

A white waveform representing a microseismic event, centered horizontally across the slide. The waveform shows a series of peaks and troughs of varying amplitudes, typical of seismic data.

Understanding reservoir processes in injection operations from advanced microseismic analysis

Bettina Goertz-Allmann, Daniela Kuehn, Kamran Iranpour, Ben Dando, Volker Oye, Robert Bauer

NORSAR

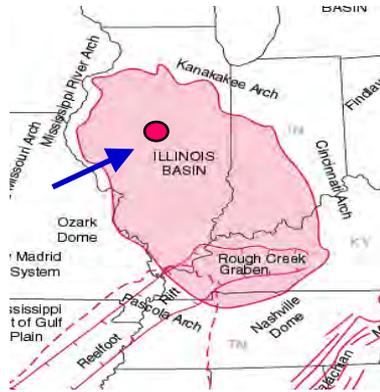
3rd Schatzalp workshop, Davos, March 8, 2019

Outline

- Reservoir characterization using microseismicity at the IBDP
 - Waveform cross-correlation and source analysis
 - Full-waveform modelling for hypothesis testing & interpretation
- Constraining event depth with examples from
 - In Salah
 - Oseberg
- Conclusions

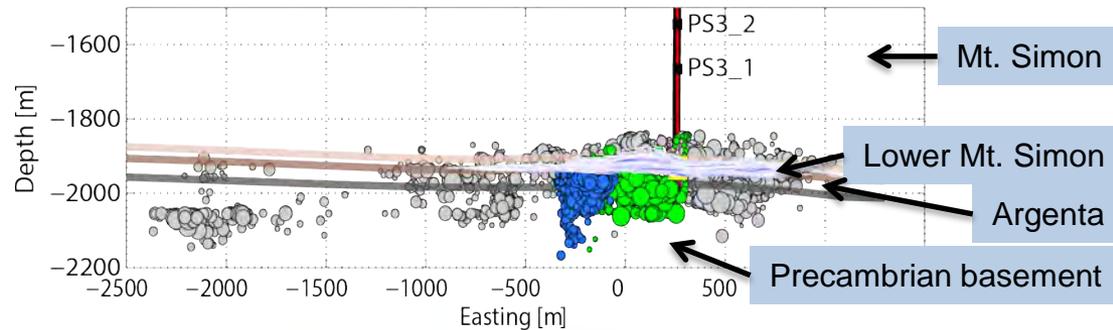
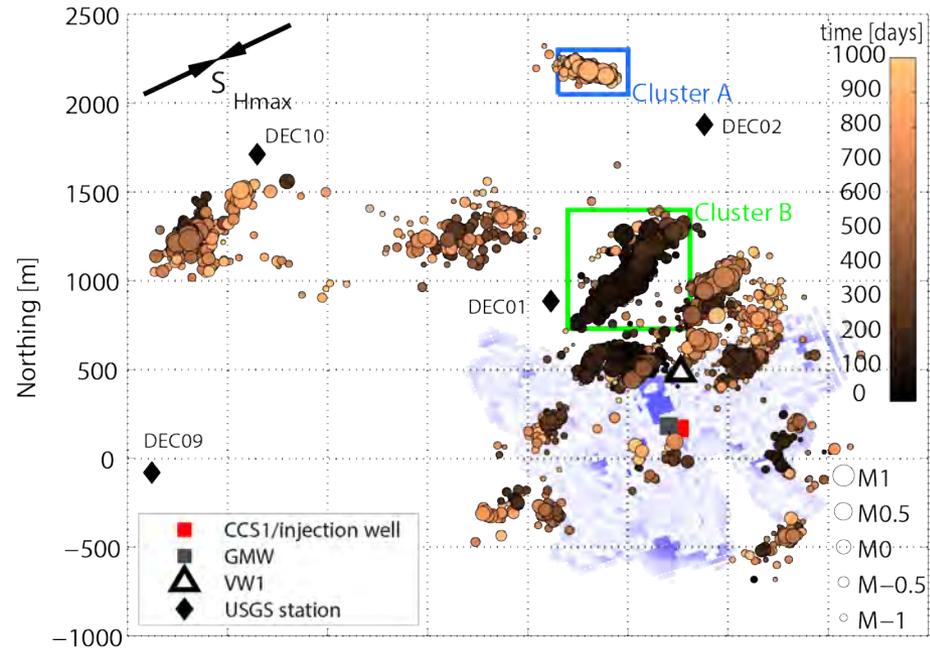


The IBDP CCS site



- Inject 1 MT of CO₂ into Mt. Simon sandstone at ~1.9 km depth over three years (end 2011-2014)
- Microseismic monitoring includes borehole & surface sensors
- Events occur in distinct clusters with heterogeneous timing

~ 4,800 microseismic events



Event characterization

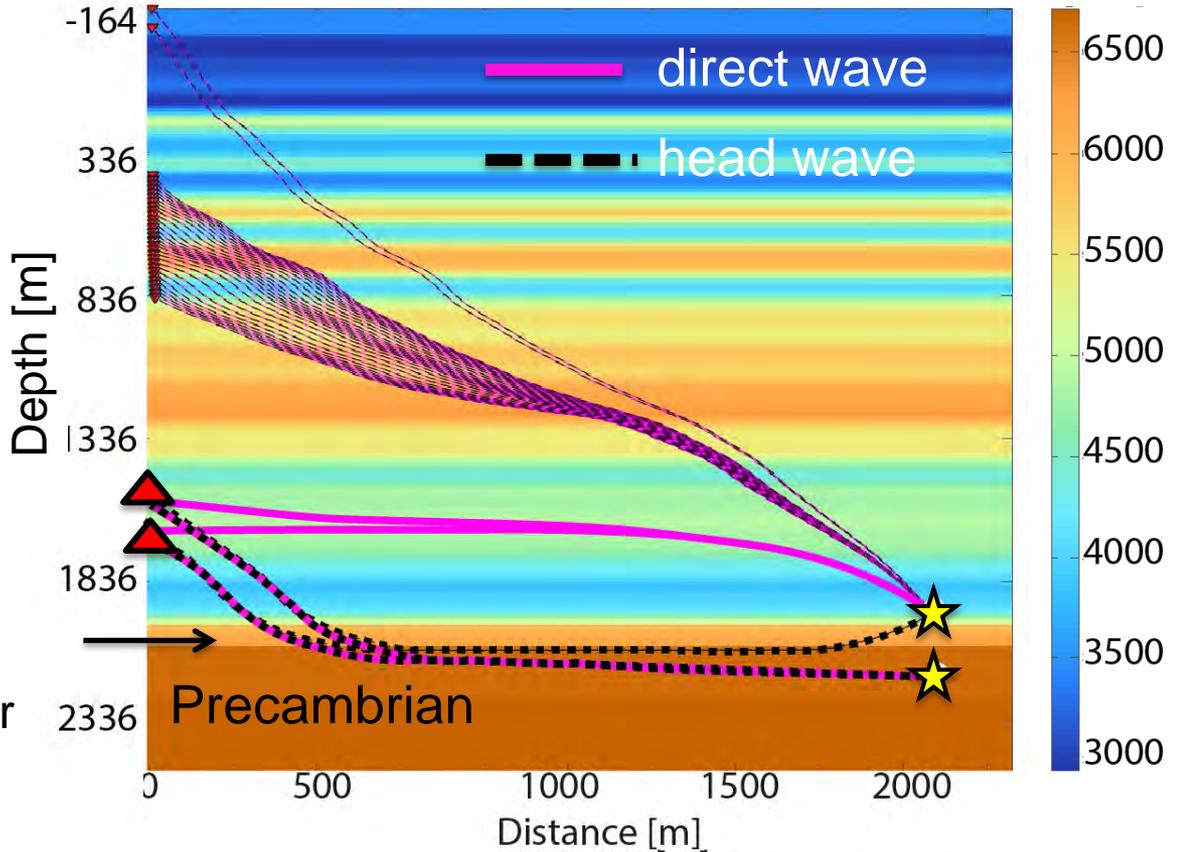
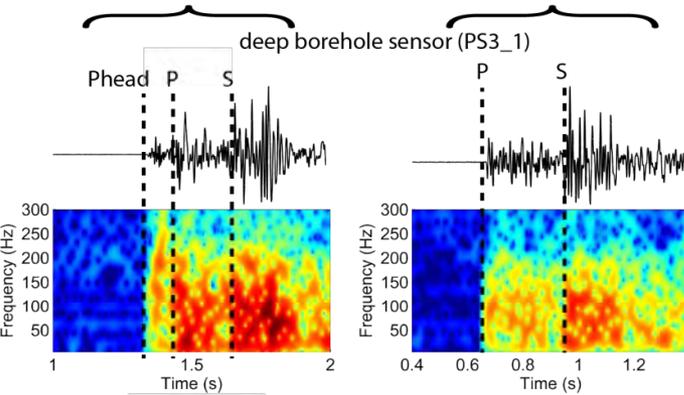
Goertz-Allmann et al. (2017), JGR

