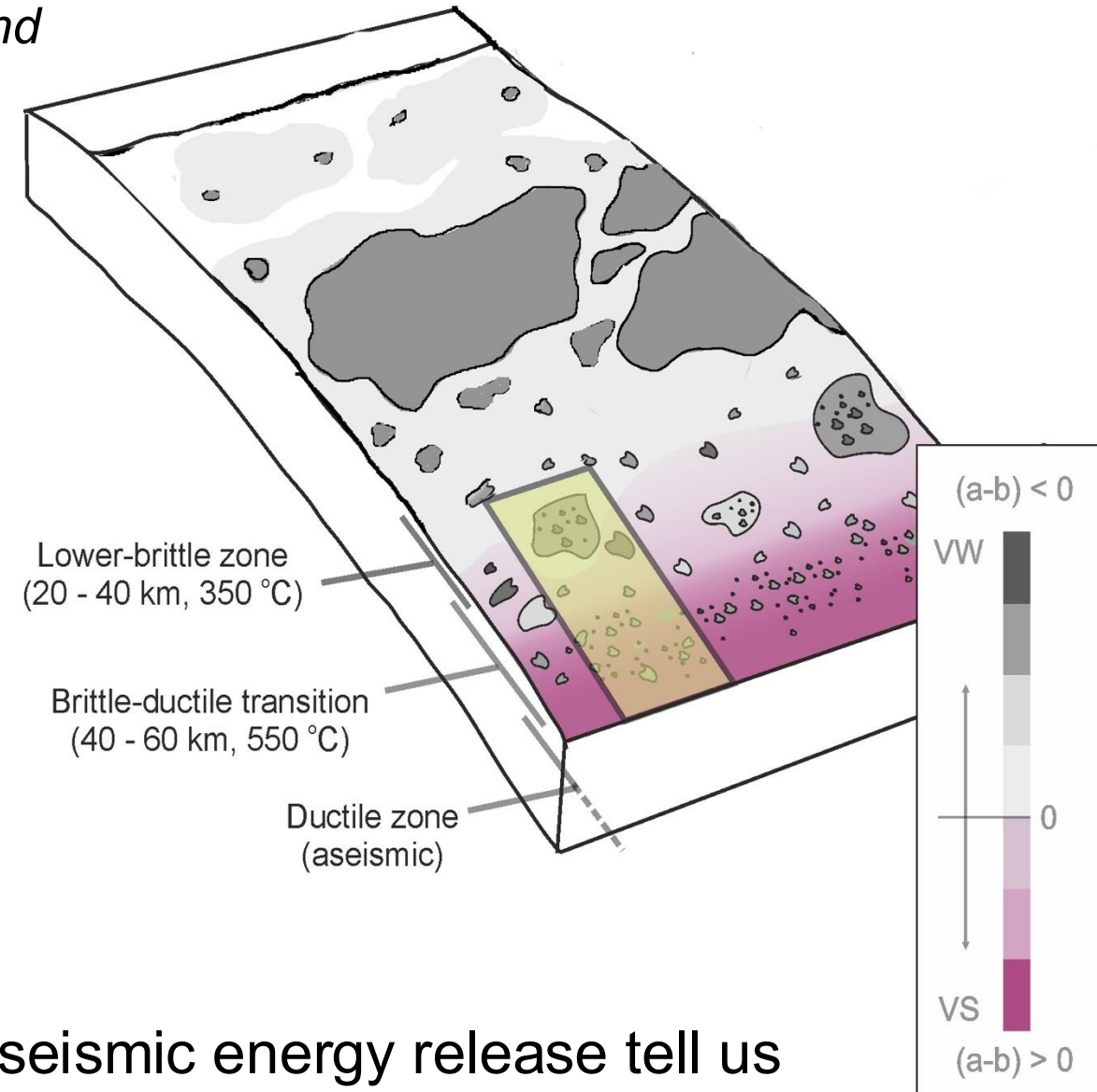
Brittle asperities (darker gray) embedded within conditionally stable (light gray) and stable sliding regions (pinkish)

⇒ Mechanical condition:

Heterogeneous frictional properties

Pronounced fault roughness

High pore pressure condition (low effective normal stress)



Does the proportion of aseismic and seismic energy release tell us about the mechanical condition at the source?

Aseismic and seismic scaling

Physical explanation

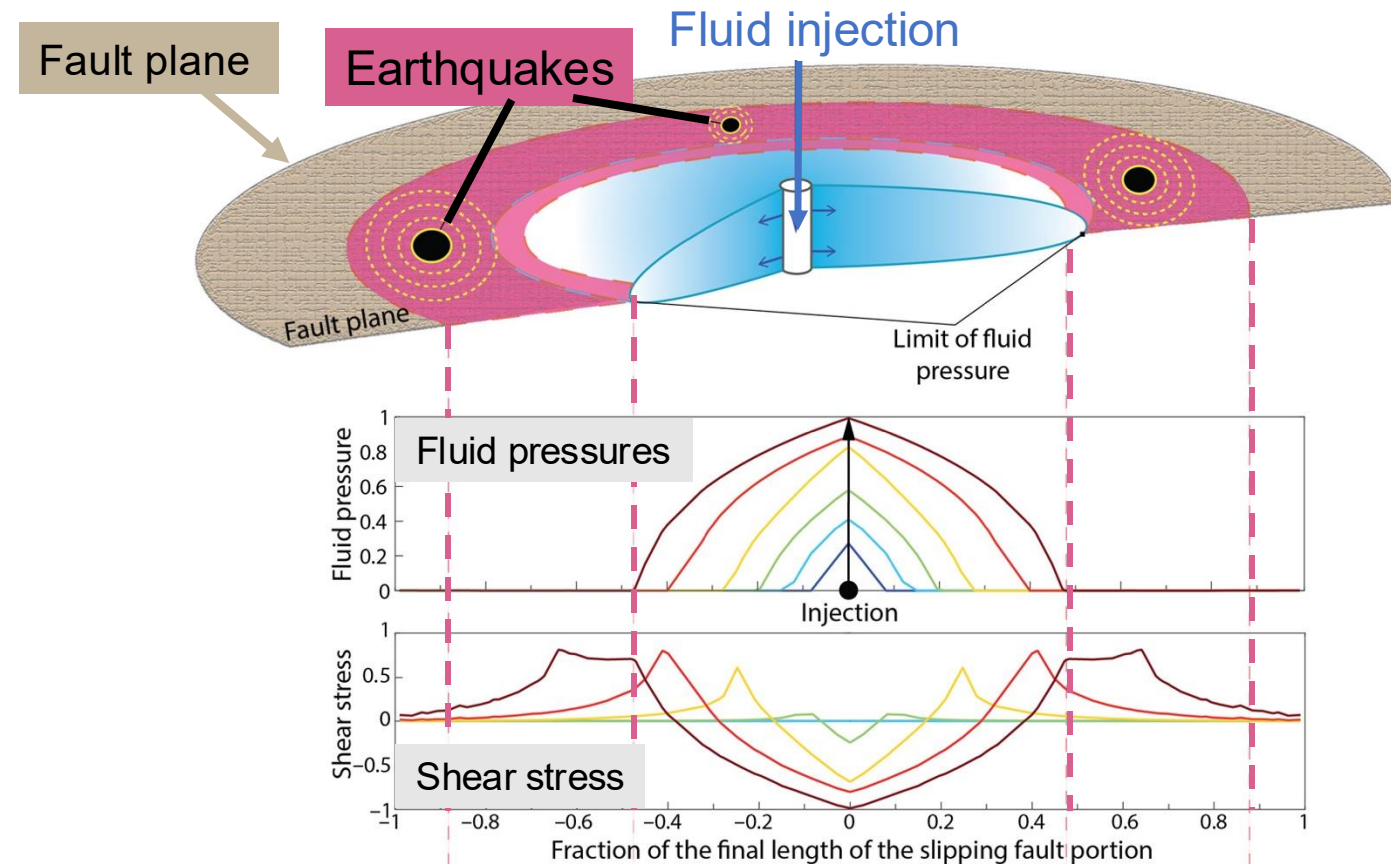
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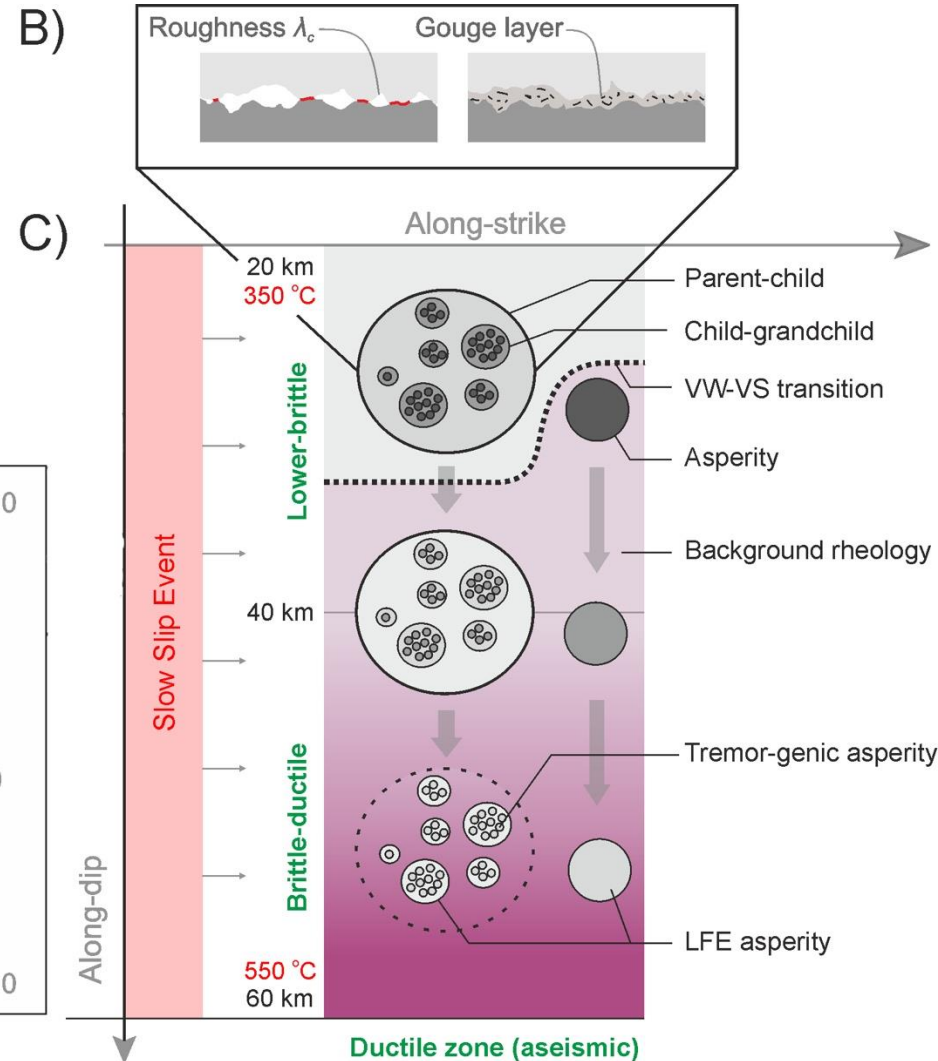
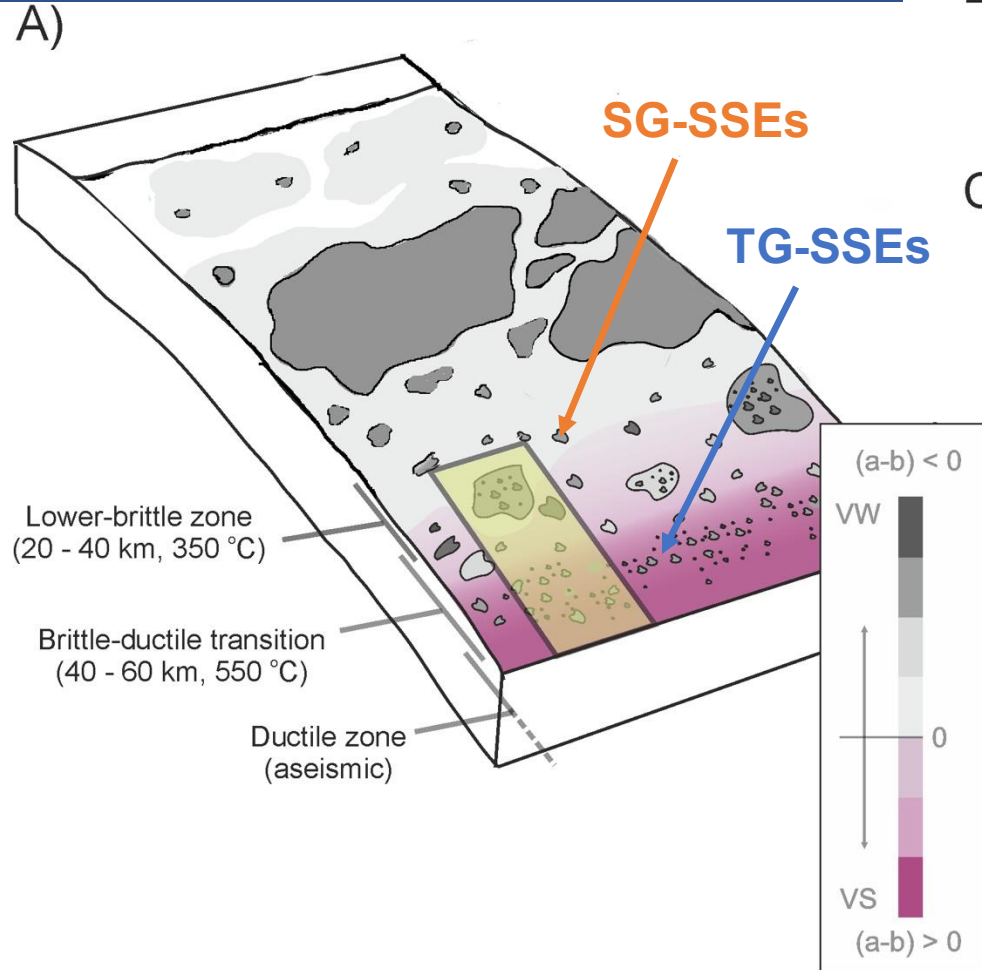
Shallower $M_0^{geod} = \beta (M_0^{seis})^{0.7}$
(outside subduction zones)