Basement vs. reservoir events

Vp [m/s]

Reservoir

Basement

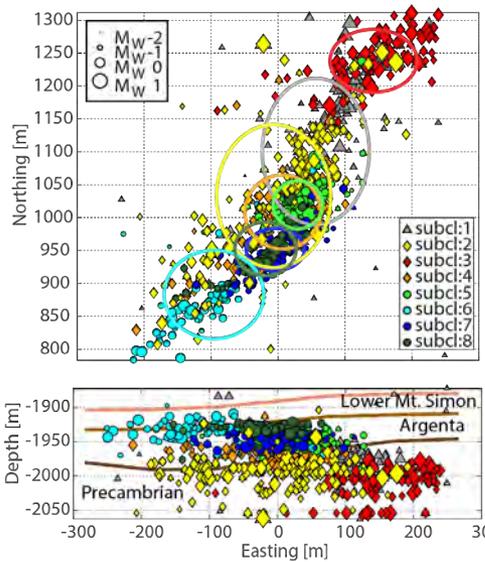


- Theoretical ray diagrams for reservoir & basement events.
- Different waveform signature: head wave and direct wave arrivals clearly visible for reservoir events.
- Events can be distinguished using waveform cross-correlation.

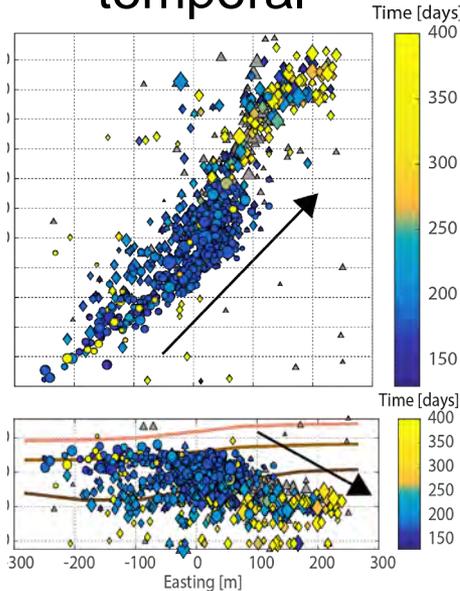


Event characterization

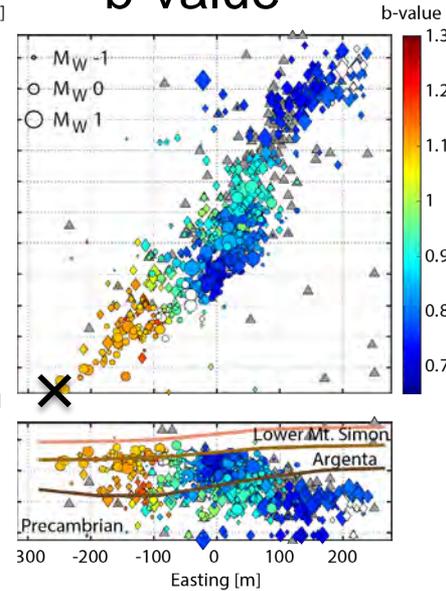
Cluster B spatial



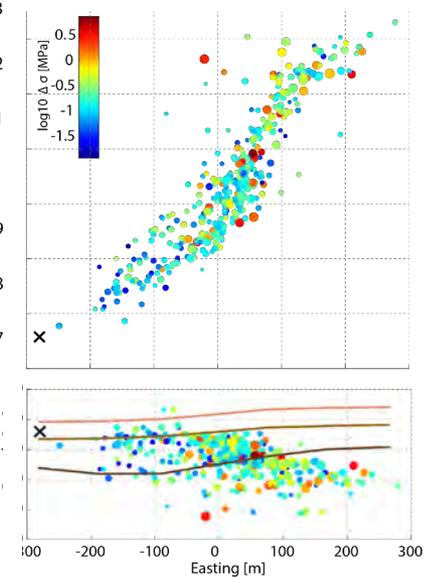
temporal



b-value



stress drop



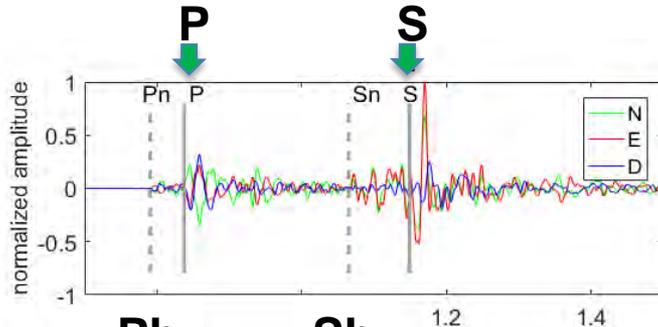
- Separation between reservoir and basement events.
 - Migration from the reservoir into the basement.
 - Decrease of b-value with distance.
 - Increase of stress drop with distance.
- Evidence for a fluid-driven process at the cluster level.
 - Possible punctual hydraulic connection between reservoir and basement (i.e., confined to faults).



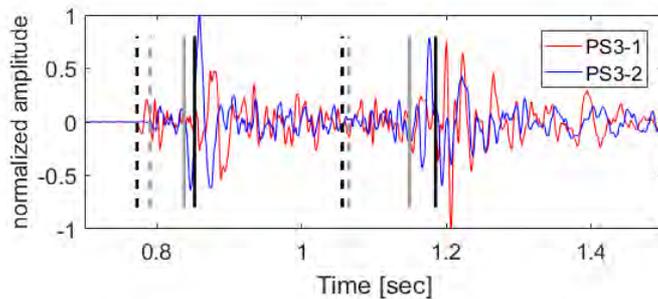
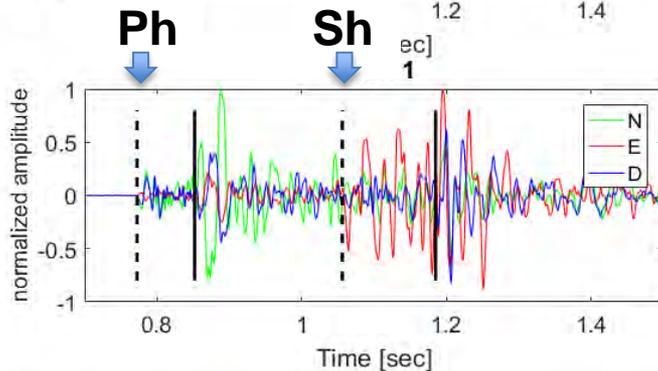
Full-waveform modelling