Cappa et al., 2019

Deeper

$$M_0^{geod} = \beta (M_0^{seis})^{0.3}$$

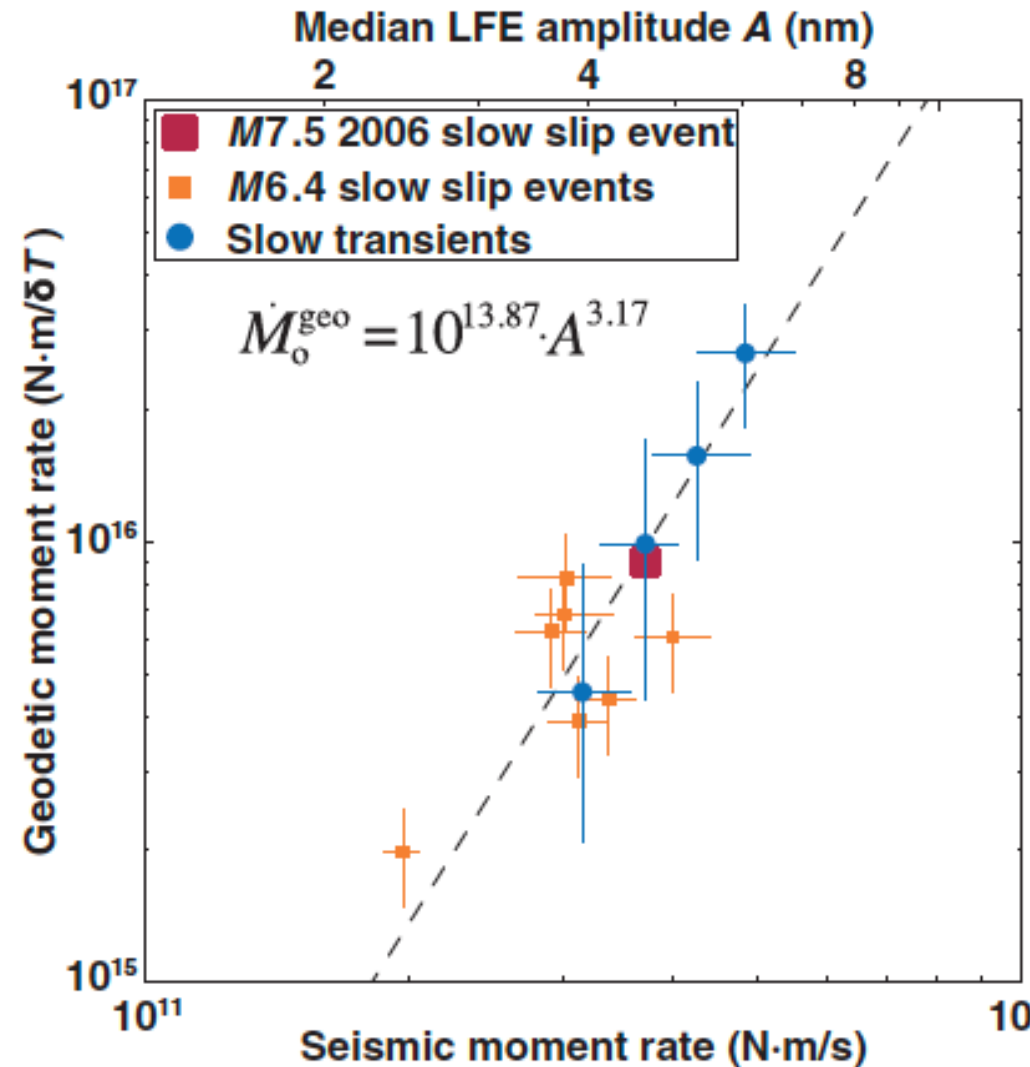


Pronounced fault roughness

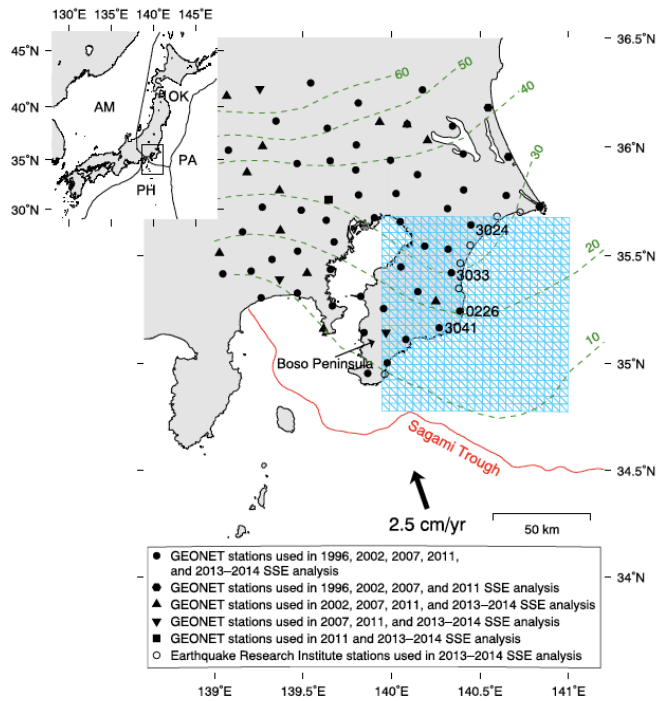
High pore pressure

Heterogeneous frictional properties dependent on temperature

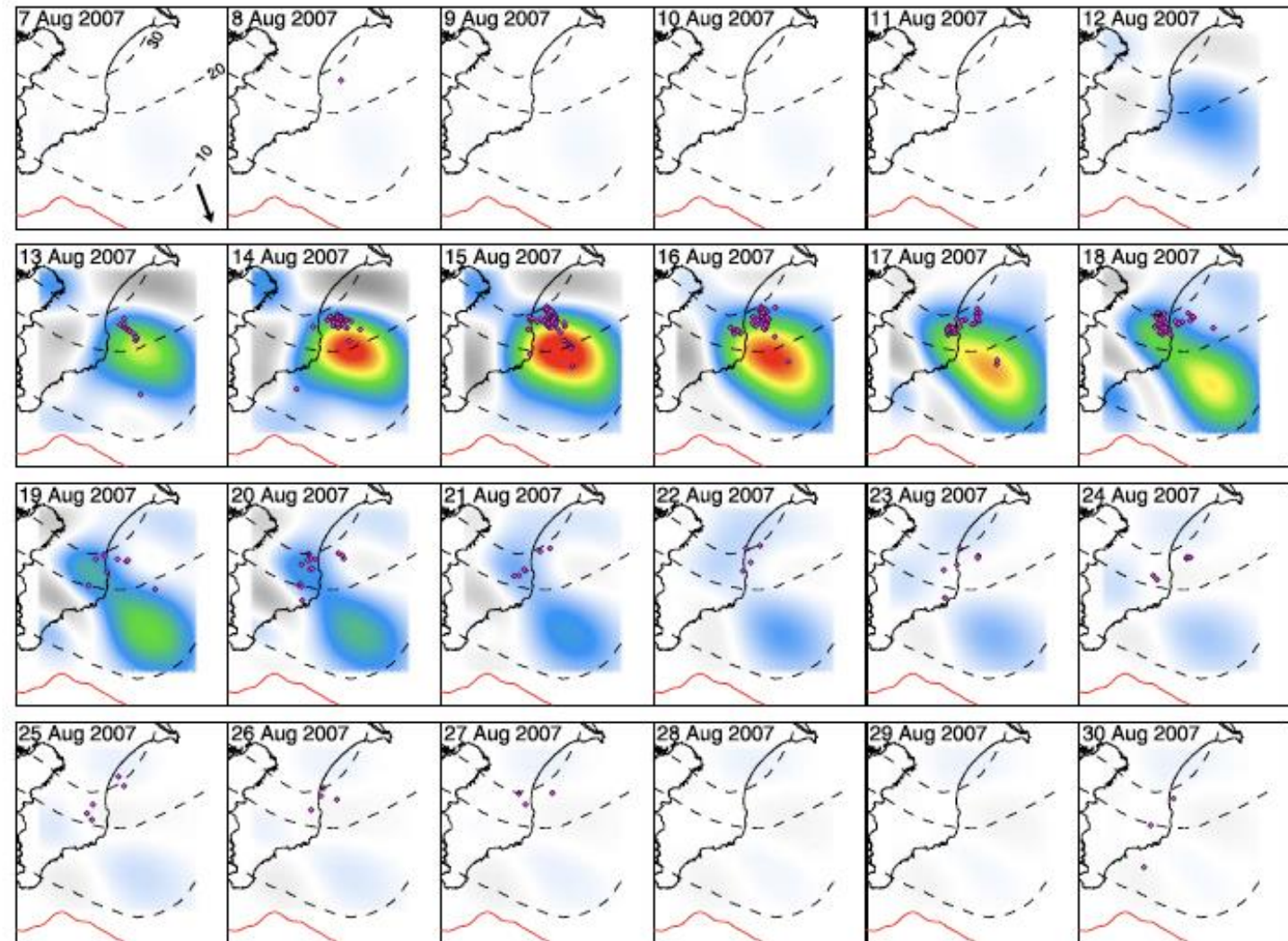
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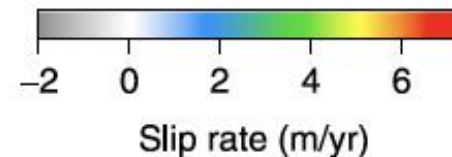
Frank and Brodsky 2019