Observed waveform example

Shallower
PS3_2



Deeper
PS3_1

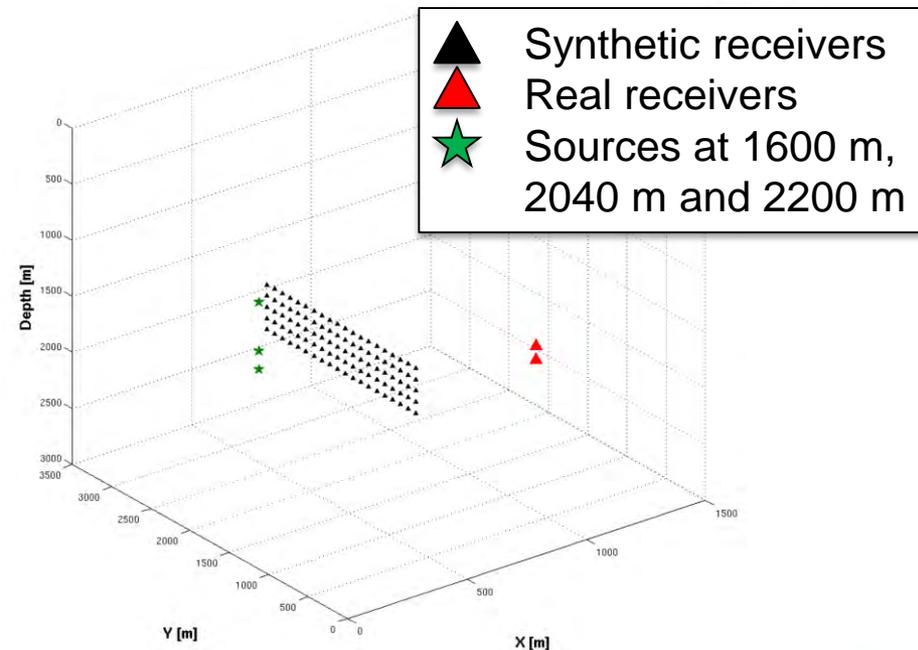
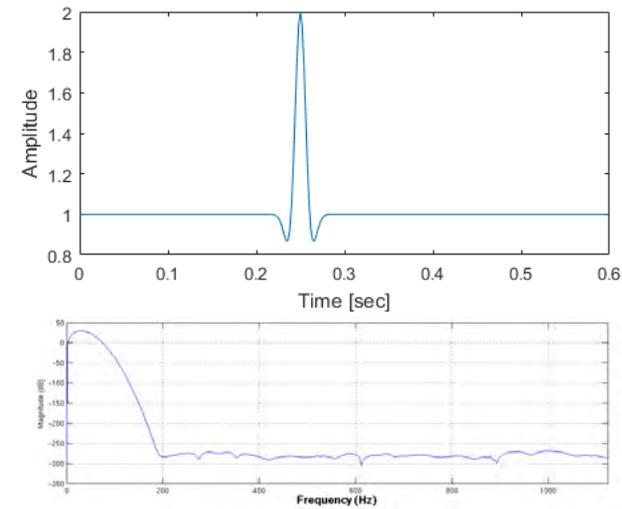
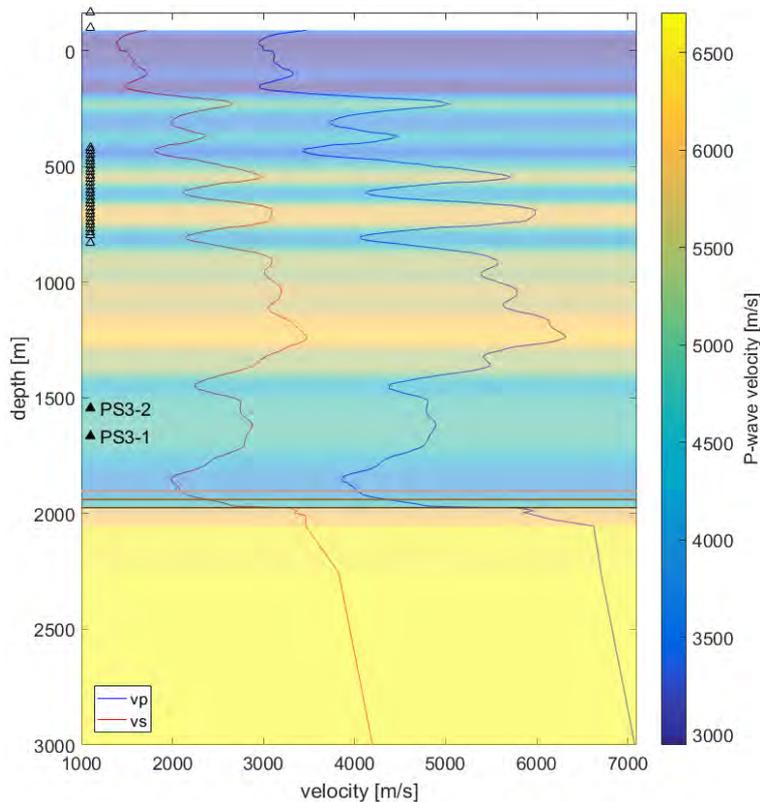


- Different phase arrivals with head wave and direct wave arrivals.
- Ph/Sh phase arrives first at deeper sensor (PS3_1).
- P/S phase arrives first at shallower sensor (PS3_2).
- Waveform modelling can help us to better understand the observed waveform characteristics.
- Gain a complete picture of the travel path of an event and helps us to select events and phases, which best sample the target area.



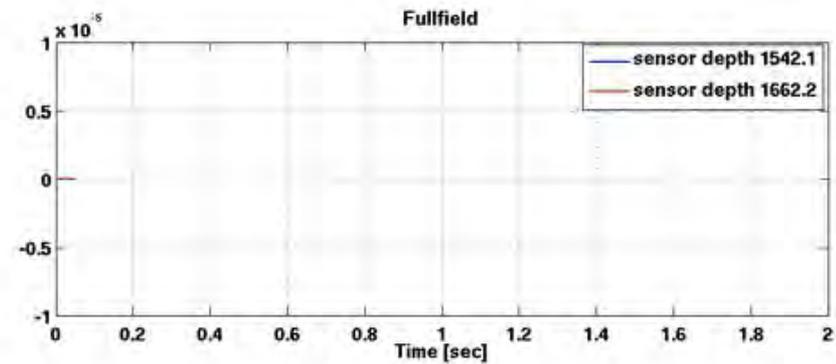
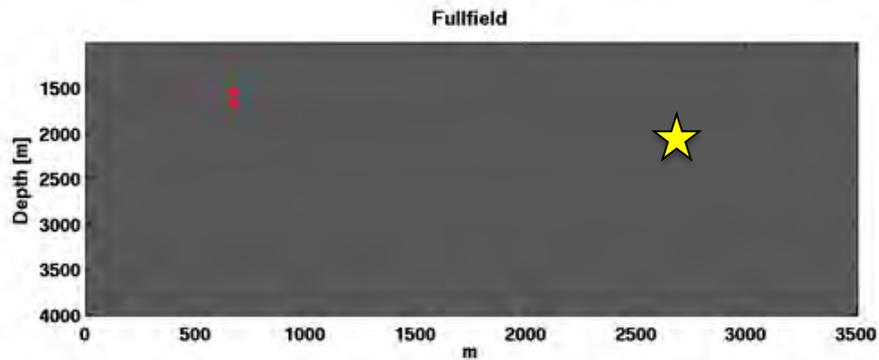
Full-waveform modelling

- 3D FD modelling using 1D velocity model
- 30 Hz Ricker wavelet.
- Compare sources placed at 1600 m, 2040 m, and 2200 m depth.