SG-SSEs in 2007
underneath Boso
peninsula, Japan

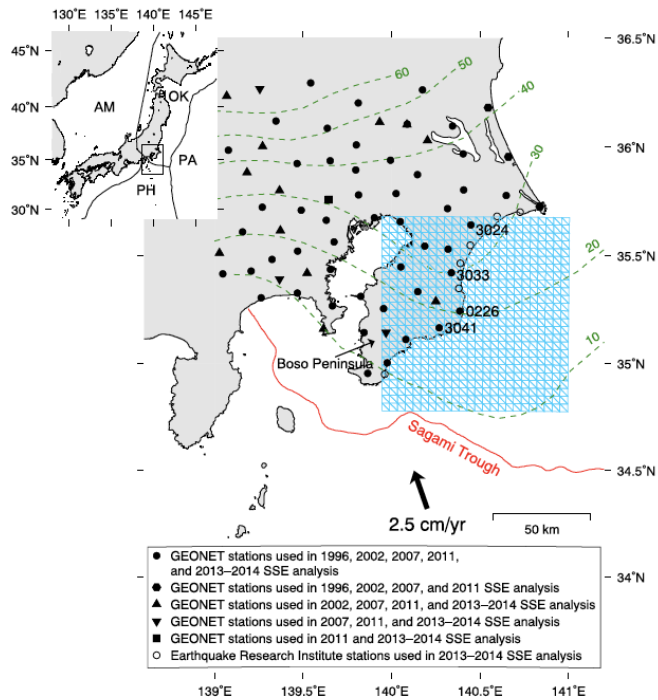


● earthquakes

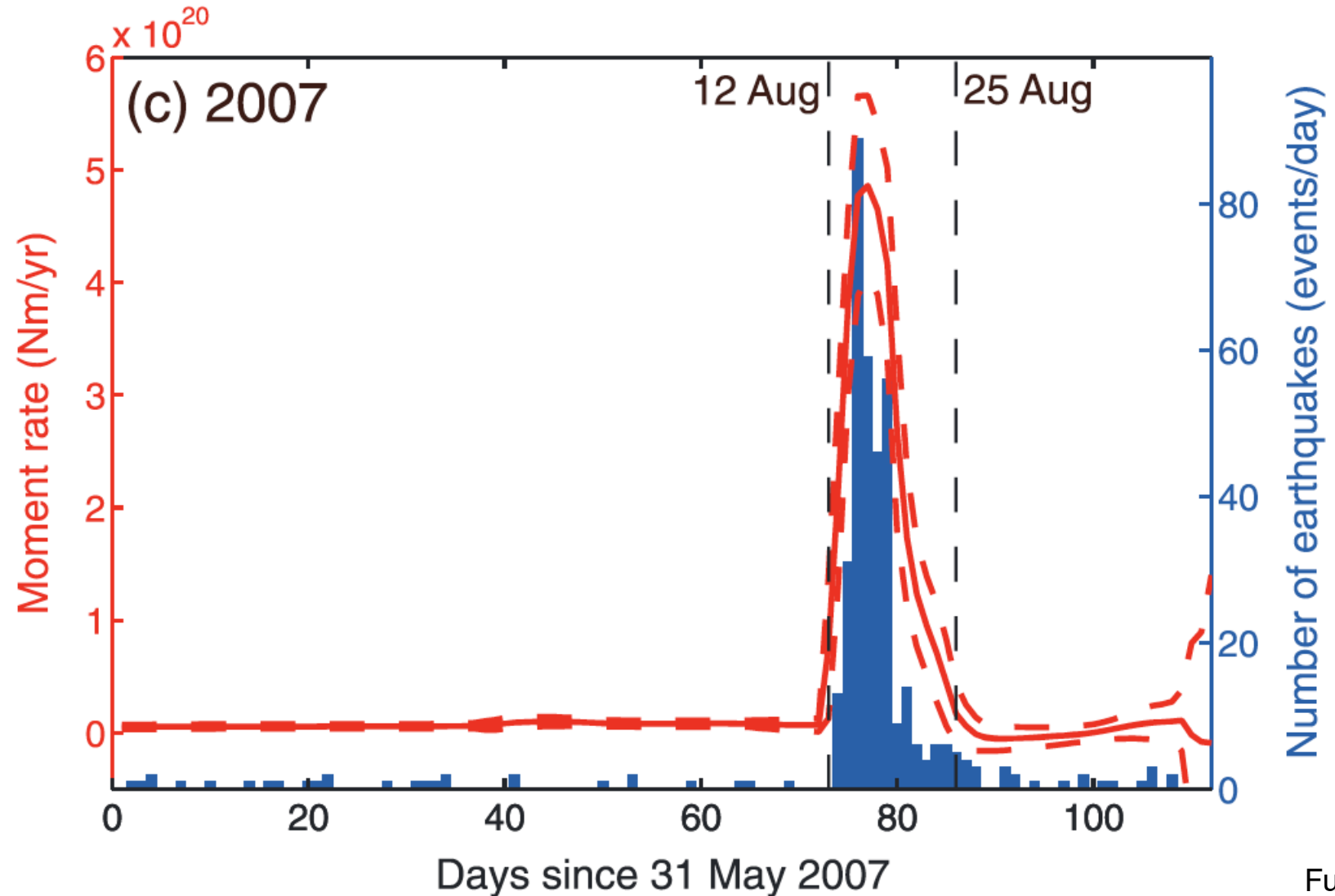


50 km

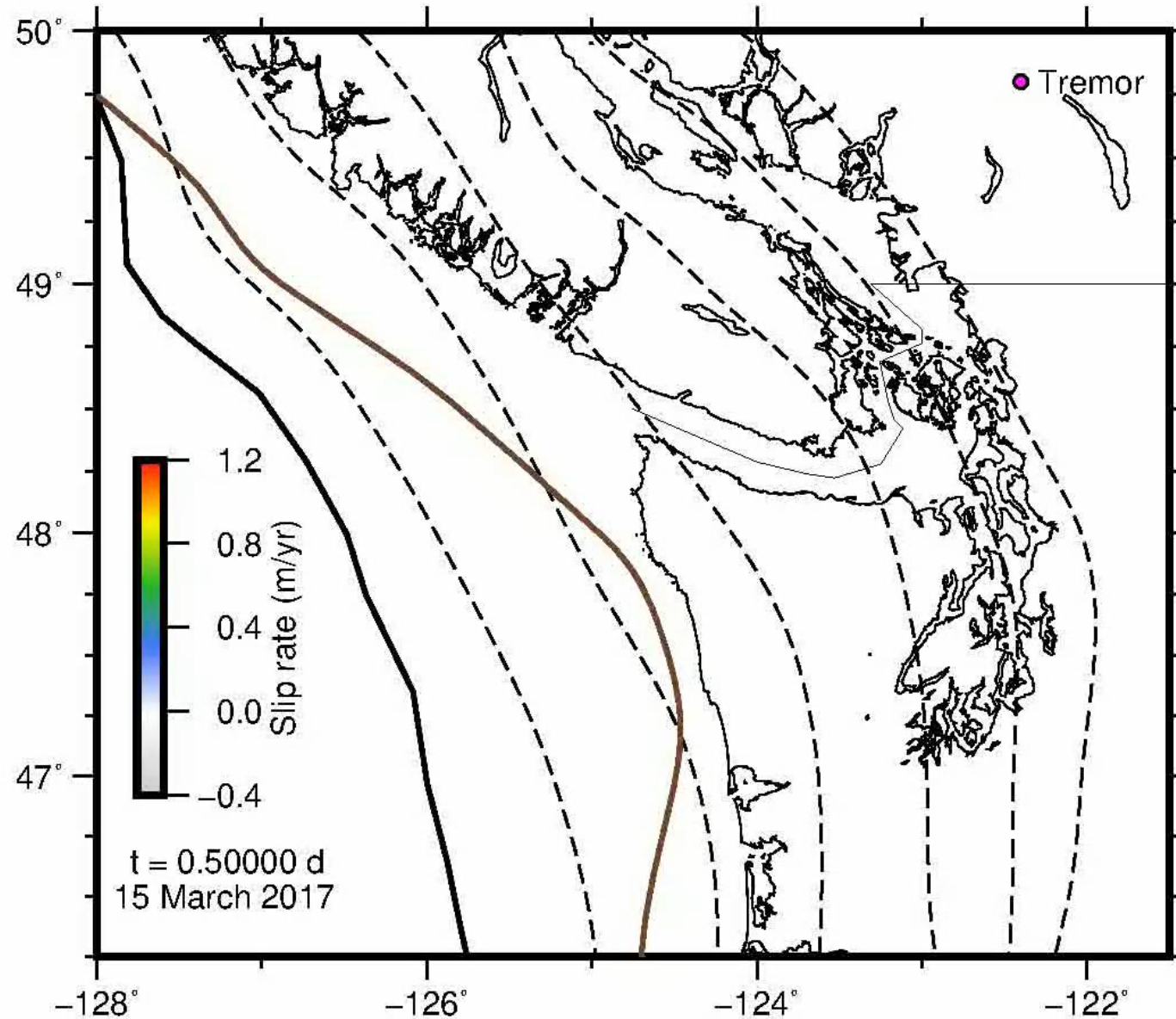
Fukuda, 2018



SG-SSEs in 2007
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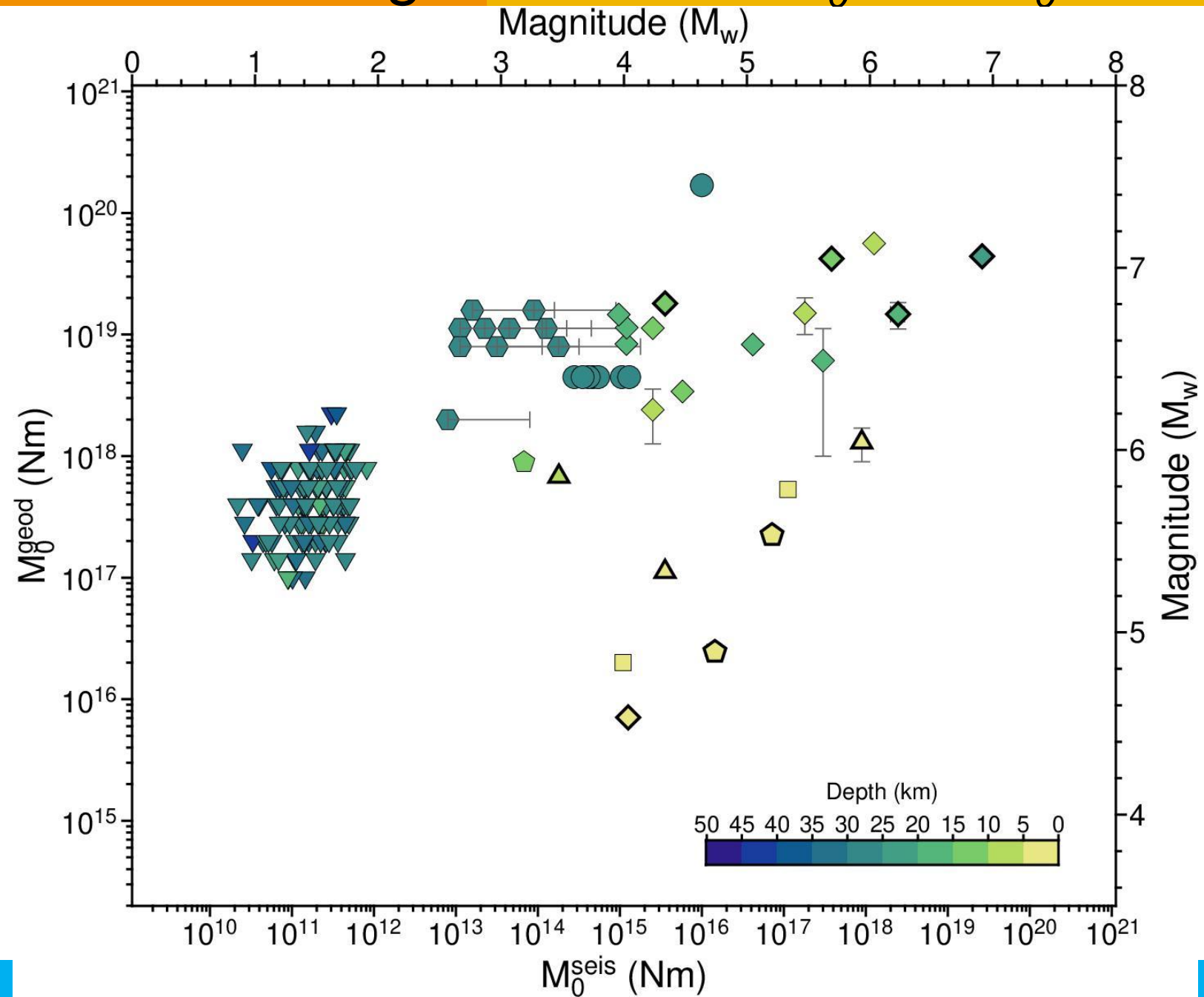


Fukuda, 2018

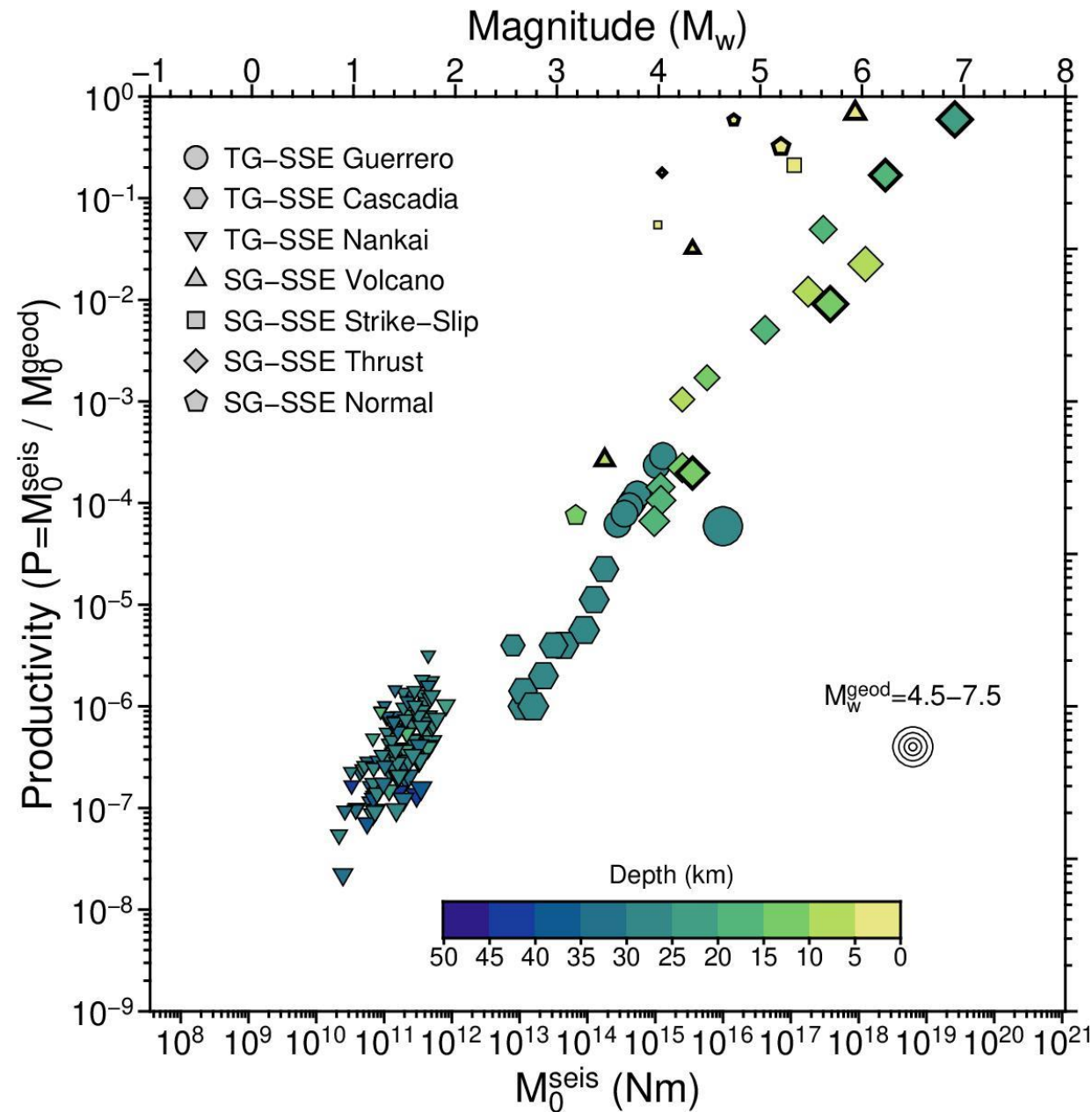


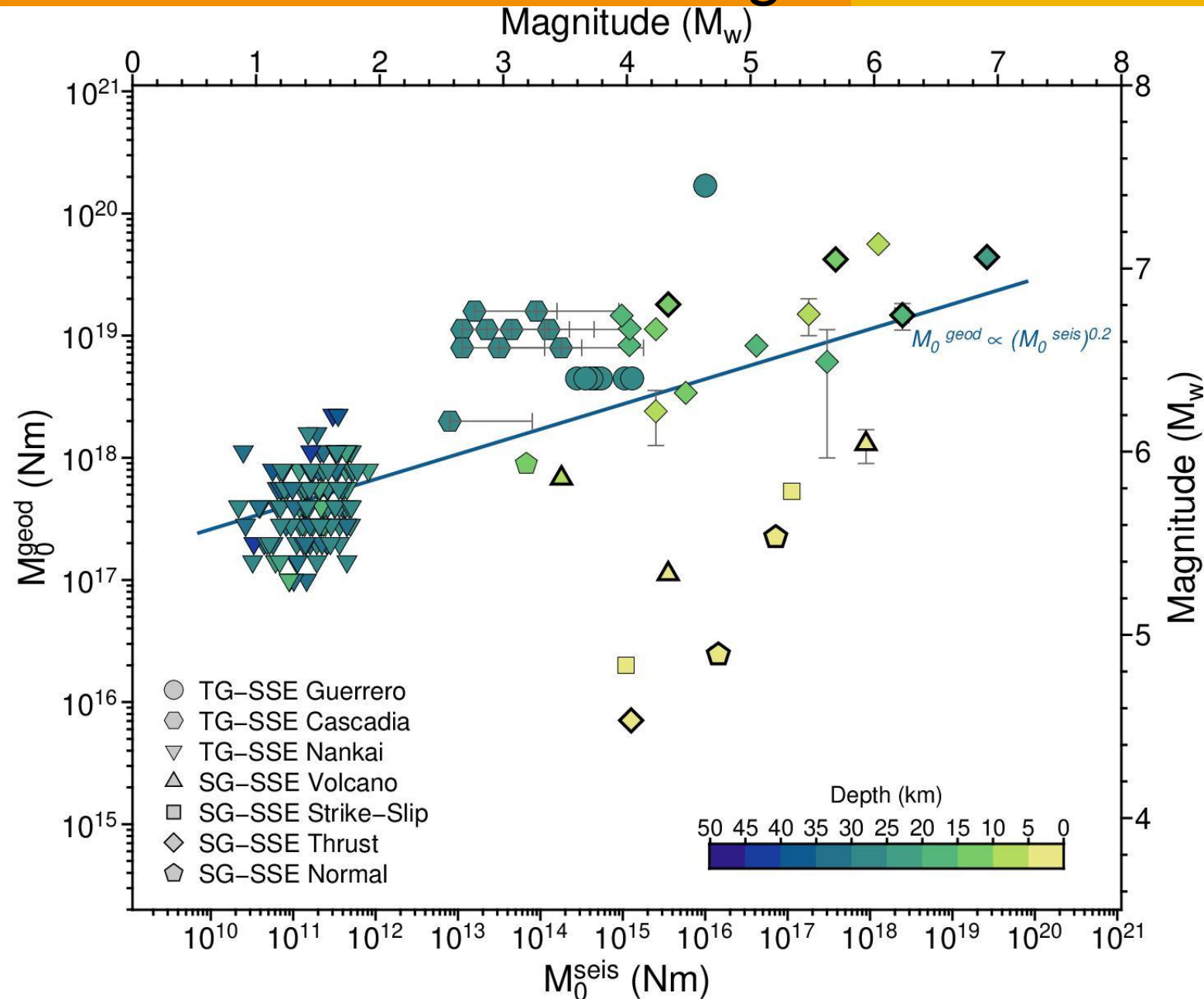
Itoh et al., 2022

- TG-SSE Guerrero
- TG-SSE Cascadia
- ▼ TG-SSE Nankai
- ▲ SG-SSE Volcano
- SG-SSE Strike-Slip
- ◆ SG-SSE Thrust
- ◐ SG-SSE Normal



Passarelli et al., 2021

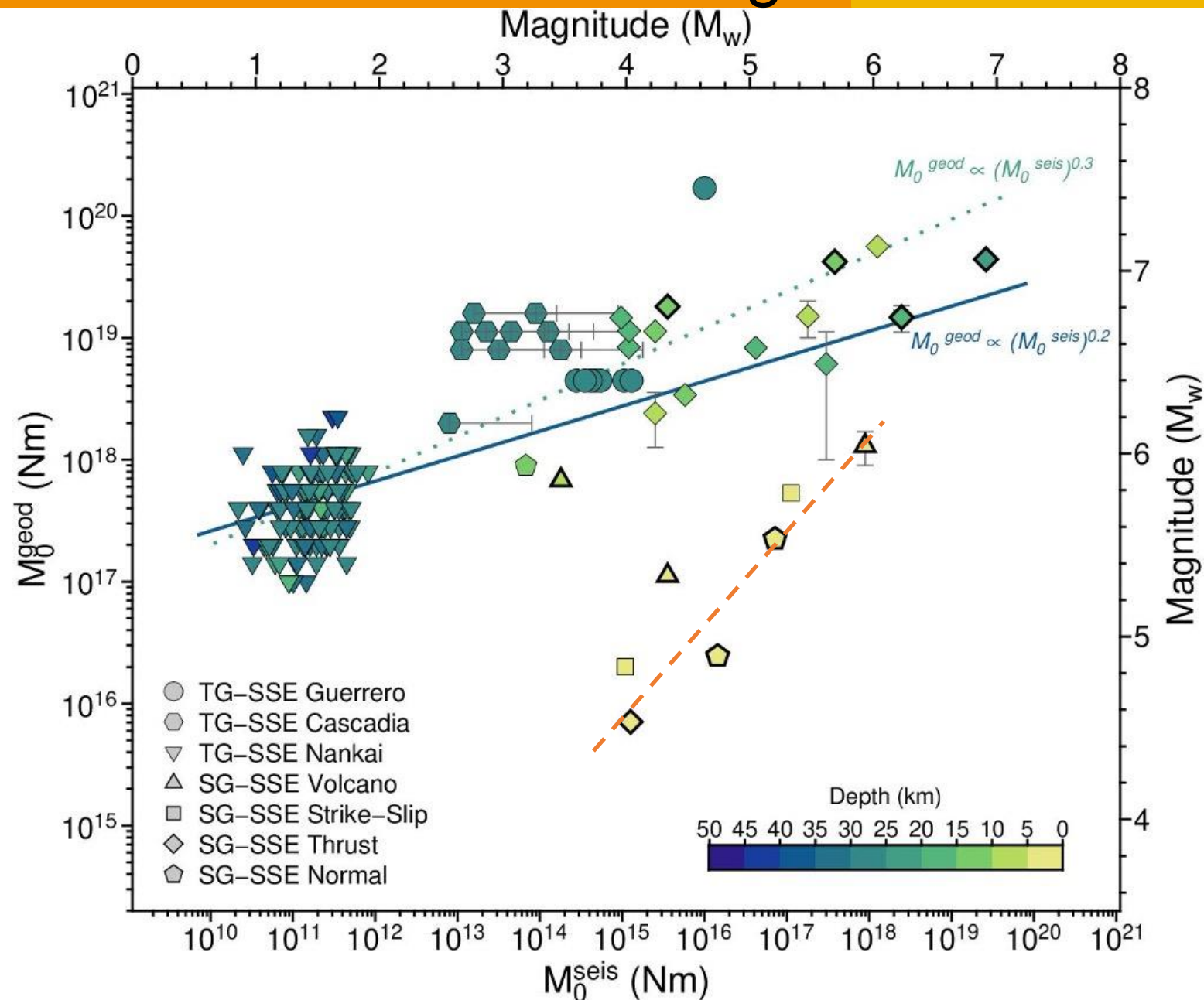




$$M_0^{geod} = \beta (M_0^{seis})^\alpha$$

$$\alpha = 0.2$$

Passarelli et al., 2021



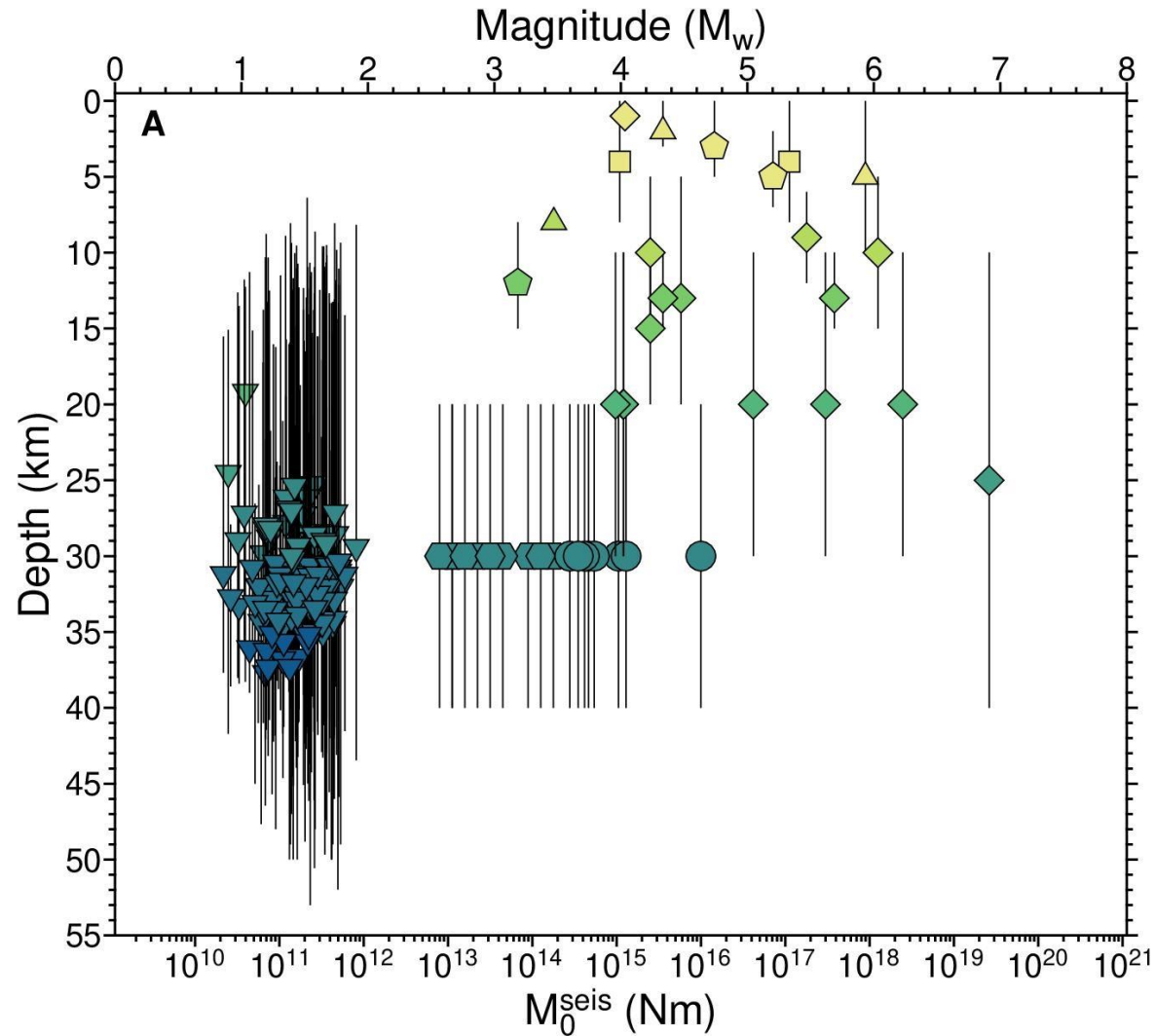
$$M_0^{geod} = \beta (M_0^{seis})^\alpha$$