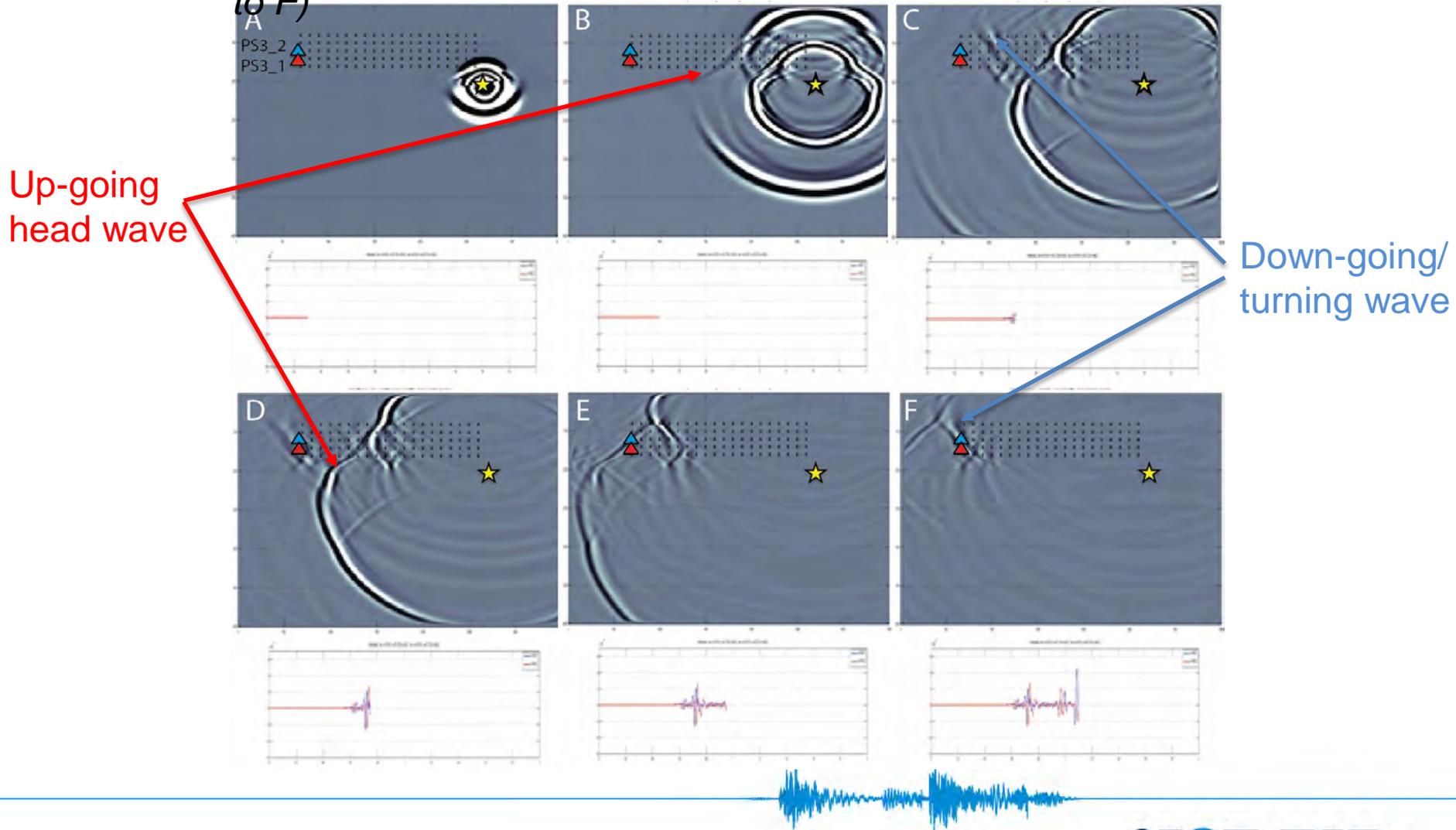
Full-waveform modelling

Source at 2040 m depth



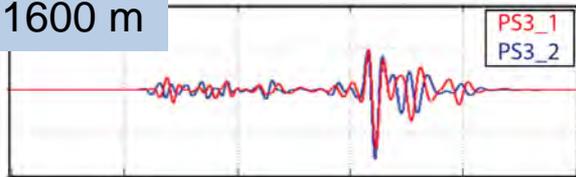
Full-waveform modelling

Sequential snap-shots of full waveform modelling (from A to F)

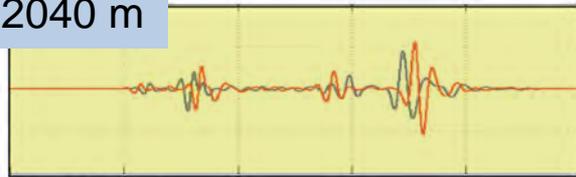


Full-waveform modelling modelled

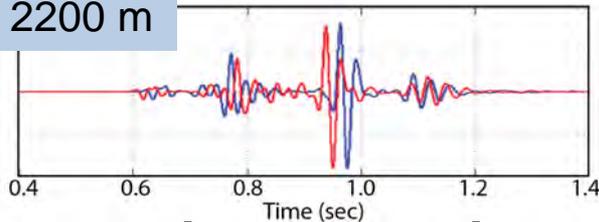
Source 1600 m



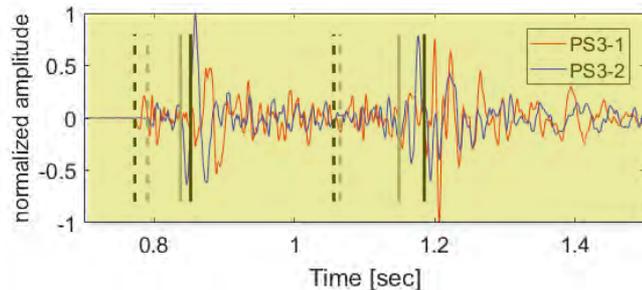
Source 2040 m



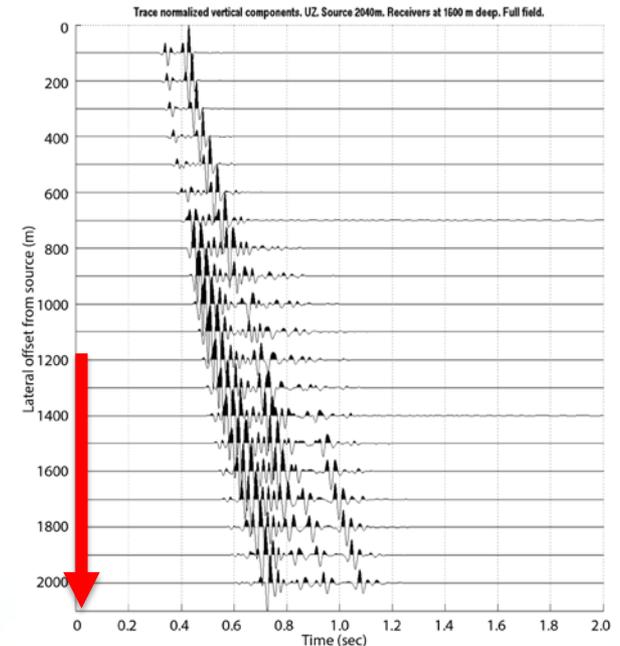
Source 2200 m



observed



- Different source depths show different signatures.
- Best match between observed and modelled data at 2040 m (reservoir/basement interface).
- Different phases can only be distinguished at larger source-receiver distances (> 1200 m).