$$\alpha = 0.3$$

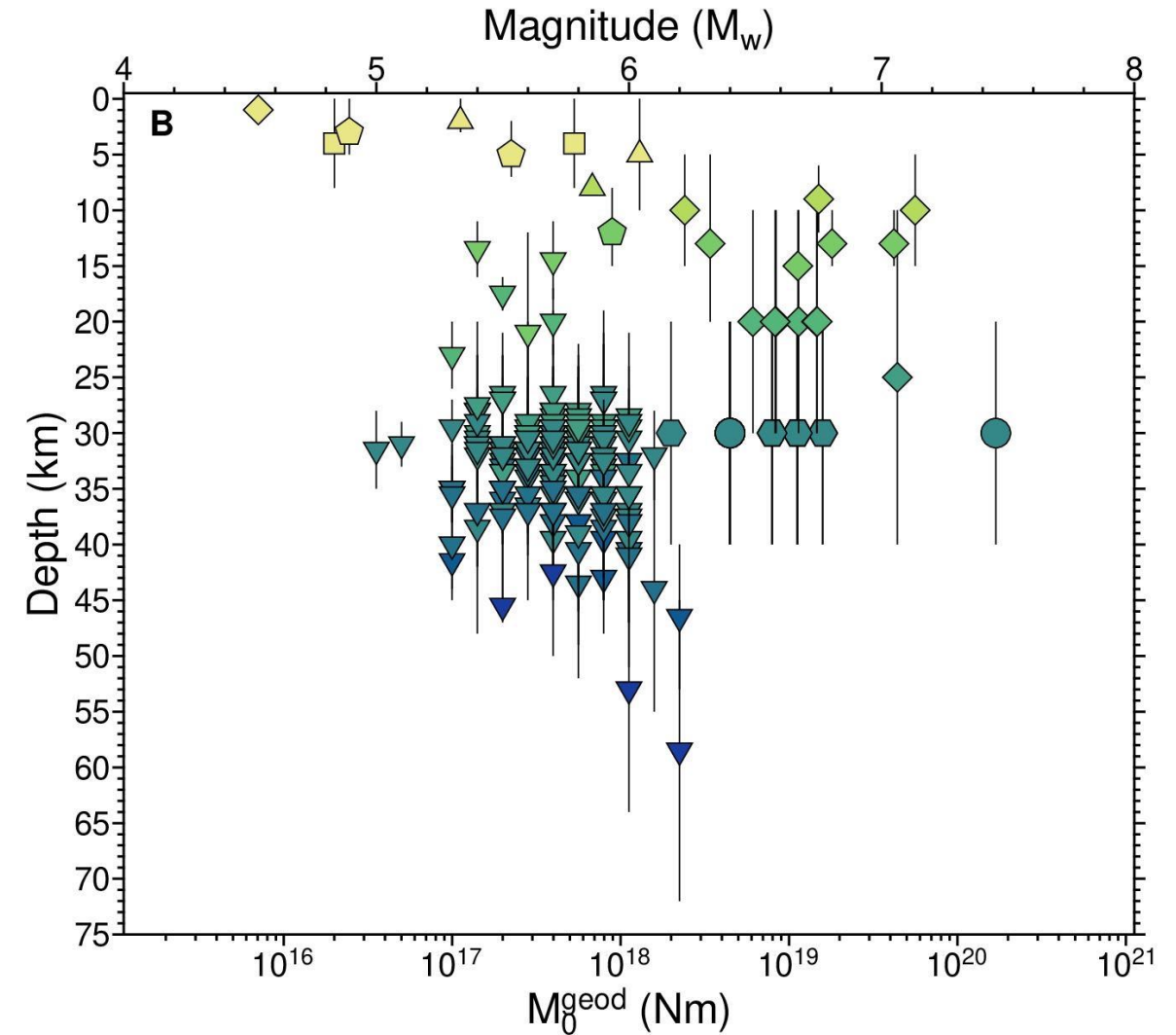
$$\alpha = 0.7$$

Passarelli et al., 2021

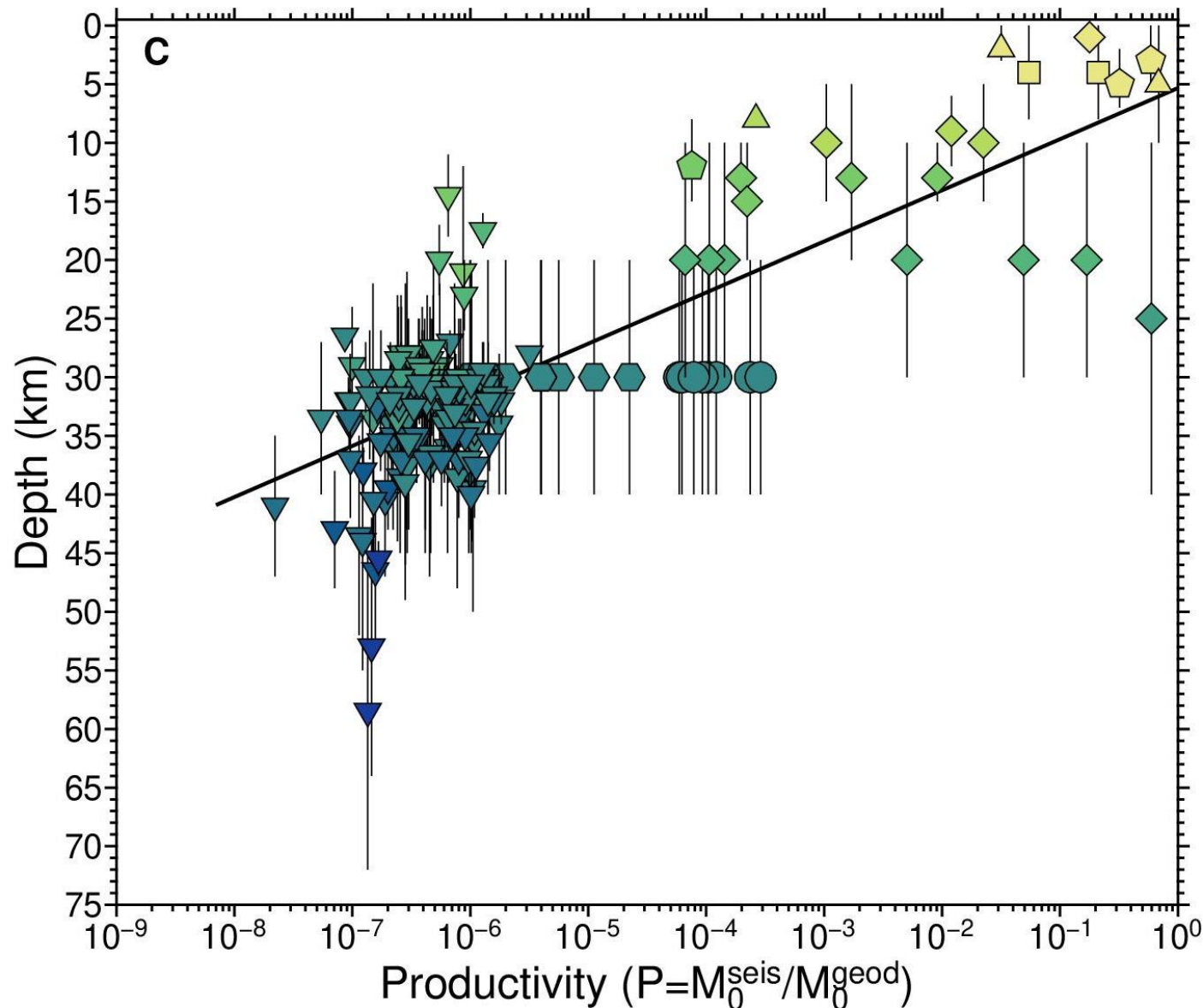
Seismic Moment



Geodetic Moment

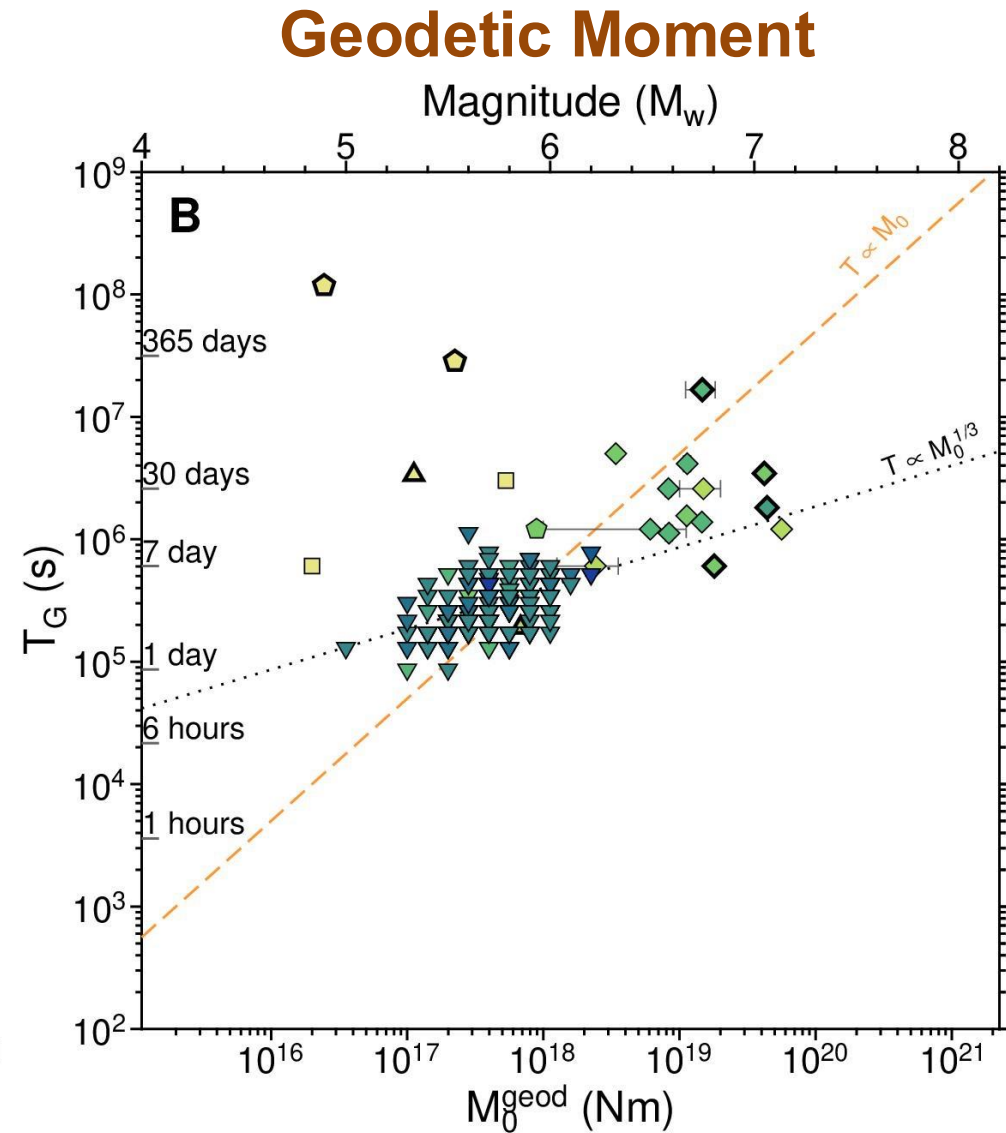
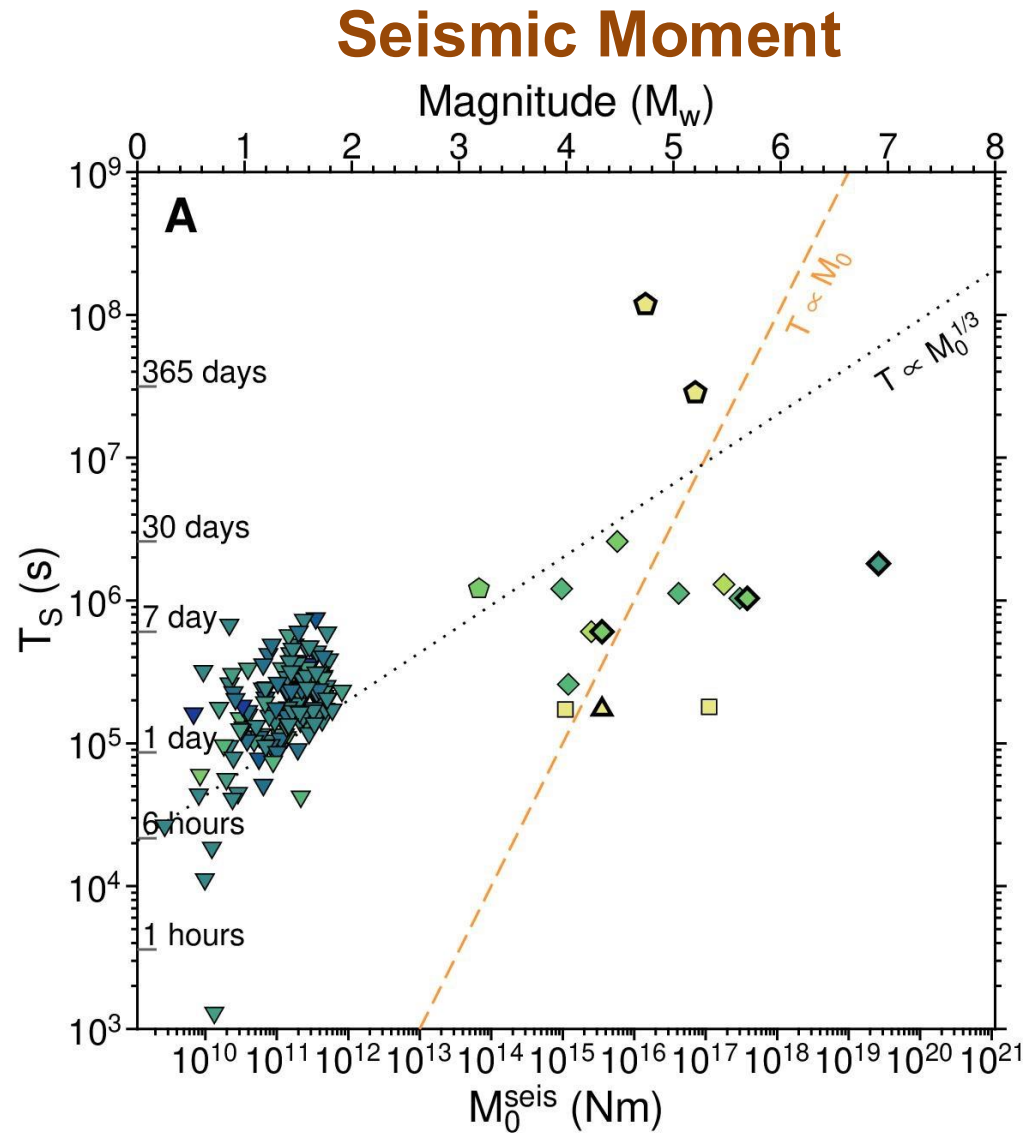