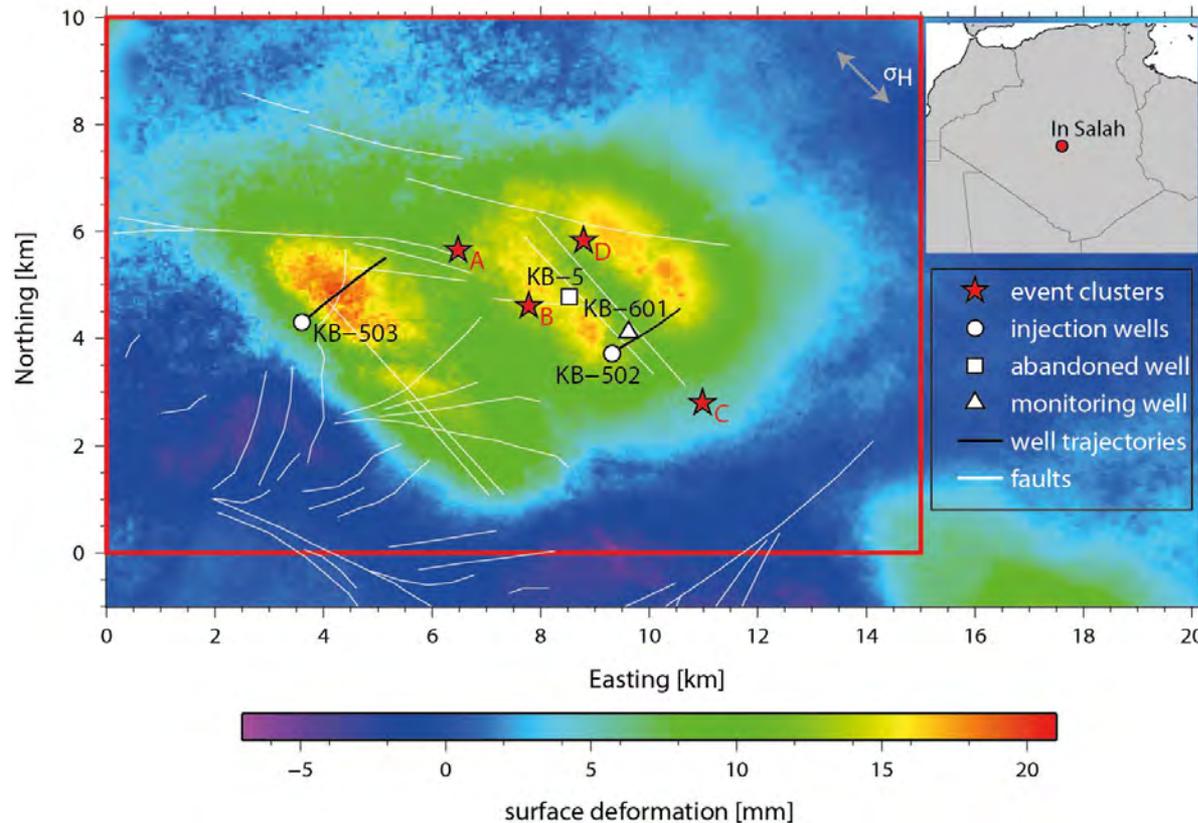


Outline

- Reservoir characterization using microseismicity at the IBDP
 - Waveform cross-correlation and source analysis
 - Full-waveform modelling for hypothesis testing & interpretation
- Constraining event depth with examples from
 - In Salah
 - Oseberg
- Conclusions



Confining microseismic event depth at In Salah

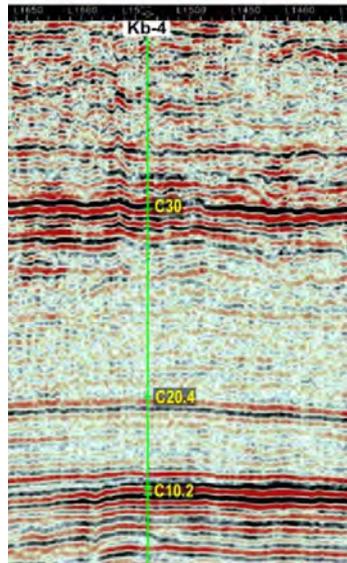
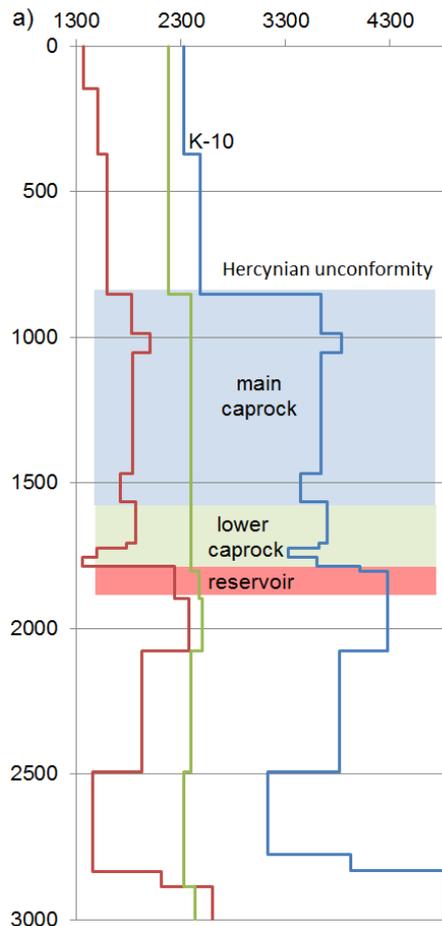


- 4 MT CO₂ injected into sandstone reservoir at 1.9 km depth.
- > 5000 microseismic events detected during injection.
- Events grouped in four clusters but no accurate locations and no depth resolution (only one geophone analysis)

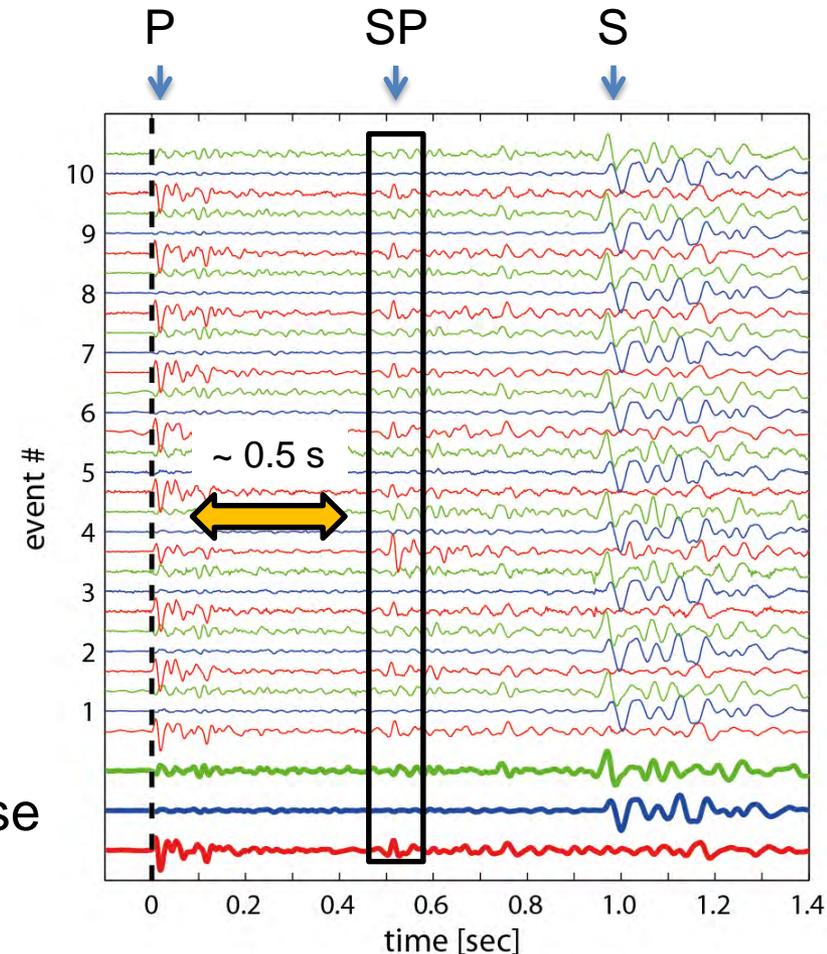
Goertz-Allmann et al. (2014)