Passarelli et al., 2021

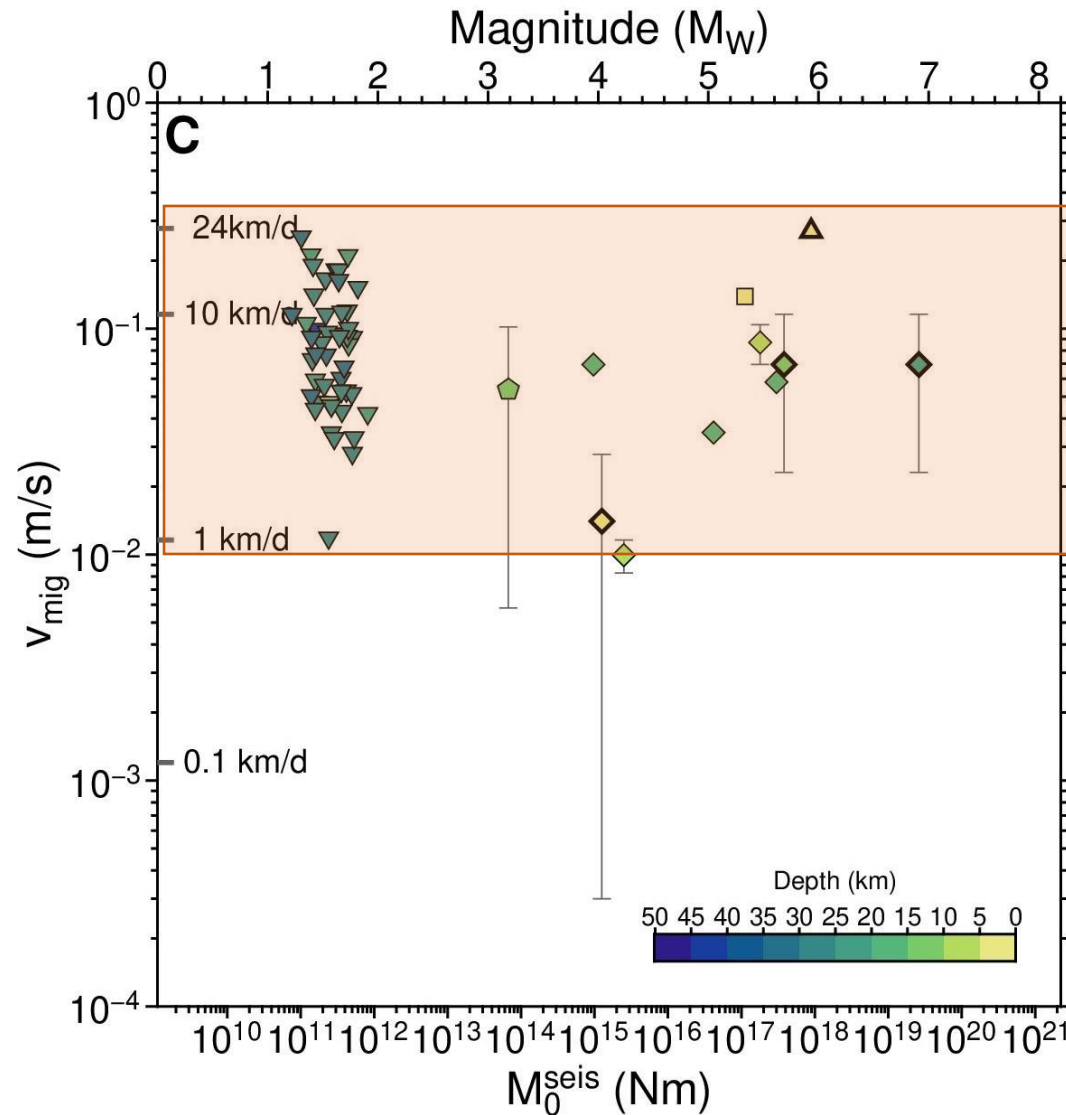


$$P = \frac{M_0^{seis}}{M_0^{geod}} \propto Depth$$

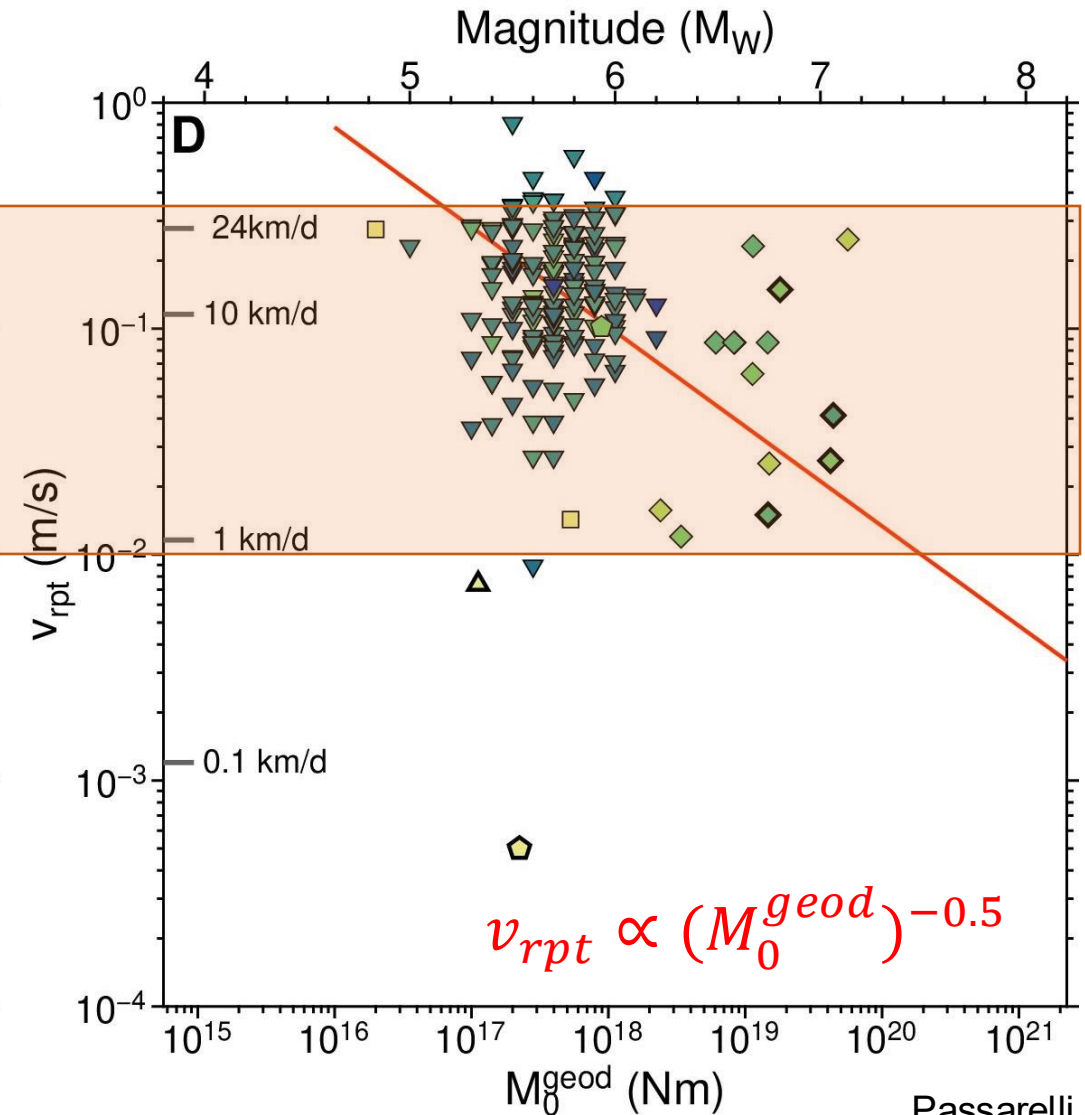
Passarelli et al., 2021



Seismic Moment

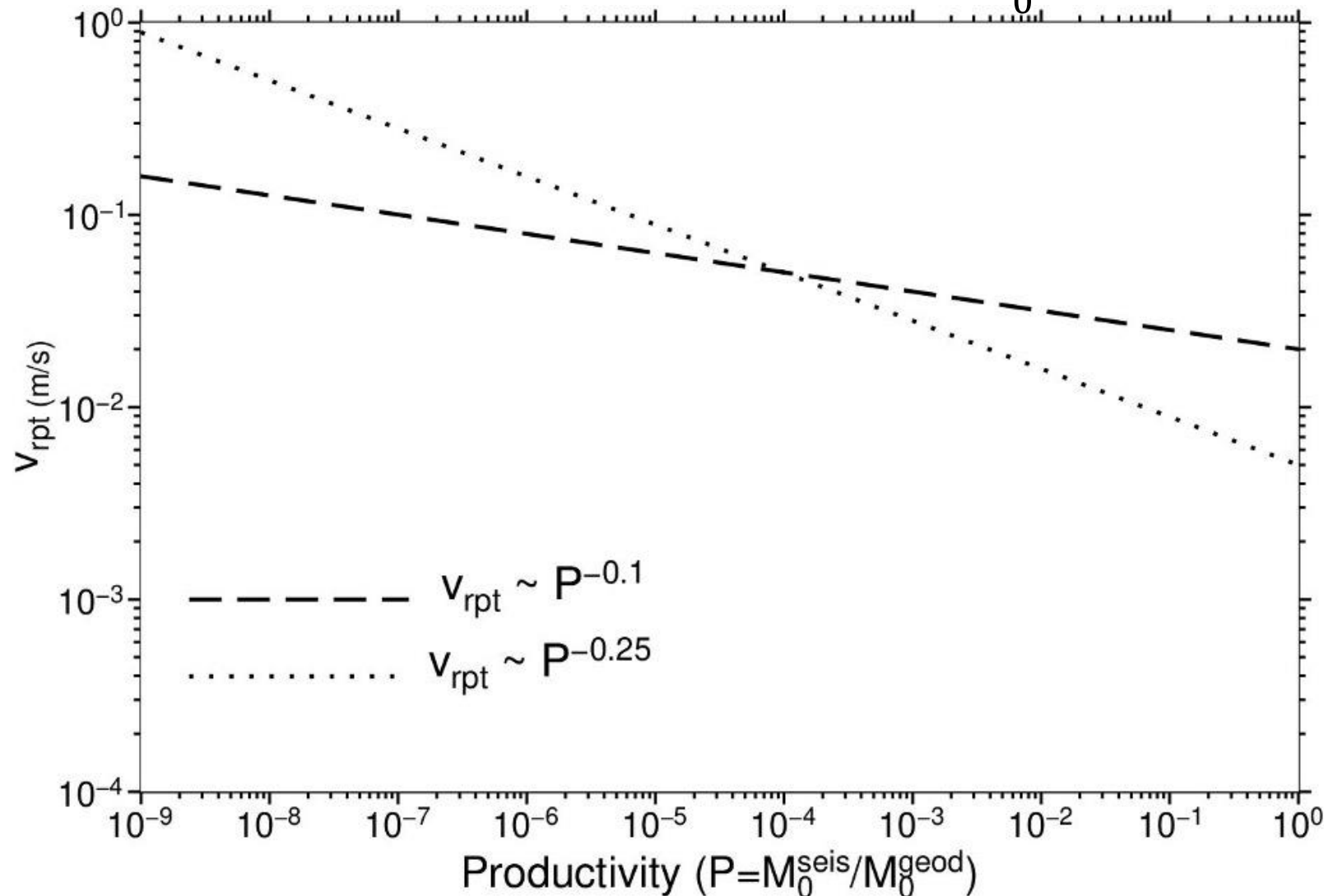


Geodetic Moment



Passarelli et al., 2021

$$v_{rpt} \propto (M_0^{geod})^{-\delta} \quad + \quad P(M_0^{geod}) = \frac{M_0^{seis}}{M_0^{geod}} \propto (M_0^{geod})^{\frac{1-\alpha}{\alpha}}$$



$$v_{rpt} \propto P^{-\gamma}$$

$$\gamma = \frac{\delta(1-\alpha)}{\alpha}$$

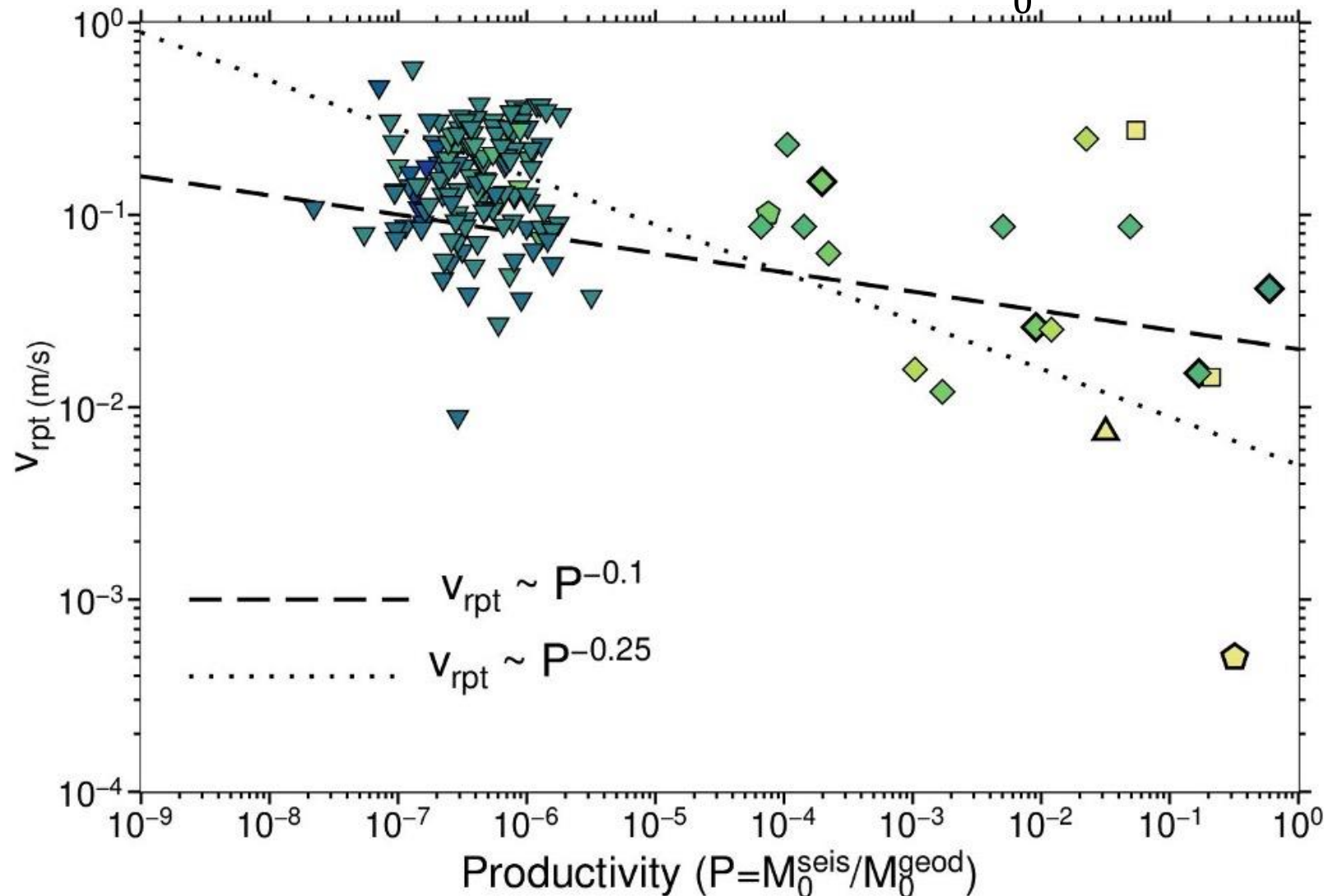
if

$\alpha = (0.2-0.3)$ (*this study*)

$\delta = (0.4, 0.6)$ (*Gao et al., 2010*)

$$\gamma = (0.1, 0.25)$$

$$v_{rpt} \propto (M_0^{geod})^{-\delta} \quad + \quad P(M_0^{geod}) = \frac{M_0^{seis}}{M_0^{geod}} \propto (M_0^{geod})^{\frac{1-\alpha}{\alpha}}$$



$$v_{rpt} \propto P^{-\gamma}$$

$$\gamma = \frac{\delta(1-\alpha)}{\alpha}$$

if

$\alpha = (0.2-0.3)$ (this study)

$\delta = (0.4, 0.6)$ (Gao et al., 2010)