Confining microseismic event depth at In Salah



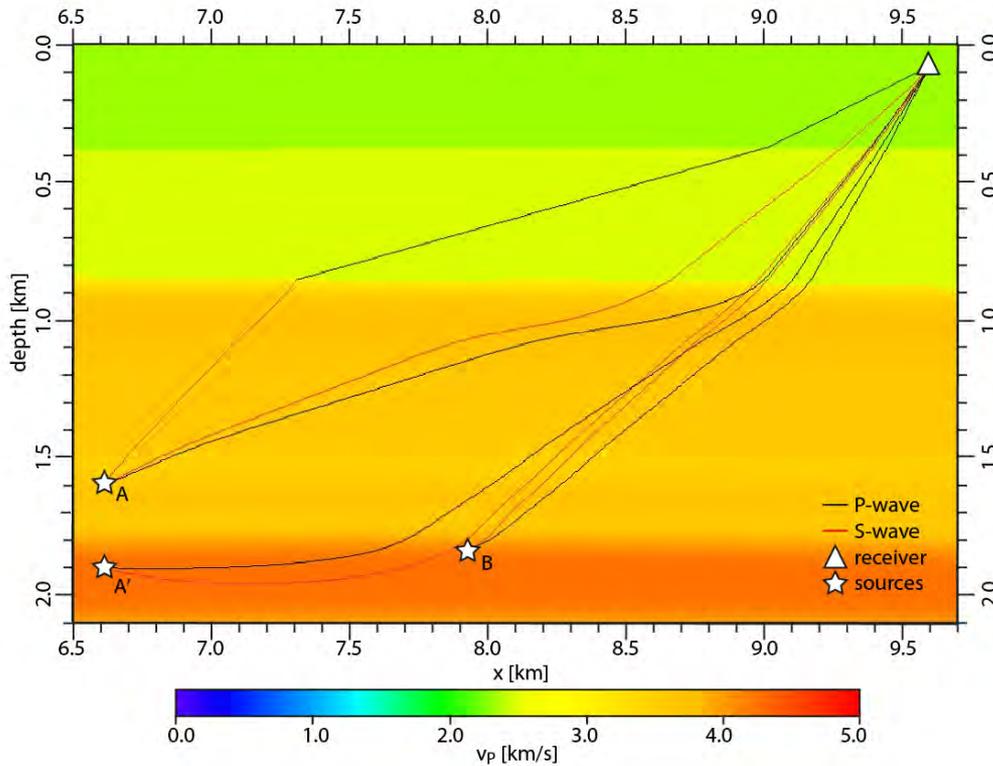
- Additional phase on Z between direct P & S
- S-to-P converted phase at strongest velocity contrast (850 m).



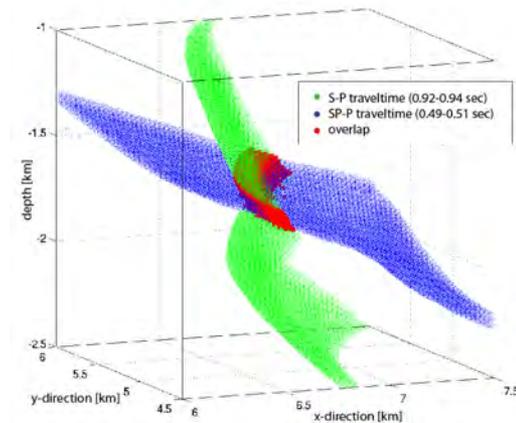
Cluster A

Confining microseismic event depth at In Salah

➤ Use 3D ray tracing to identify converted SP.



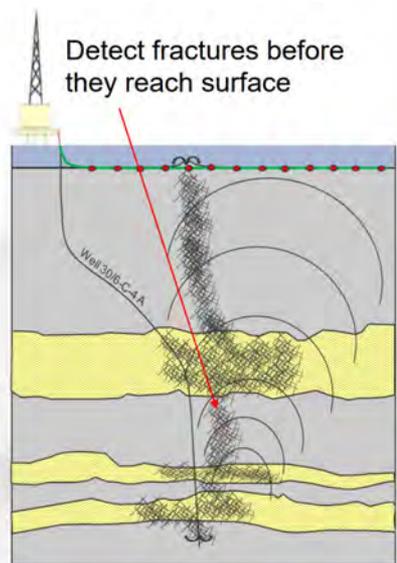
- Waveforms at A and A' have similar S-P traveltimes but converted phase only matches real data at shallower position A.



- Cluster A at about 1.7 km (well above the reservoir but still within lower cap rock).

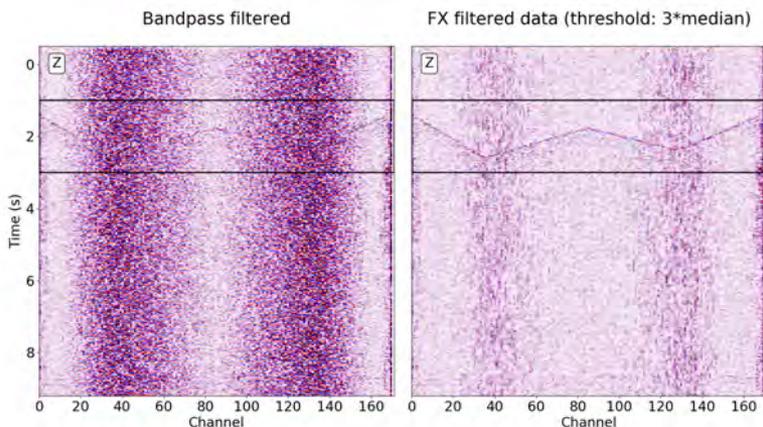


Improving microseismic event depth at Oseberg

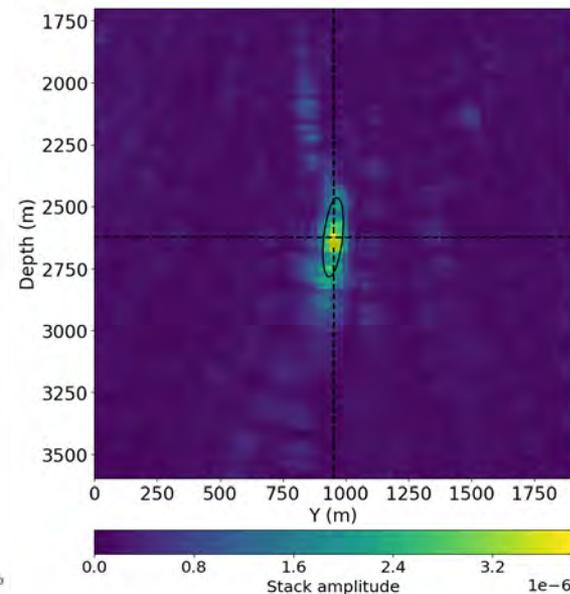
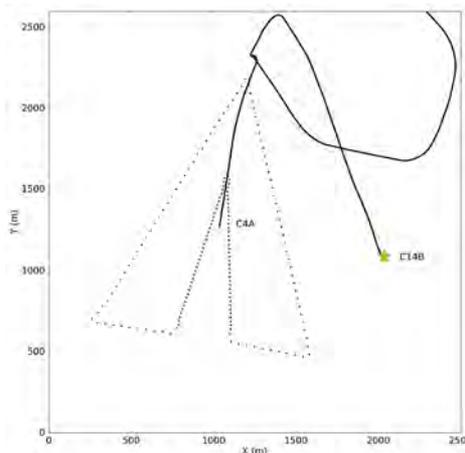


Source: Equinor

- Offshore recordings are often contaminated by a variety of noise sources
- This effects detectability and location precision
- Noise can mask part of a network:
 - Decreases effective aperture, reducing location precision (depth)
- We need smart ways of removing this noise to improve our depth constraints



172 sensors



Bussat et al. 2016

Dando et al. 2016

Conclusions

- During CCS operations: most important is event depth resolution to verify seal integrity and to map reactivated fractures outside of the reservoir.
- Reservoirs are generally thinner than depth uncertainty from standard seismological methods → additional constraints need to be exploited!
- Exploiting information contained in later arrivals / multipathing.
- Requires waveform modelling and ray-tracing for hypothesis testing and confirmation.
- Advanced noise removal techniques may be necessary in offshore operations.
- Decatur:
 - Connection between reservoir and basement
 - Confirm a source at the reservoir/basement interface
- In Salah:
 - Information on caprock integrity
 - Despite very inadequate network coverage
- Oseberg:
 - Distinguish in zone and out-off zone events



Thank you for your attention!

Contact: Bettina Goertz-Allmann

Email: bettina@norsar.no

Acknowledgements

The partners in the In Salah CO₂ Storage Joint Industry Project - BP, Equinor (formerly Statoil), and Sonatrach.

Part of this work on Decatur was supported as part of the Center of Geological Storage of CO₂, an Energy Frontier Research Center funded by the U.S. Department of Energy, Office of Science.