$$\gamma = (0.1, 0.25)$$

SG- and TG-SSEs have robust moment scaling

Deeper $M_0^{geod} = \beta (M_0^{seis})^{0.3}$ lower P

Shallower $M_0^{geod} = \beta (M_0^{seis})^{0.7}$ higher P

Earthquake productivity P scale consistently with *depth*

$$P = \frac{M_0^{seis}}{M_0^{geod}} \propto Depth$$

Data are inconclusive for $M_0^{geod} - T_G$ scaling

$$M_0^{geod} - T^3 \text{ or } M_0^{geod} - T ??$$

No $M_0^{seis} - T_s$ (duration of seismicity) scaling

$$M_0^{seis} \not\propto T_s^n$$

No $v_{mig} - M_0^{seis}$ scaling

$$v_{mig} \not\propto M_0^{seis}$$

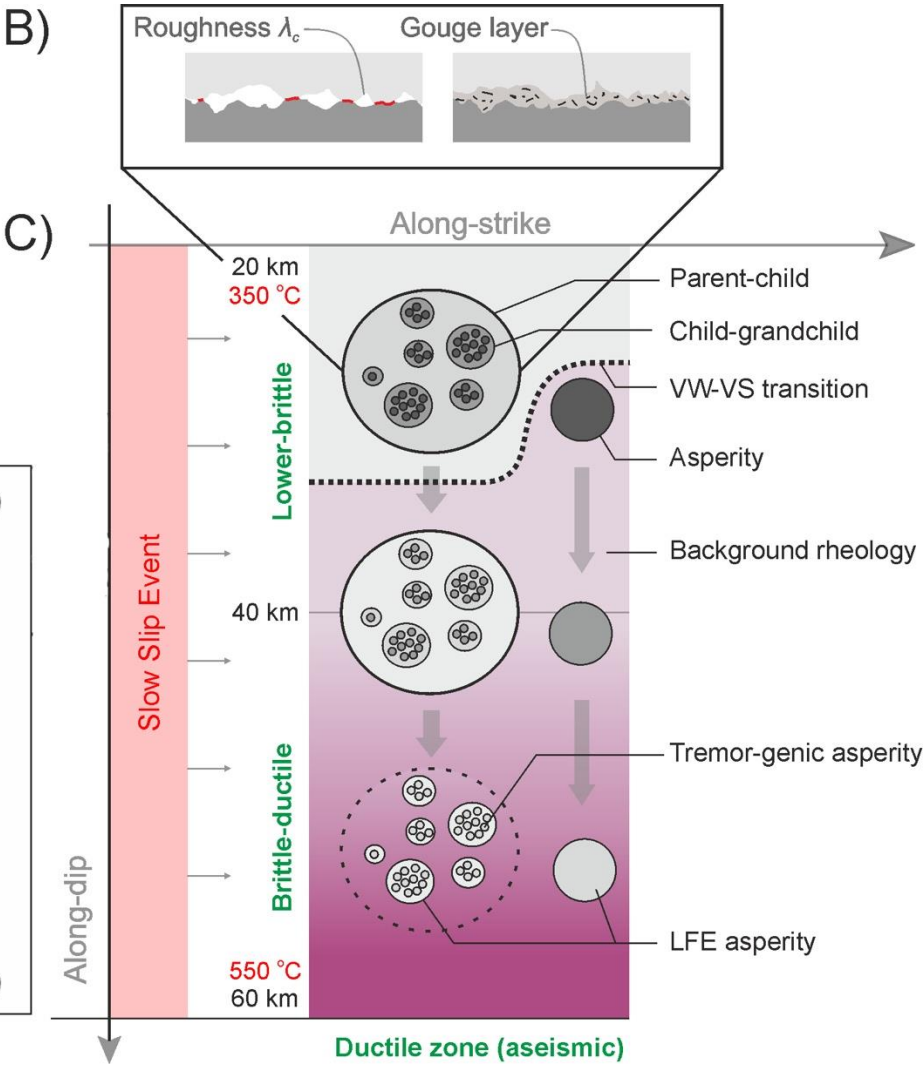
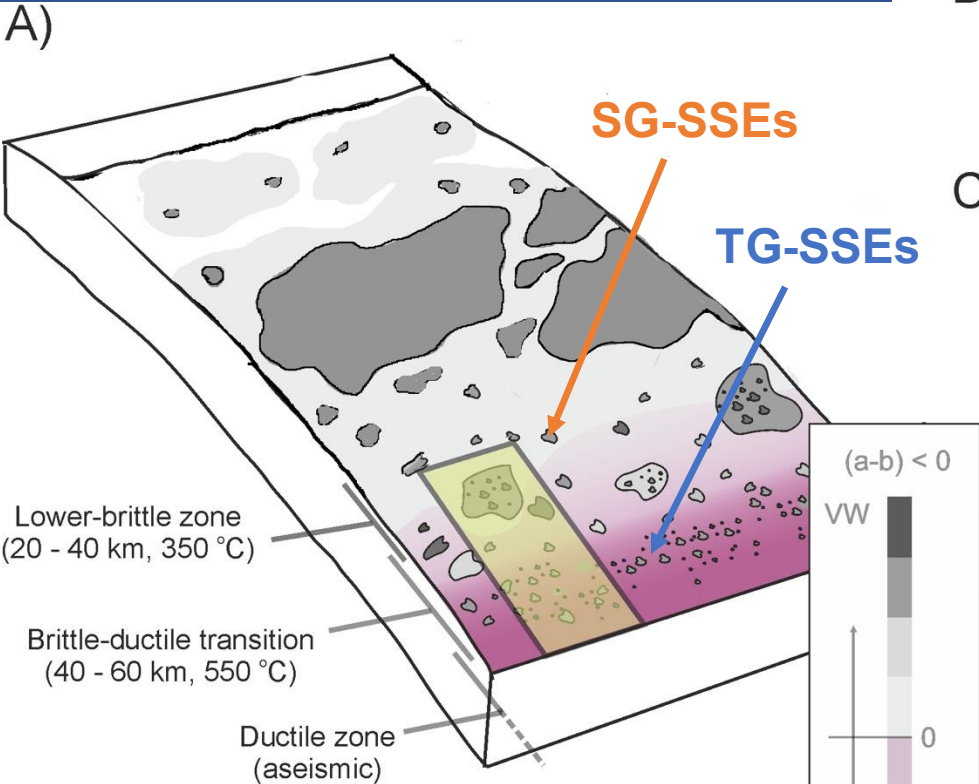
Our data compatible w/ the $v_{rpt} - M_0^{geod}$ scaling

$$v_{rpt} \propto (M_0^{geod})^{-0.5}$$

Earthquake productivity P scale consistently w/ v_{rpt}

$$v_{rpt} \propto P^{-\gamma}$$

Deeper $M_0^{geod} = \beta (M_0^{seis})^{0.3}$



Pronounced fault roughness

High pore pressure

Compatible w/

$$P = \frac{M_0^{seis}}{M_0^{geod}} \propto Depth$$
$$v_{rpt} \propto P^{-\gamma}$$

Heterogeneous frictional properties dependent on temperature

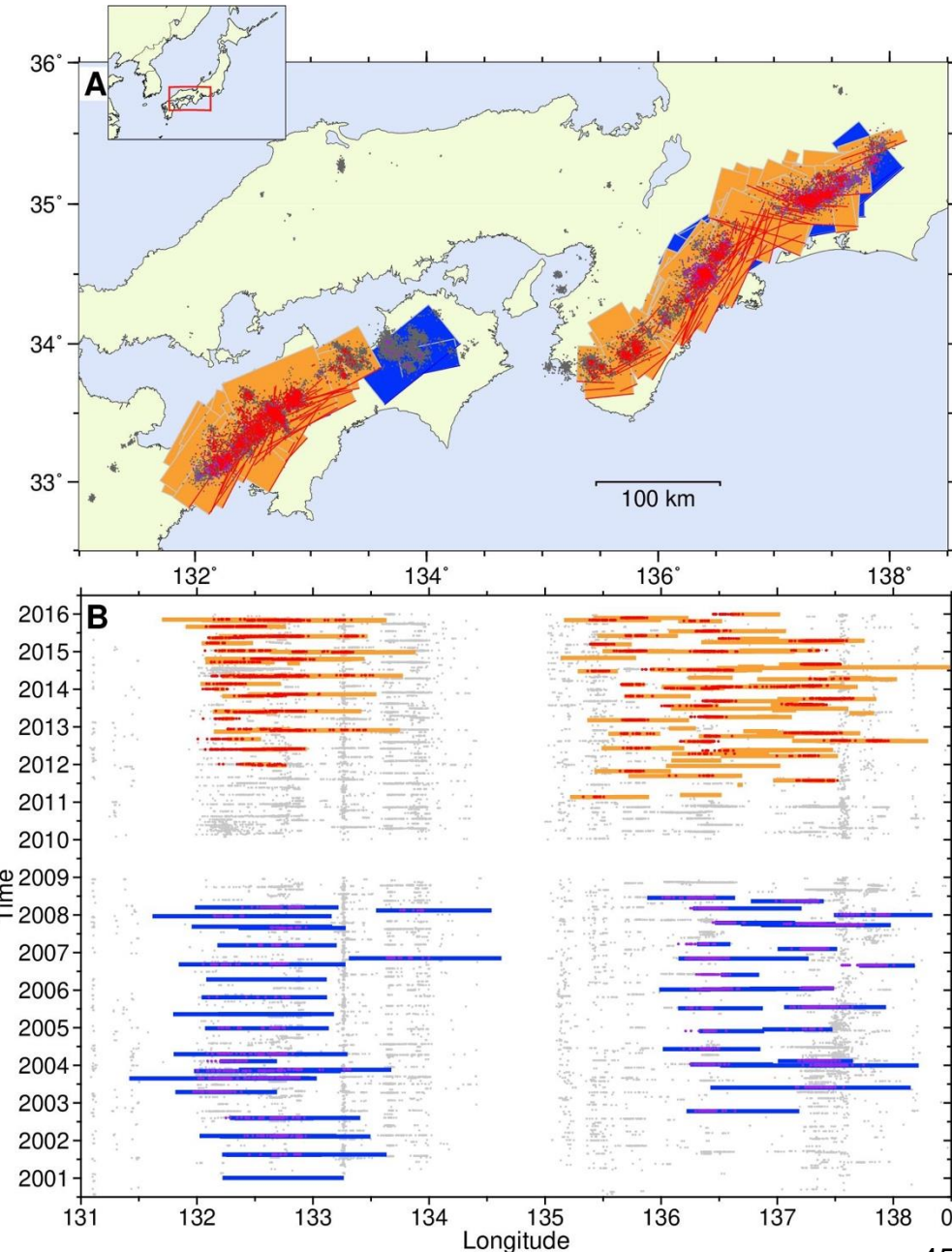
TG-SSEs:

- 10 SSEs at Cascadia subduction zone (*literature*)
- 8 SSEs at Guerrero subduction zone (*literature*)
- 174 SSEs at Nankai subduction zone (*database*)

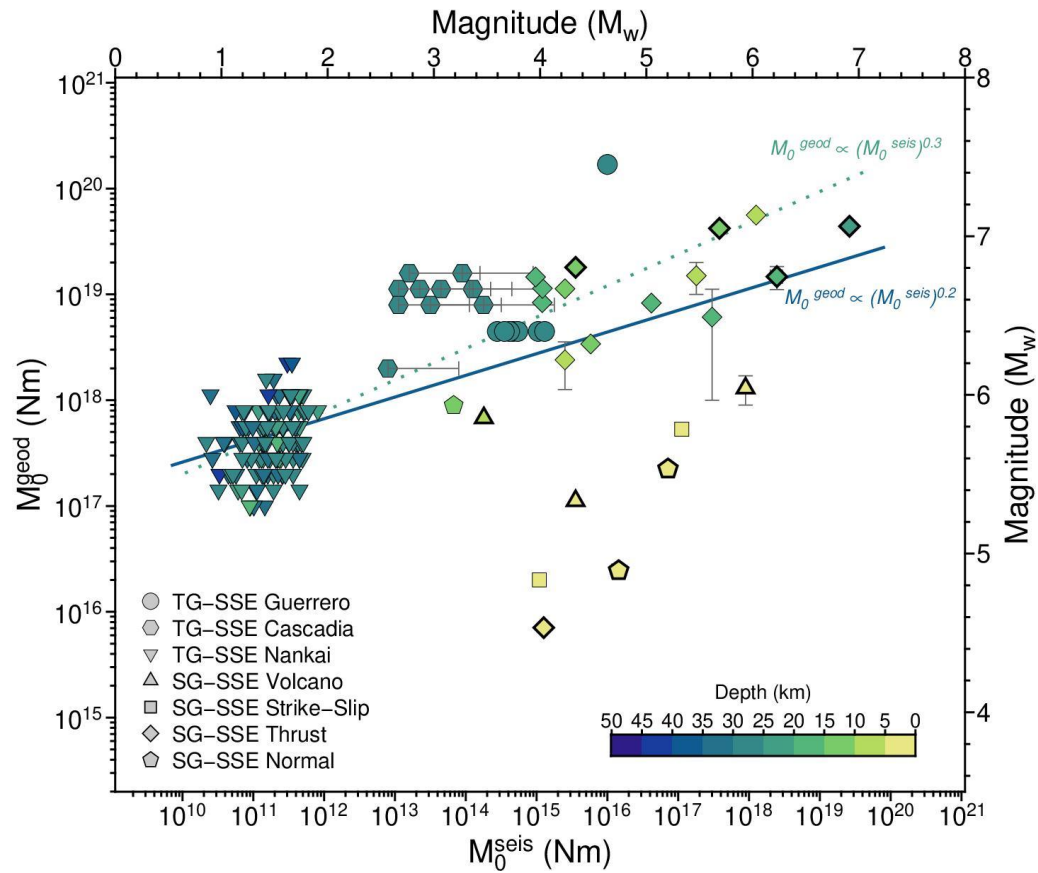
SG-SSEs:

- 3 SSEs at volcanoes (Kilauea, Mt Etna and Miyake)
- 2 Strike-slip fault (San Andreas and Salton Throat)
- 3 Normal fault (Pollino, Alto Tiberina and L'Aquila)
- 1 Thrust (Wood Island Washington State) (*literature*)
- 14 Thrust in subduction zone (Ecuador, Chile, Japan)

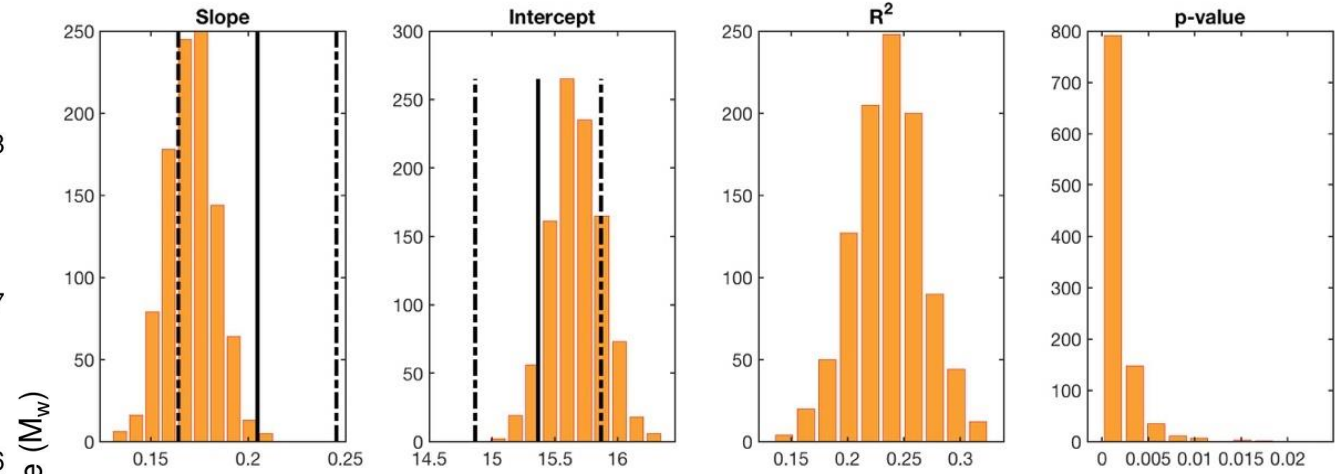
Database dominated by events in subduction zone



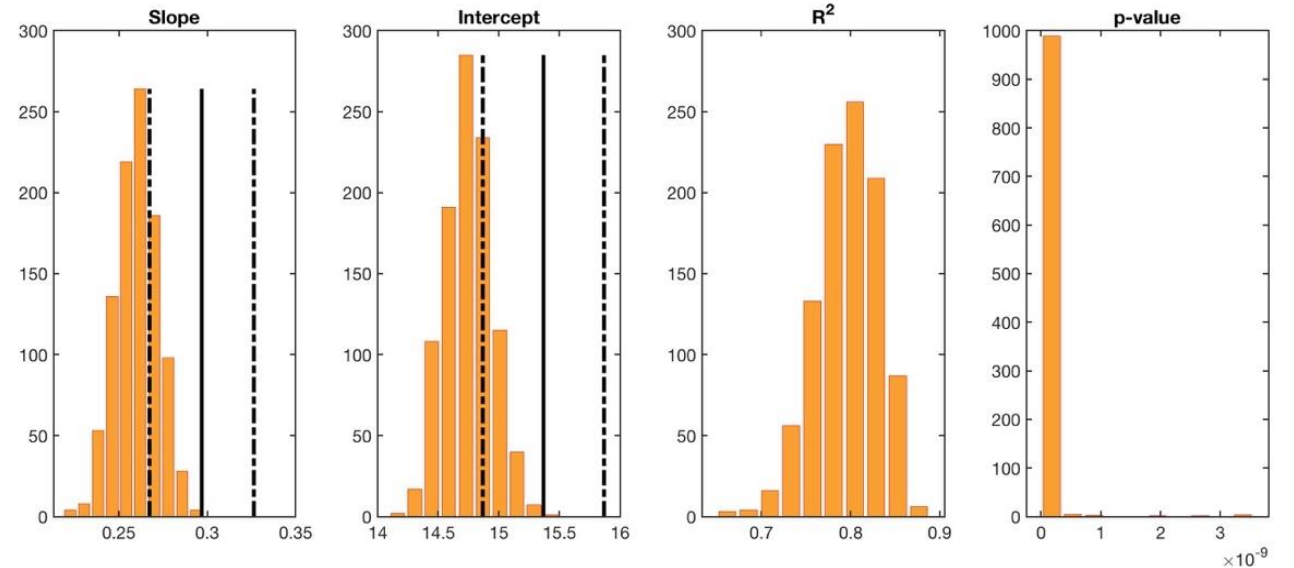
Statistical test



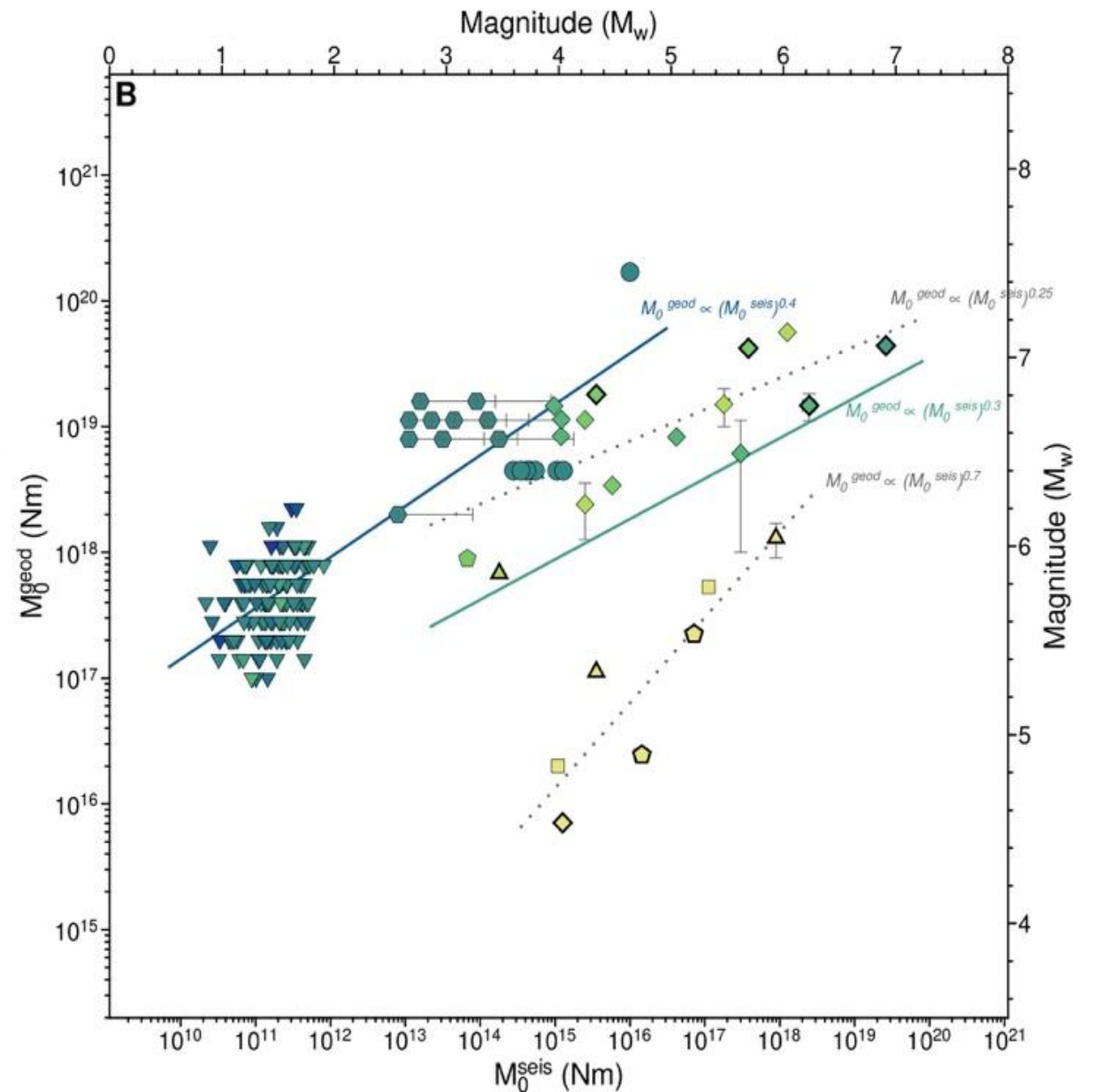
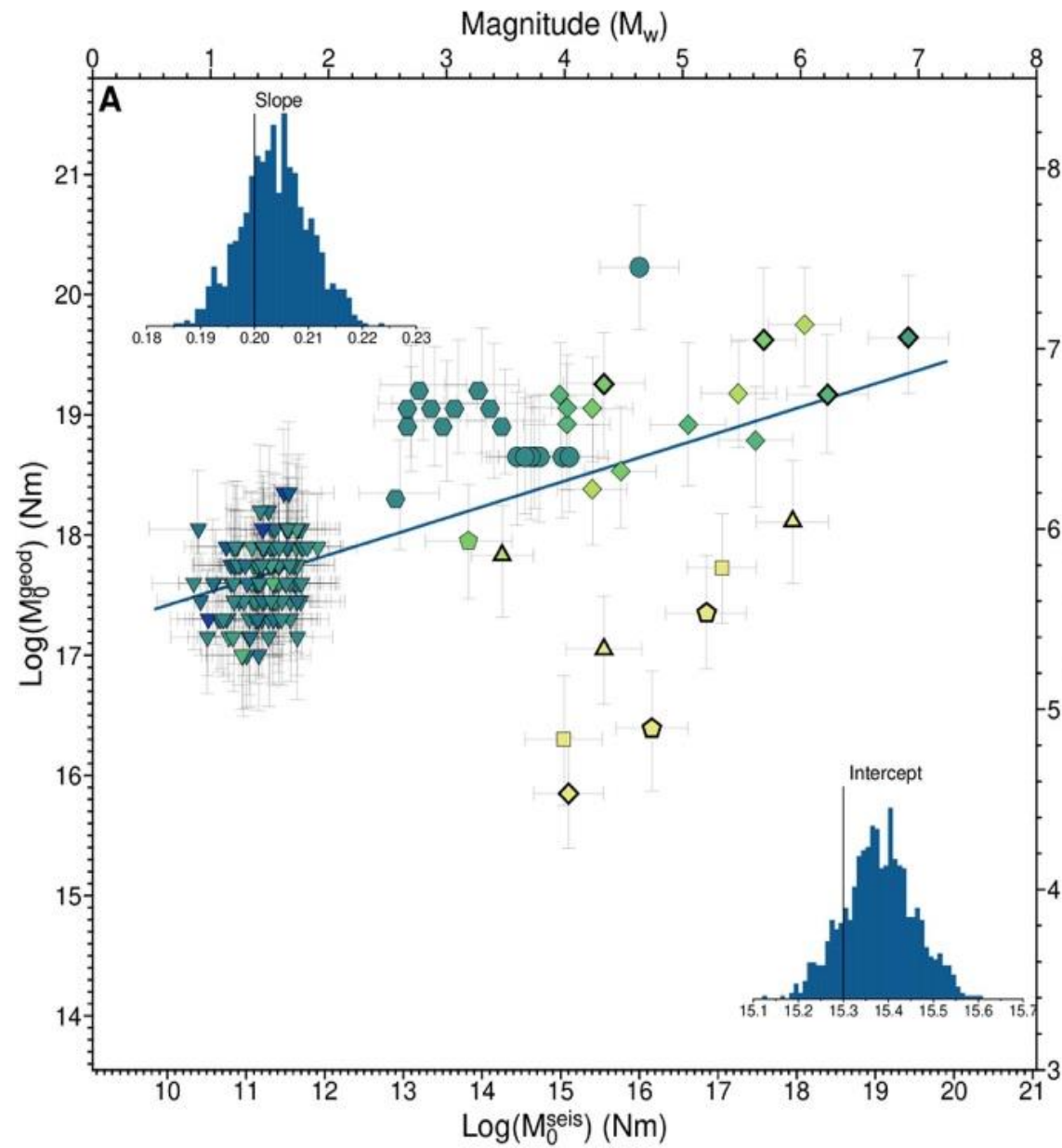
All data

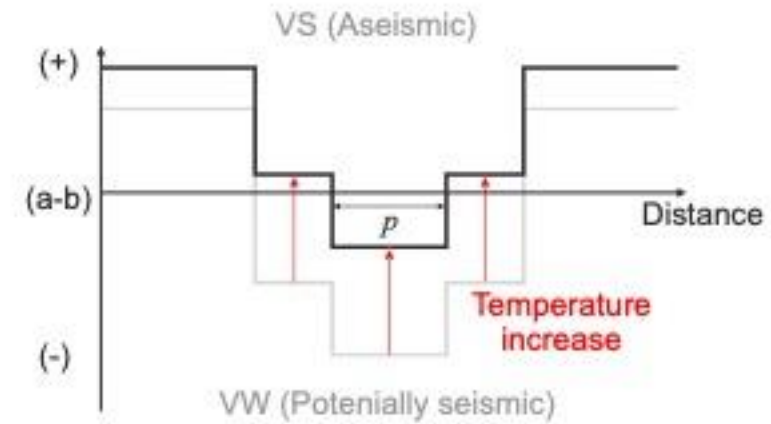
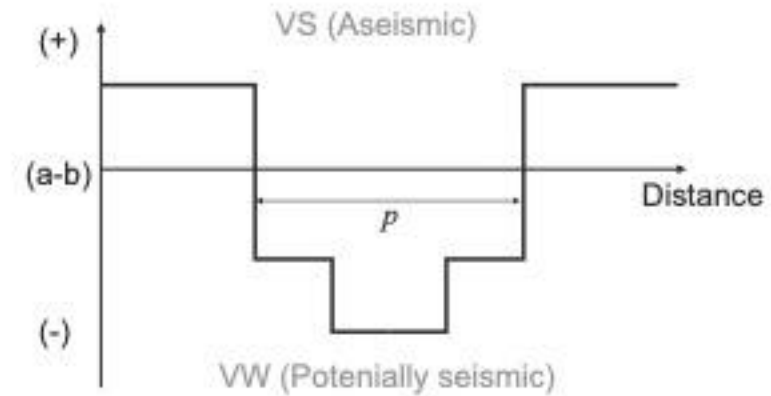
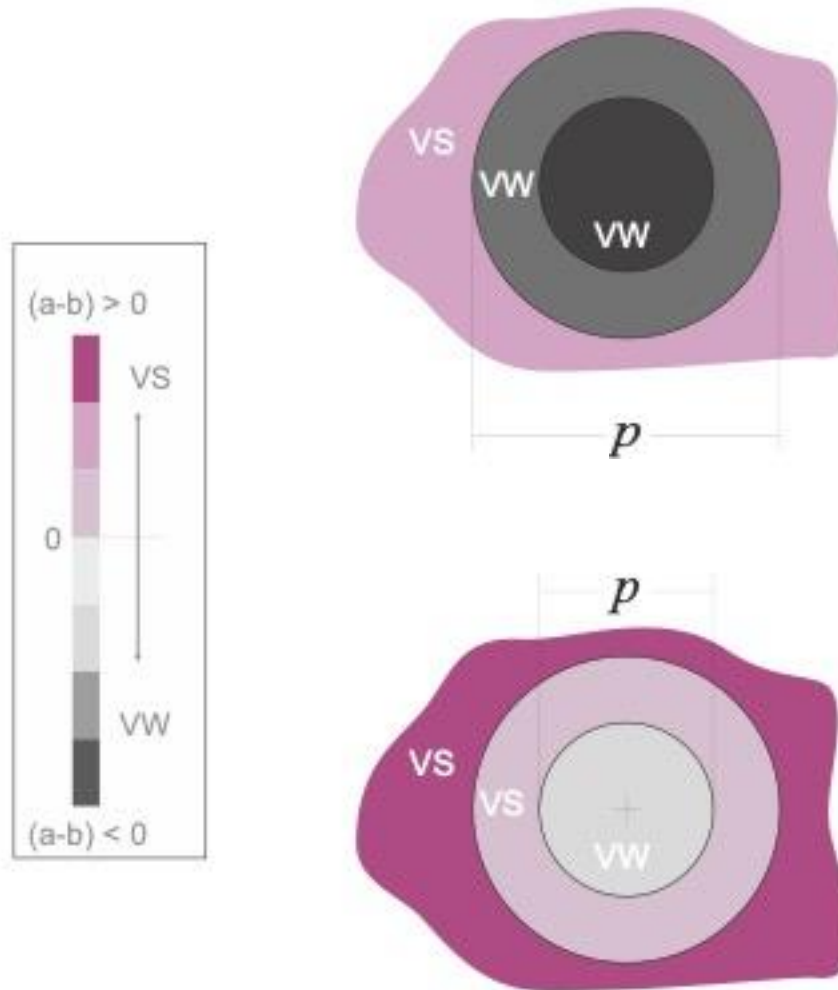


No shallow SG-SSEs



Statistical test





$$h^* = \frac{\mu D_c}{(b-a)\sigma_{eff}}$$