Data for this project were provided, in part, by work supported by the U.S. Department of Energy under award number DE-FC26-05NT42588 and the Illinois Department of Commerce and Economic Opportunity.

We are grateful for support through the Climit program of GASSNOVA project No. 616065.

Data provided by the Midwest Geological Sequestration Consortium (MGSC), funded through the U.S. DoE National Energy Technology Laboratory (NETL) and the State of Illinois.



8.03.2019

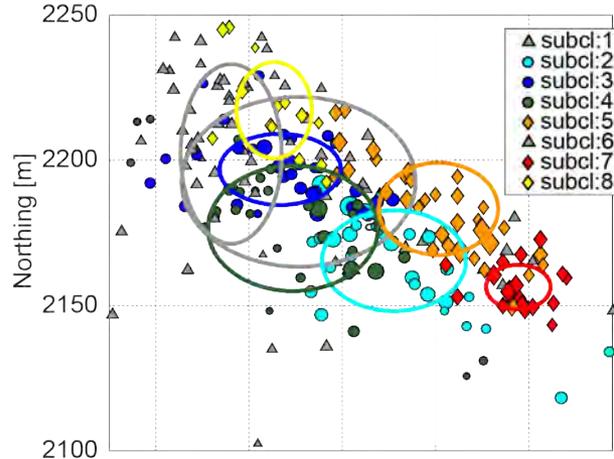
Schatzalp induced seismicity workshop

NORSAR

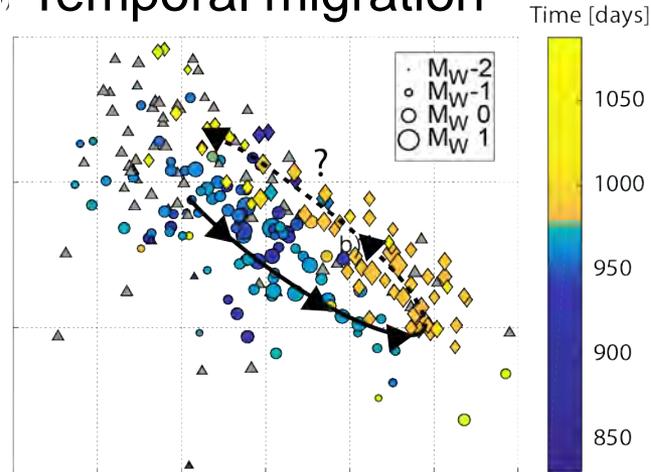
Microseismic event characterization

Cluster A

a) Spatial distribution



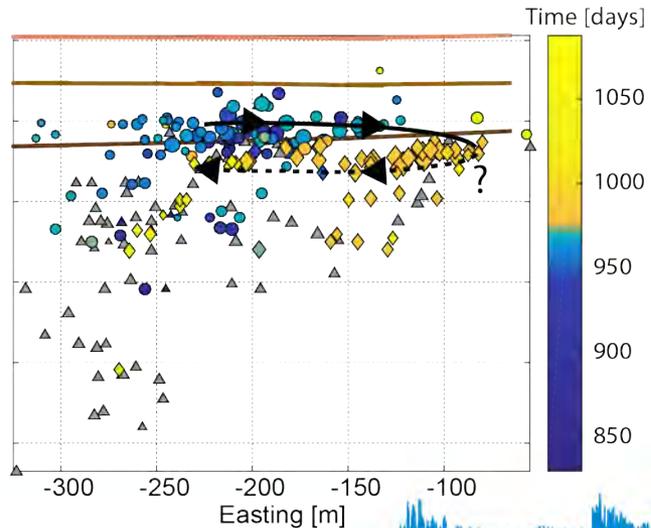
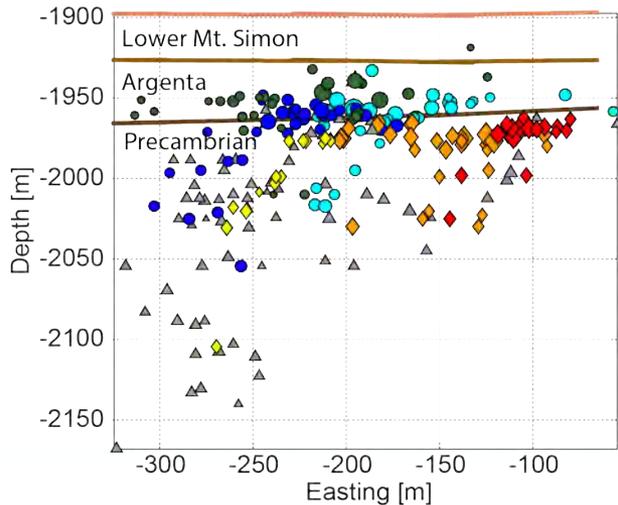
b) Temporal migration



- Separate events occurring within different layers:

Cold = reservoir

Warm = basement



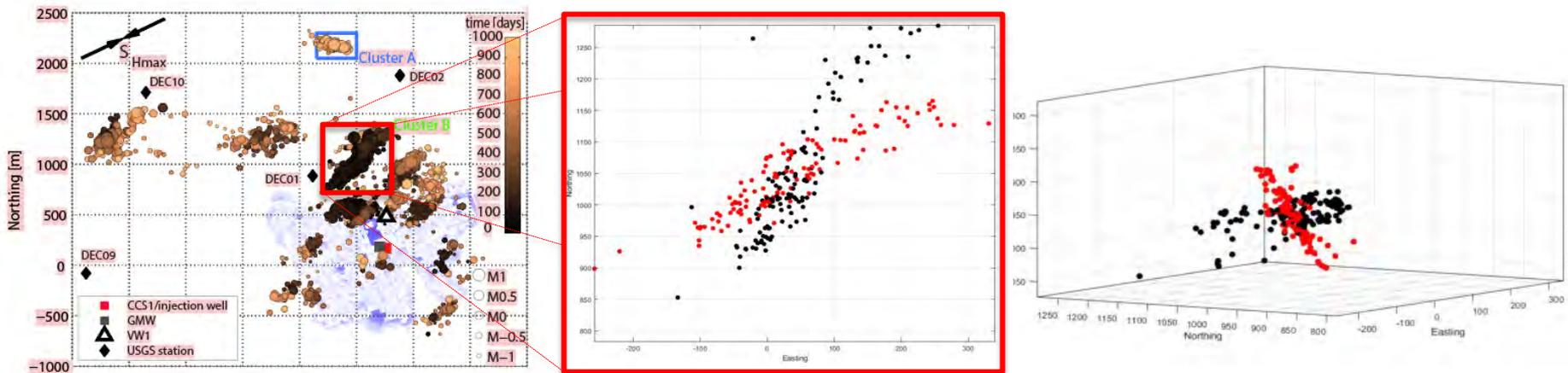
- Migration of events from the reservoir into the basement over the course of 100-200 days.

Goertz-Allmann et al. (2017), JGR



Relative event locations

Preliminary results of improved relative event locations by developing a modified relocation method.



- Accurate event locations are necessary for any kind of interpretation
- Change of cluster orientation
- Planar feature

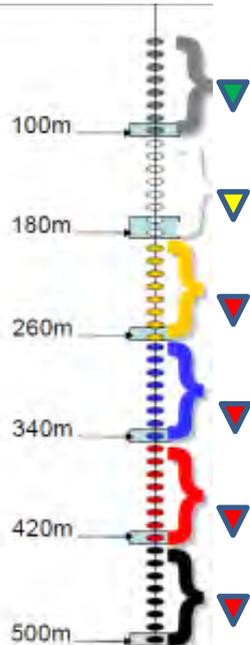
→ **Fracture?**

- Old event locations
- Relocated events



Microseismic array at KB-601

2009-2011

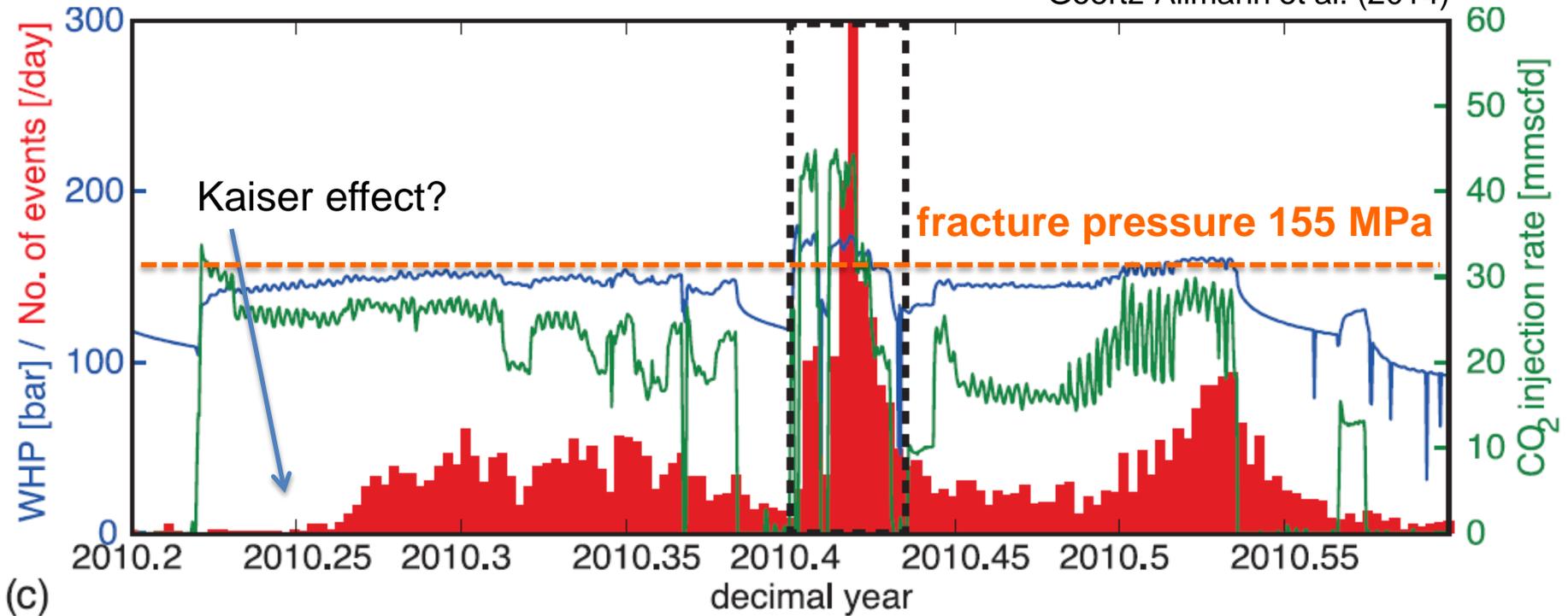


- Downhole array of 48 3C geophones between 30-500 m depth
- 6 geophones were connected to 3 digitizers
- GPS timing problems and strong electronic noise
- **Only uppermost geophone provided reliable data**



Comparison of events and injection data

Goertz-Allmann et al. (2014)

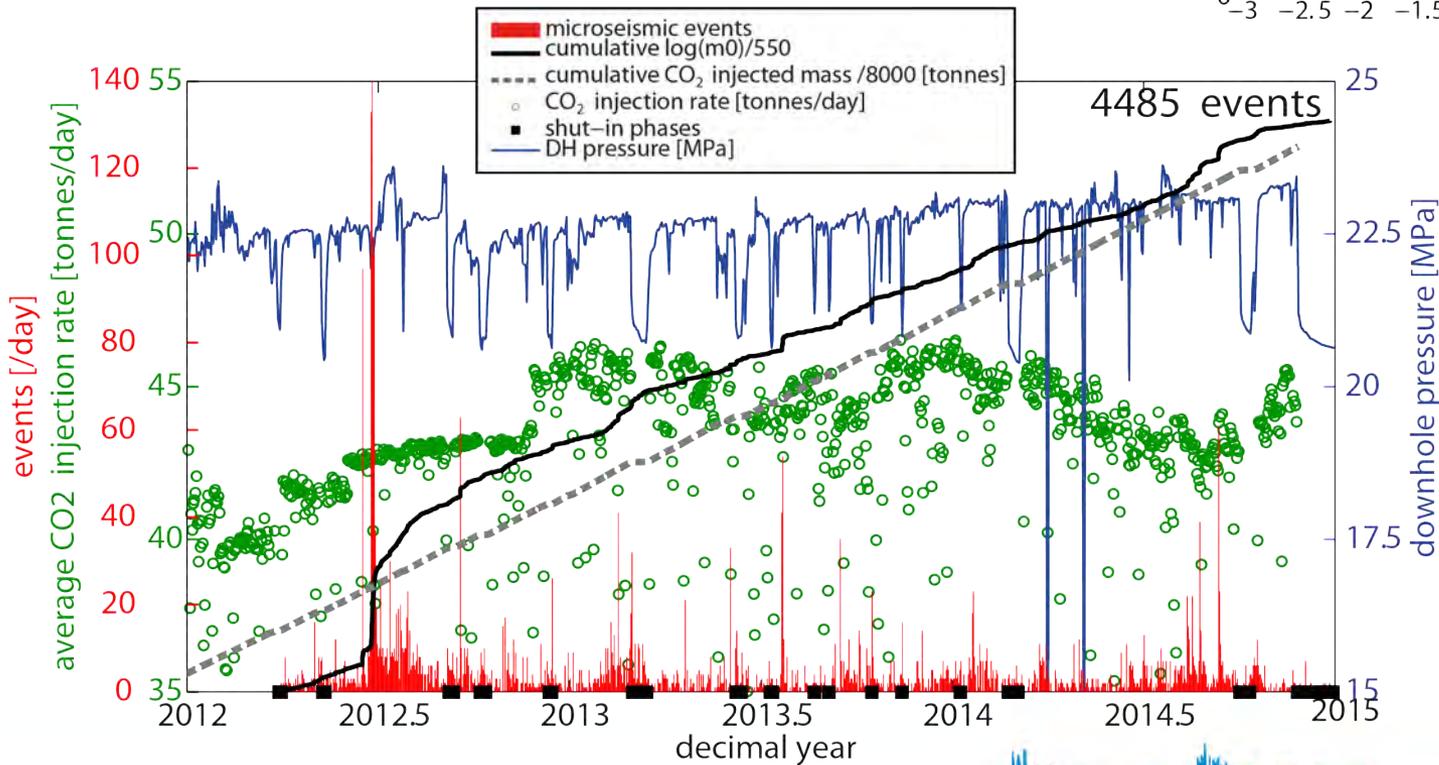
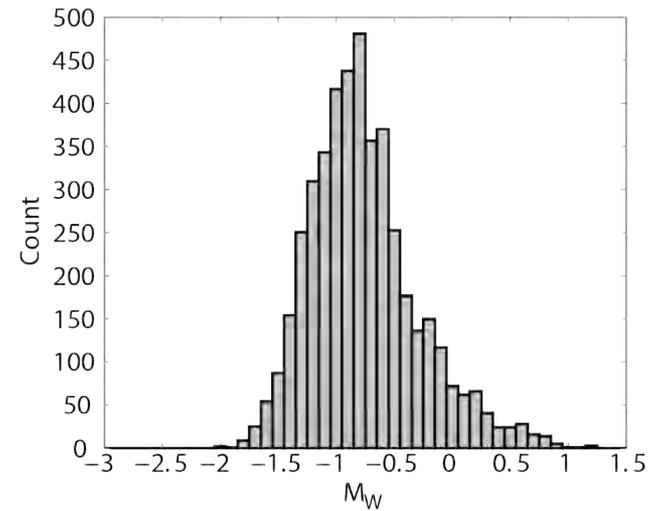


- High correlation between occurrence of microseismic events and injection rate
- Periods of **matrix injection** and **fracture injection**



The Decatur CCS site

- Most events with $M_w < 0$.
- Injection at very low pressure (< 1 MPa)
- No obvious correlation with plume migration – events far from the injection



Ray-tracing using QSEIS

